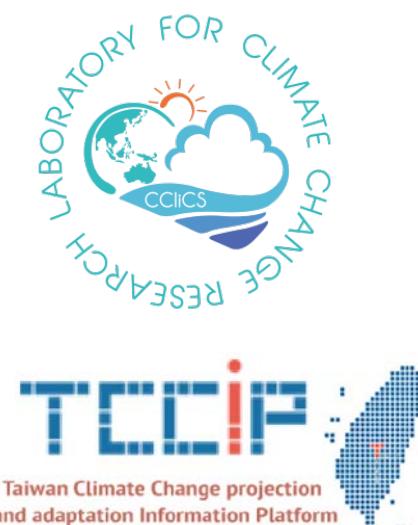


East Asian Climate Variability and Change

(A Brief Review)

Huang-Hsiung Hsu

Laboratory for Climate Change Research
Research Center for Environmental Changes
Academia Sinica

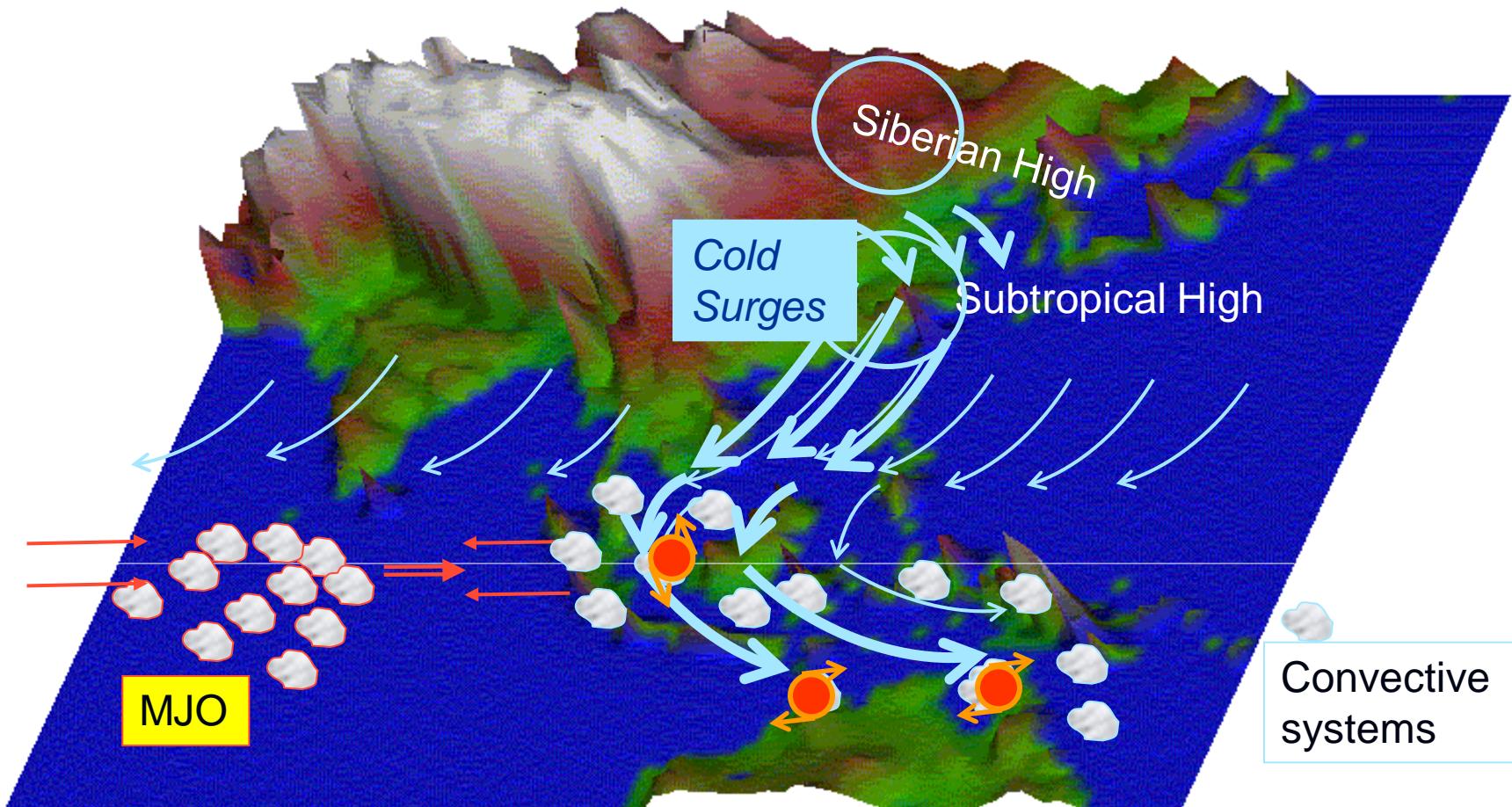


Consortium for Climate Change Study
(CCICS/MOST)
Taiwan Climate Change Projection and
Information Platform (TCCIP/MOST)

DSWS, Tsukuba, October 2017



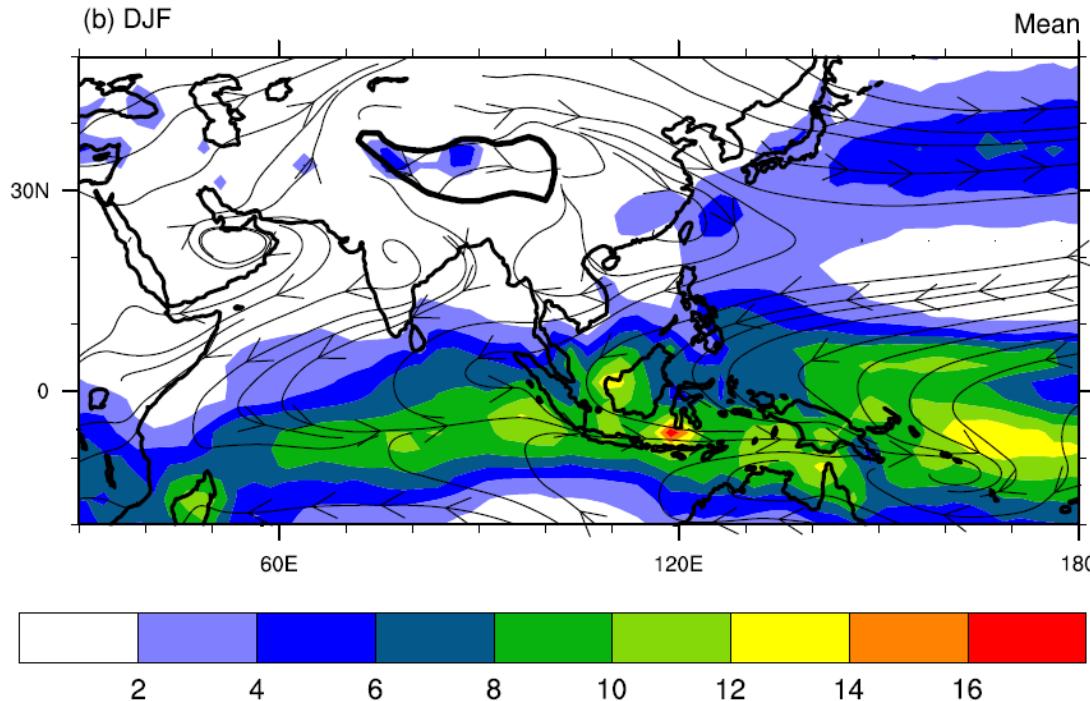
East Asian Winter Monsoon



courtesy of C.-P. Chang

Winter Climate Variability

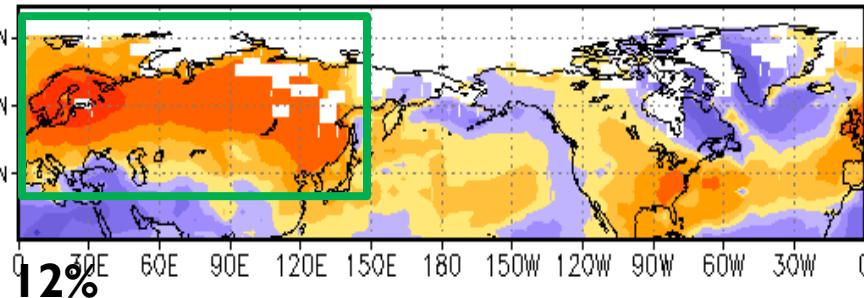
Need to understand as part of global system



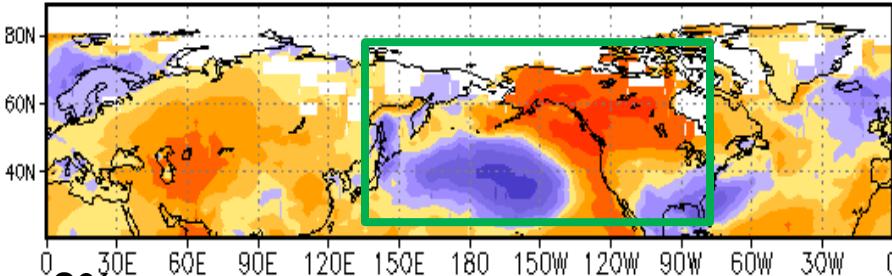
Regime Shift and Change in Dominant Modes Mechanisms?

CRU+Hadisst temperature 1941-2002 (JFM)

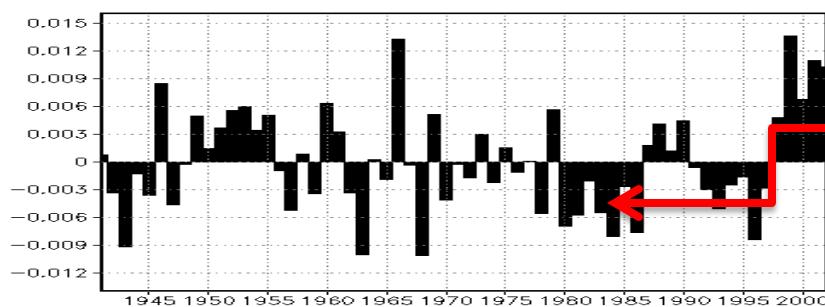
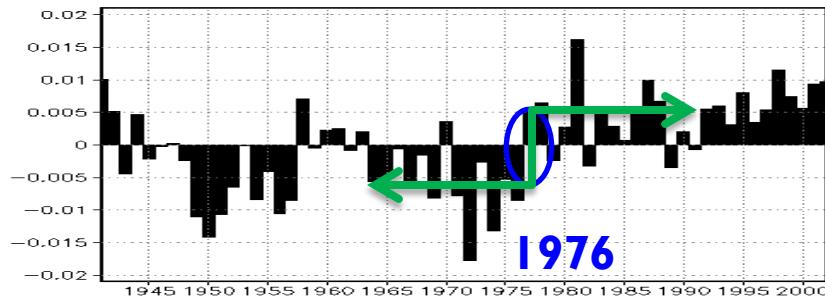
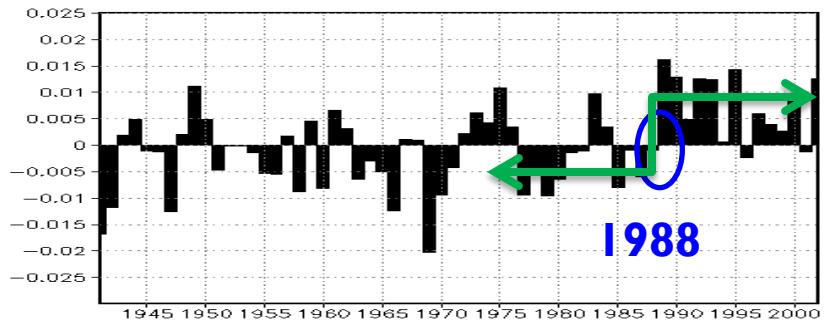
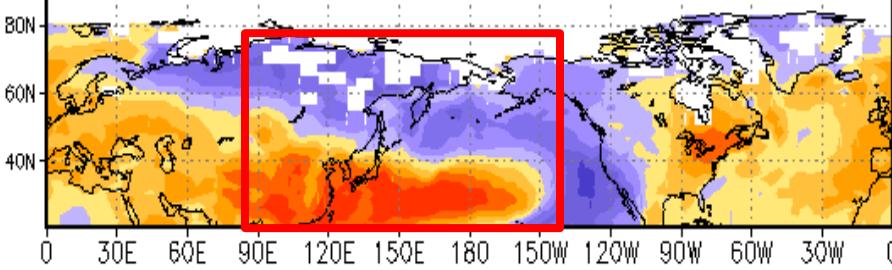
15%



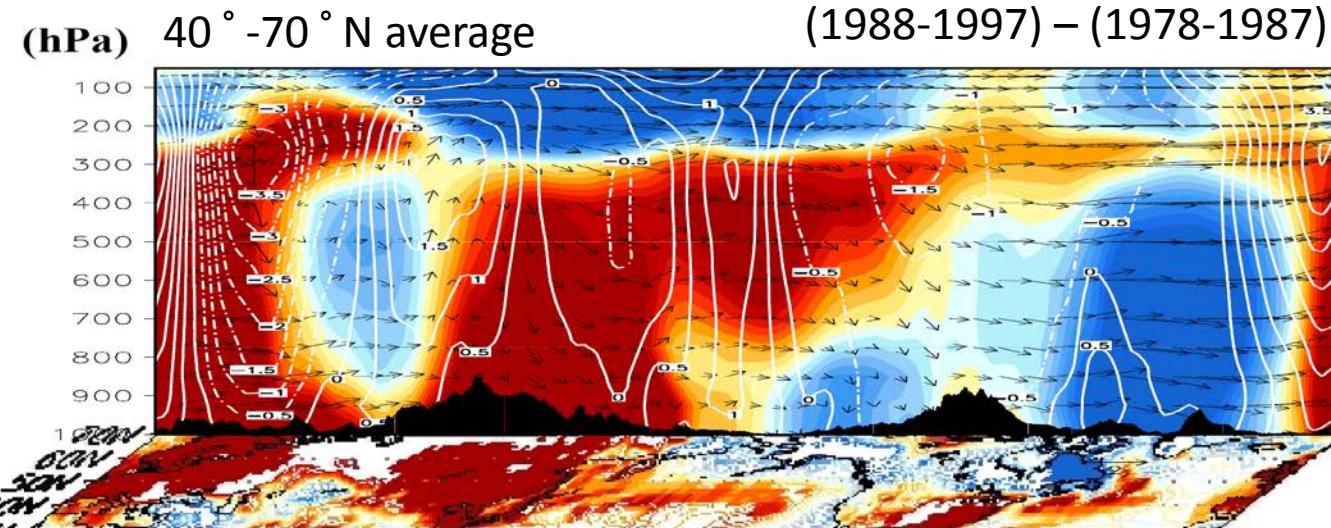
12%



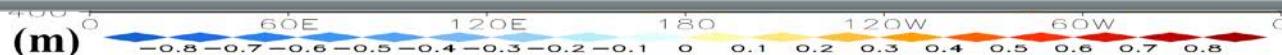
8%



Late 1980s Warming in East Asia (associated with NAO/AO)



- Synchronous regime shift occurred in the whole troposphere, on and under ocean surface. What cause this?
- A dominant hemispheric (global) pattern, which has not been well understood?
- Mutual enhancement through ocean-atmosphere teleconnection?



shaded: temperature

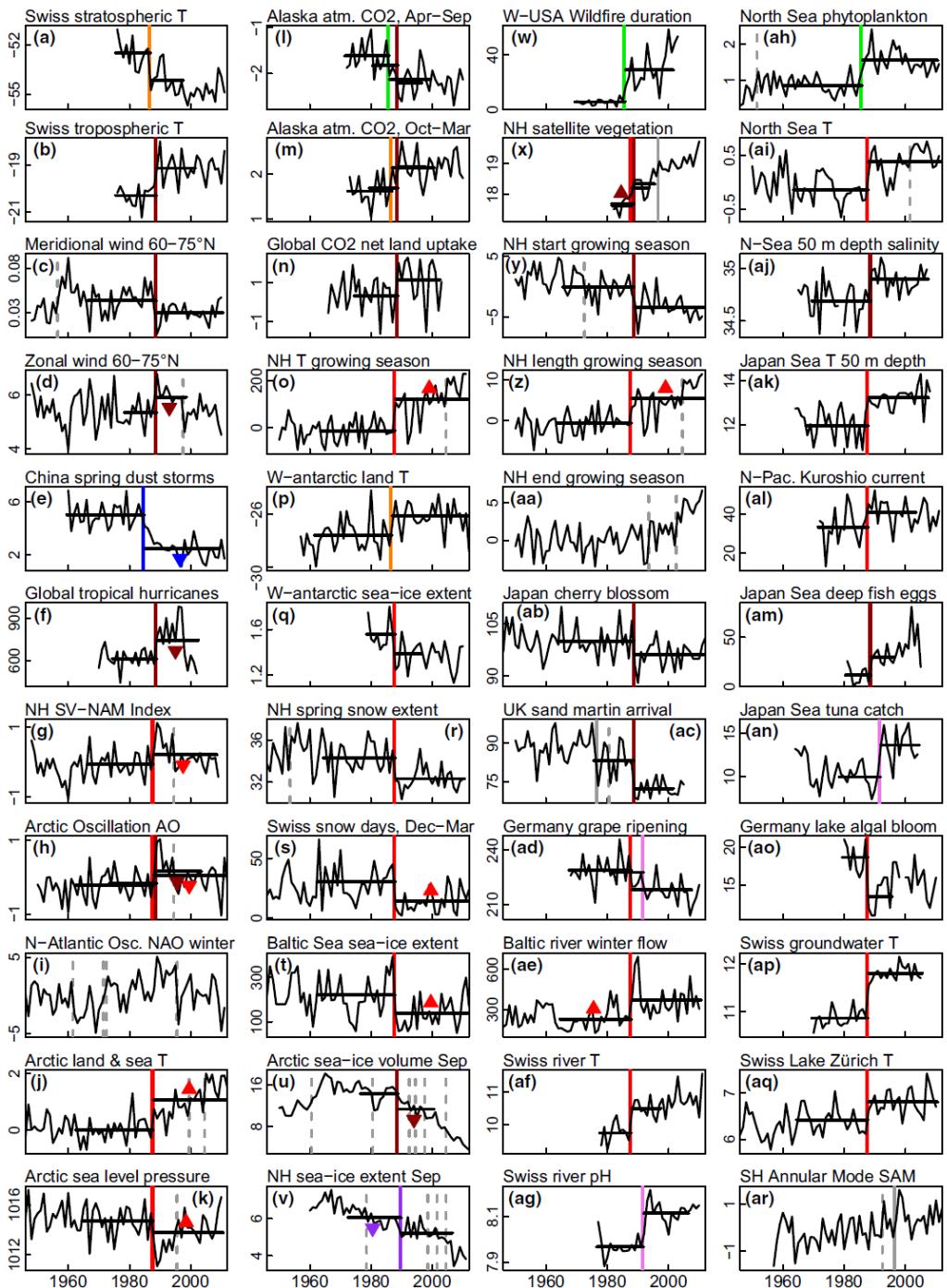
contour: v-wind

arrow: u-w wind

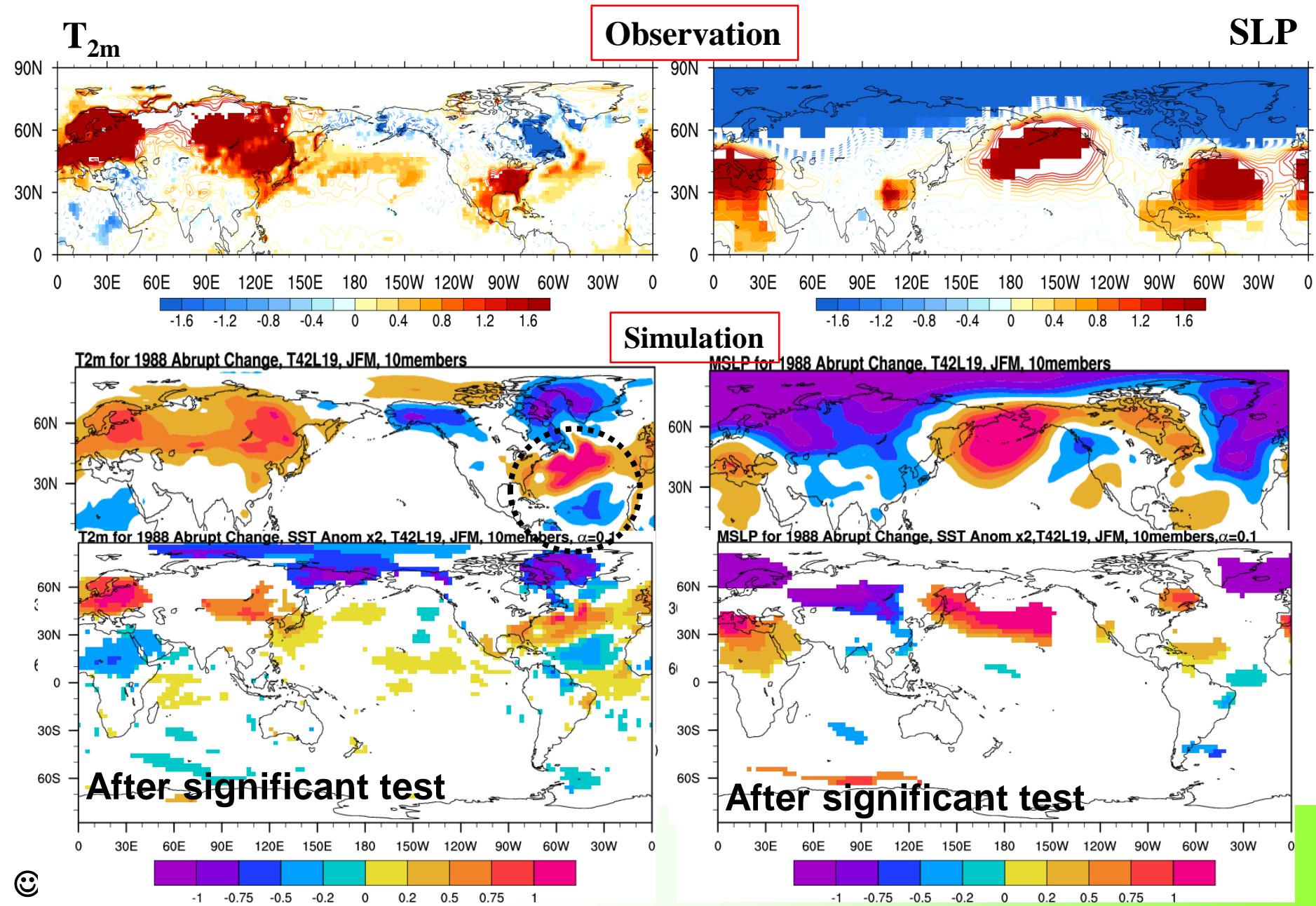
Global impacts of the 1980s regime shift

Reid et al., 2015
Global Change Biology,
doi: 10.1111/gcb.13106

- China dust storm (decreased)
- Global Tropical cyclone (increased)
- CO₂ land uptake (increased)
- Antarctic sea ice/Baltic sea ice (d)
- NH spring snow (decreased)
- Starting growing season (NH)
- Japan Sea T 50m (higher)
- Japan cherry blossom (earlier)
- Japan Sea tuna catch (more)
- Kuroshio current (stronger)
- German grape ripening (earlier)
- North Sea phytoplankton (increased)
- AO (negative → positive)
- ...



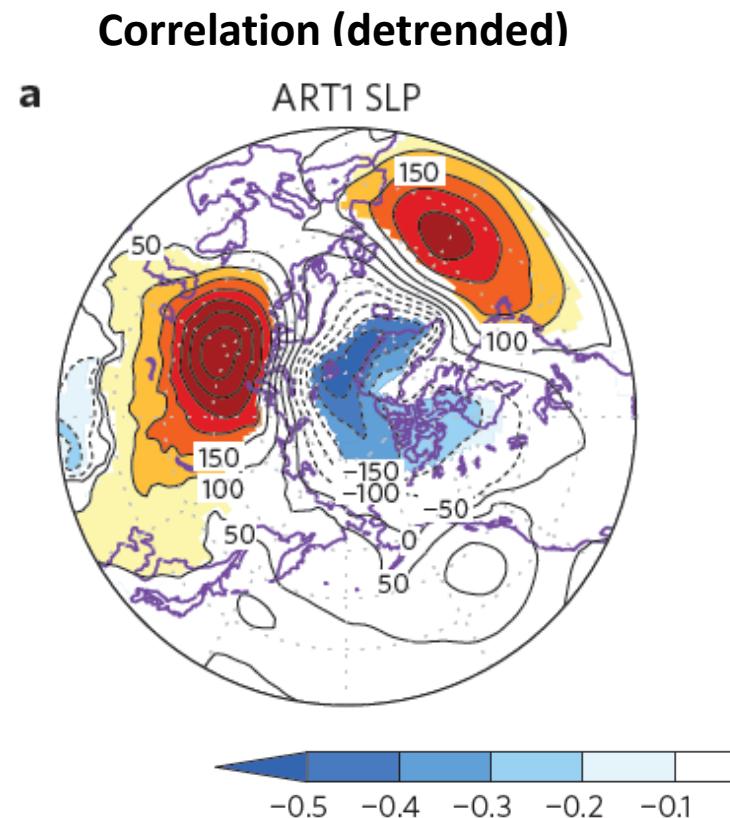
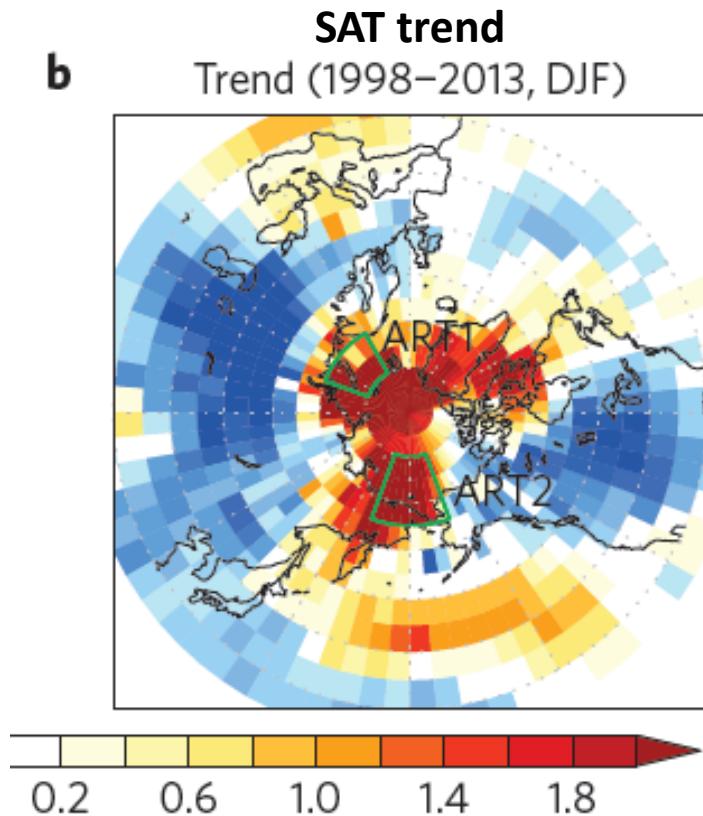
EA warming can be simulated by SSTA forcing in the North Atlantic



East Asia winter and Barents-Kara Sea SAT after 1979/80 (interannual time scale)

“...severe winters (1979/80 – 2013/14) across East Asia are associated with anomalous warmth in the Barents–Kara Sea region ...”

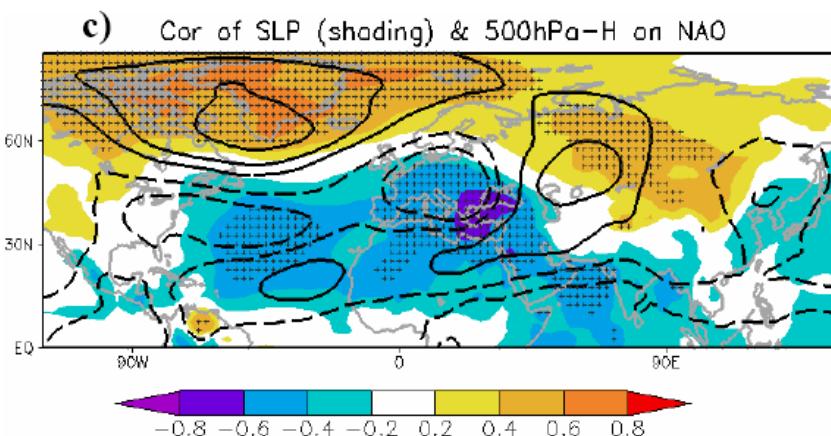
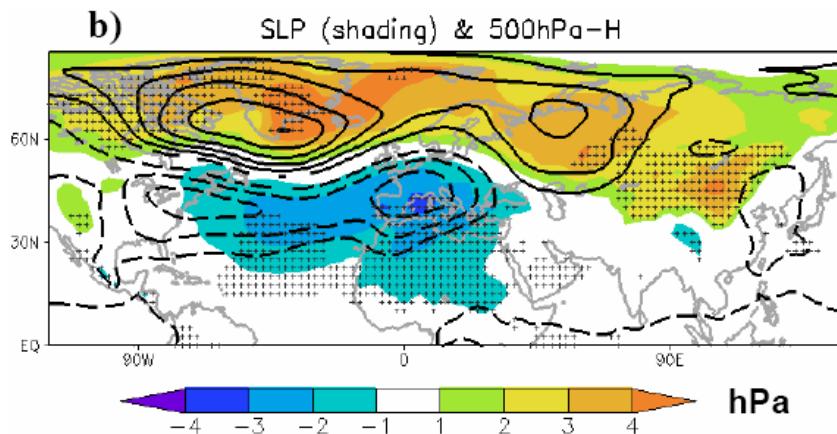
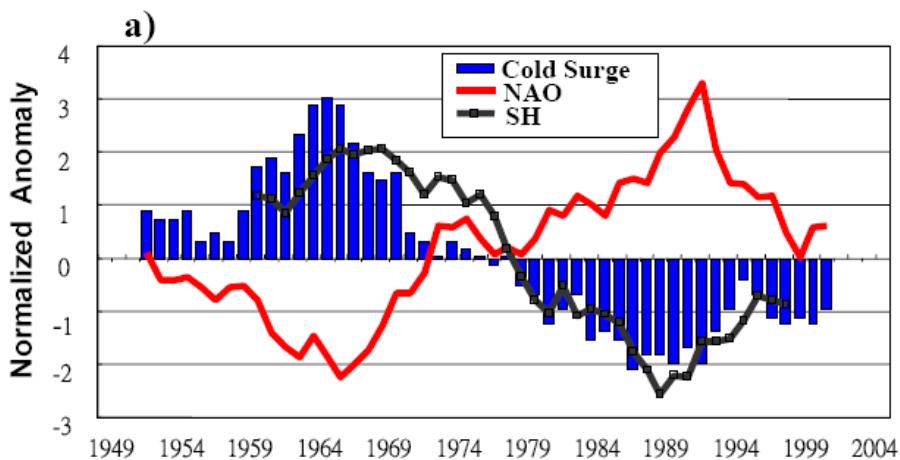
(Kug et al. 2015, Nature Geoscience; similar in Mori et al. 2015; Kim and Son, 2016; Petoukov and Semenov 2010, ...)



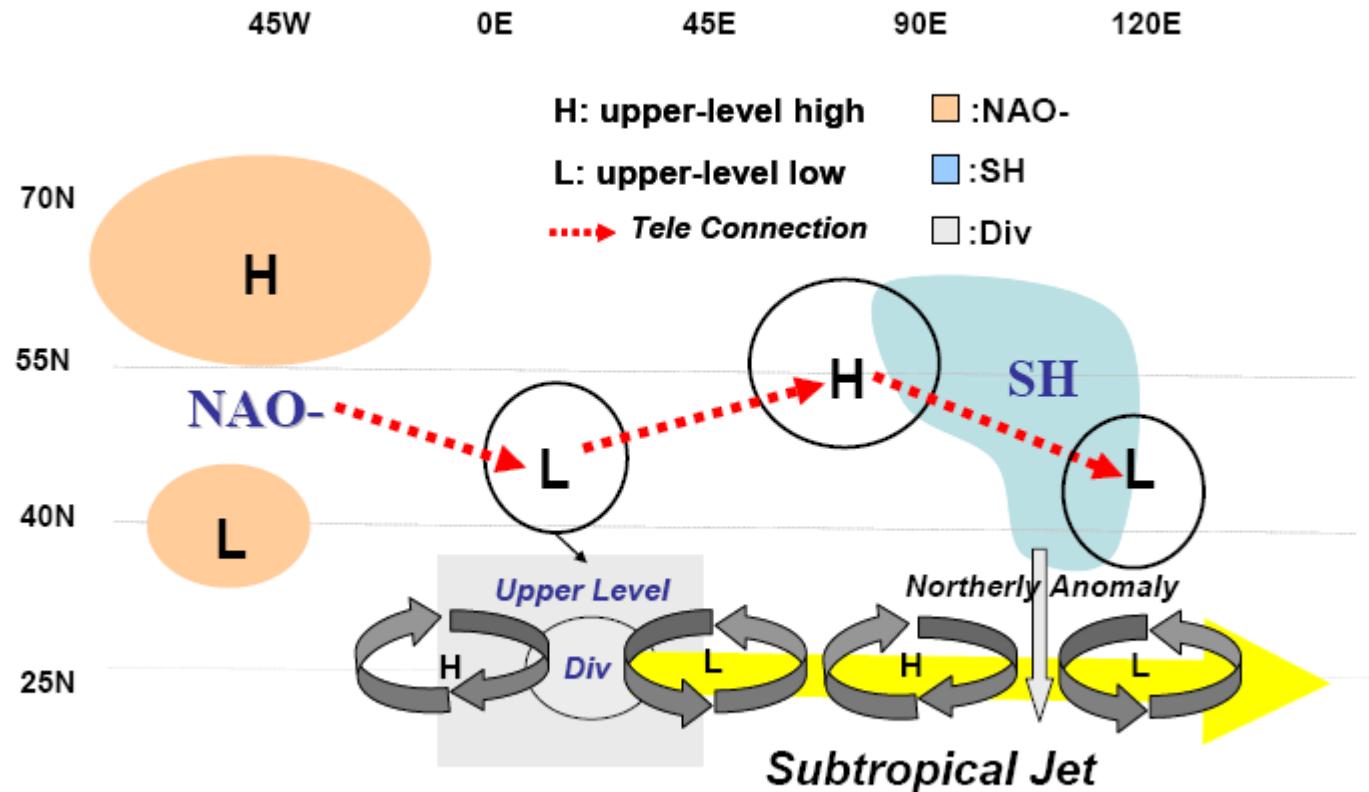
Positive Decadal Correlation: NAO and Cold Surge Frequency in Taiwan (correlation ~ -0.9)

(Hong, Hsu, Chia, and Wu, 2008, GRL)

- Similar pattern in interannual and interdecadal time scale
- Cause and Effect?
 - NAO/AO like perturbation
 - Barents-Kara Sea SSTs
 - others



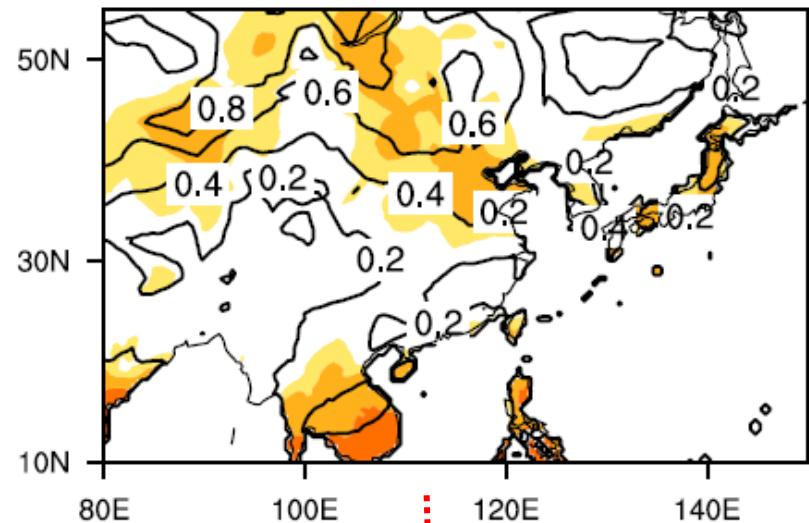
- Downstream impacts (of North Atlantic/Europe) on East Asian climate/weather through Rossby wave-like propagation (waveguides)
- How will this change in the warming future?



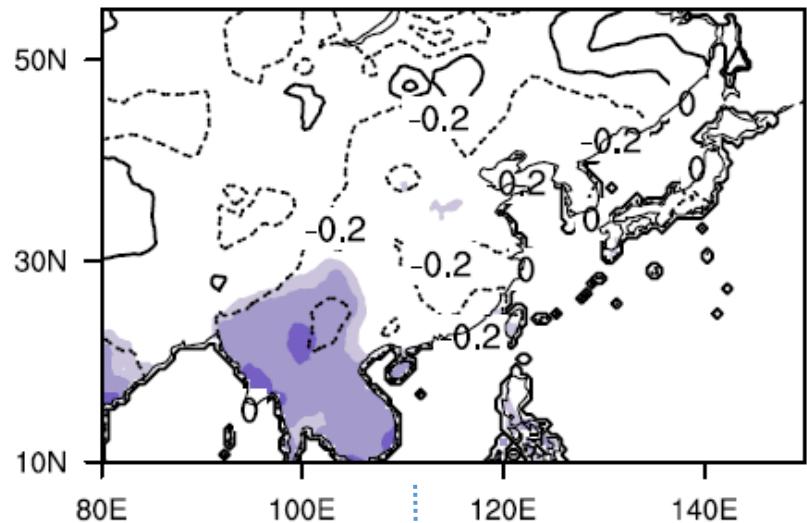
Combined effect of ENSO and AMO (Hao and He, 2017, J. Climate)

SAT & ENSO-like

a) CRU & warm ENSO-like



b) CRU & cold ENSO-like



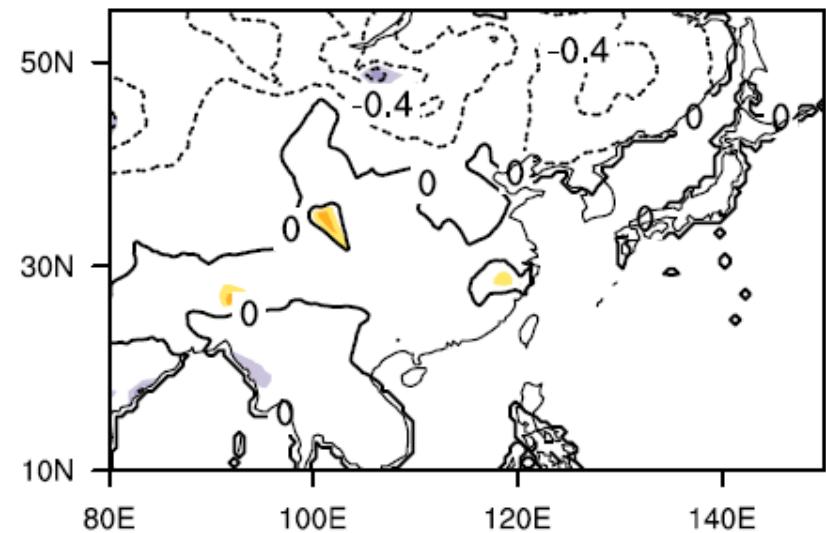
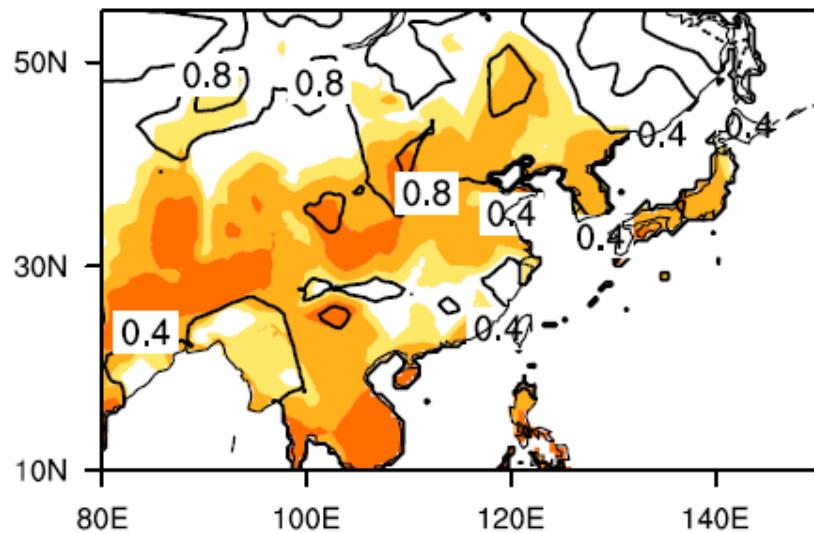
enhanced

a) warm AMO + warm ENSO-like

CRU SAT

cancelled

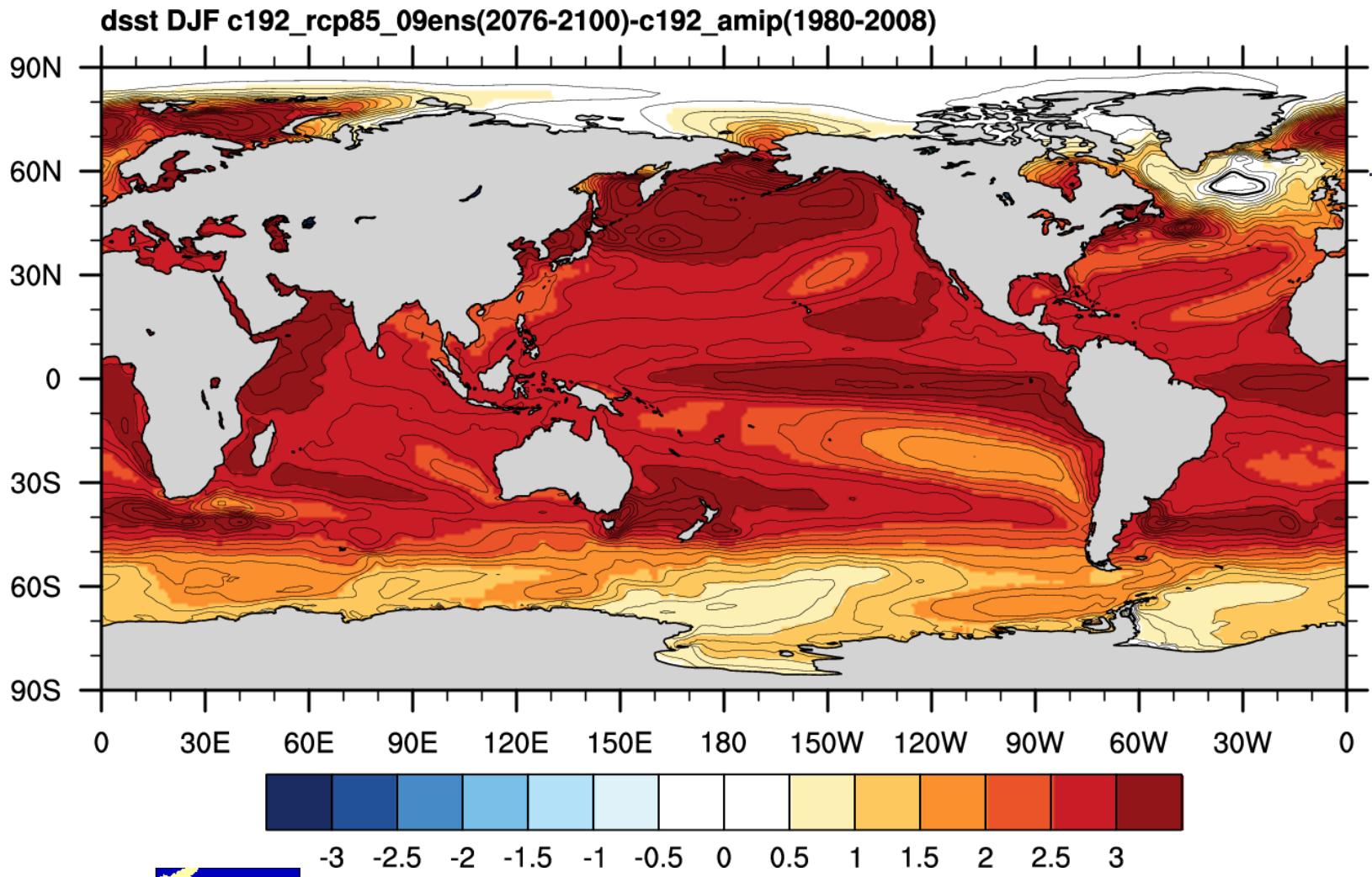
b) warm AMO + cold ENSO-like



Future Change Projection

HiRAM/GFDL: 1979-2008, 2075-2100 (RCP8.5)

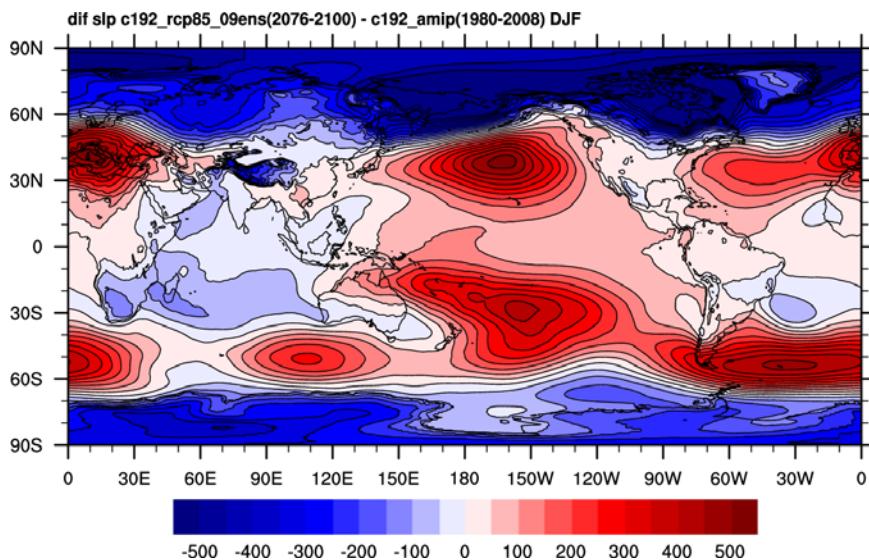
Same ensemble SSTa used in MRI-AGCM20



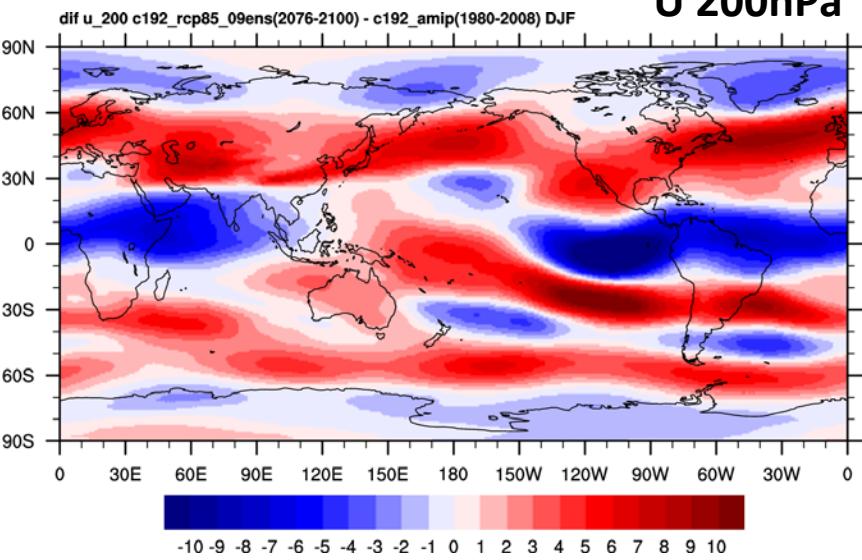
Projected Future Changes under RCP8.5

(using 50/25-km HiRAM forced by ensemble SST changes)

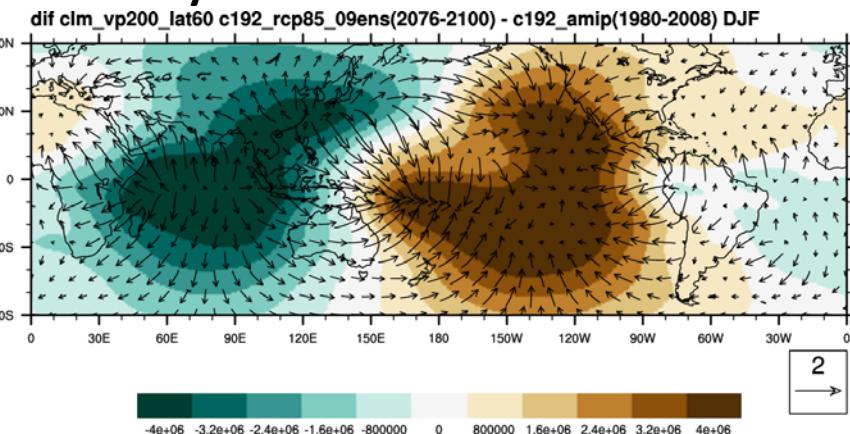
SLP



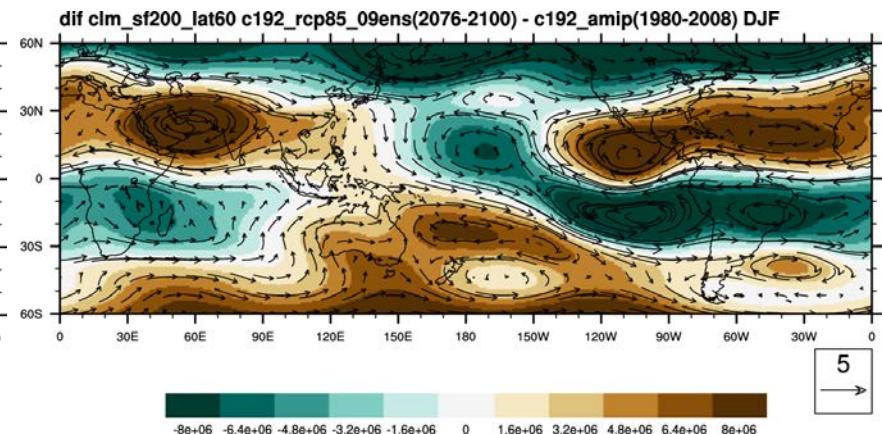
U 200hPa



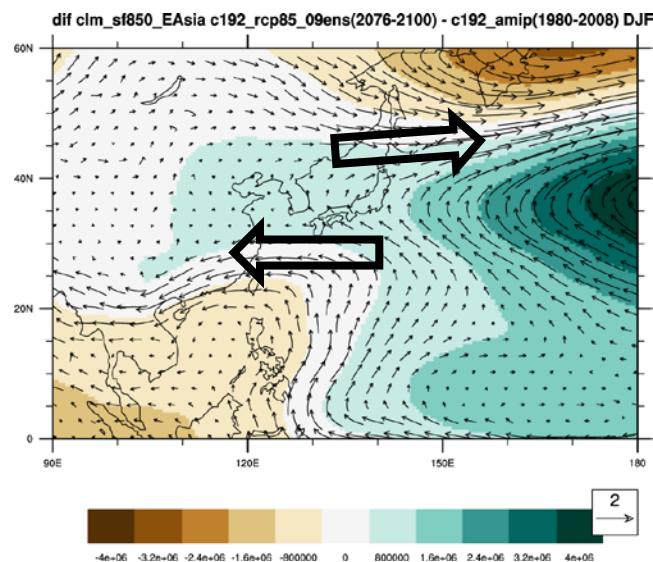
Velocity Potential 200 hPa



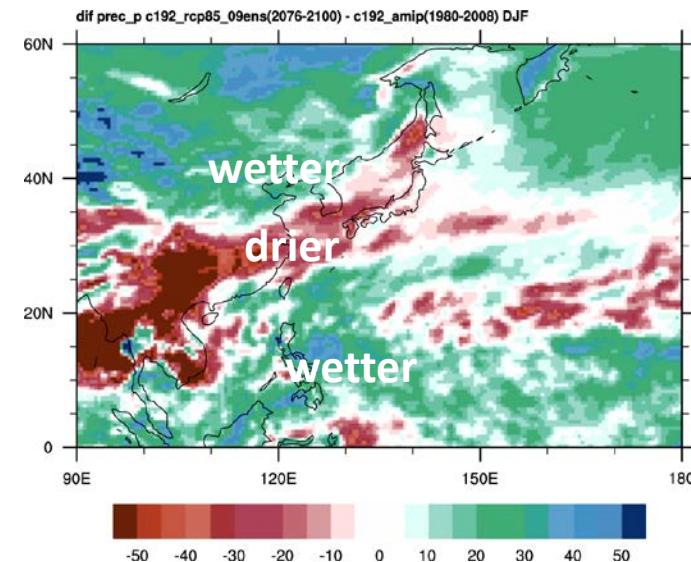
Streamfunction 200 hPa



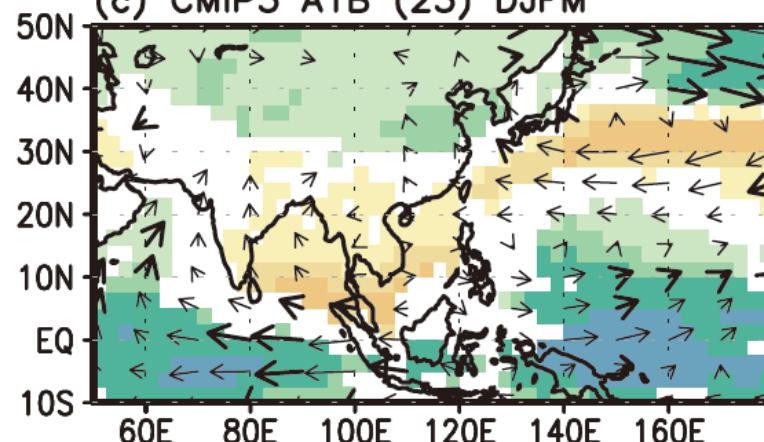
850 hPa Stream Function



Precipitation

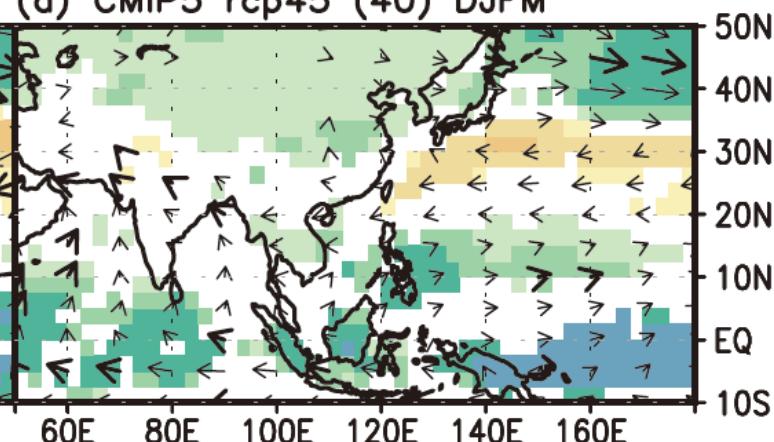


(c) CMIP3 A1B (23) DJFM



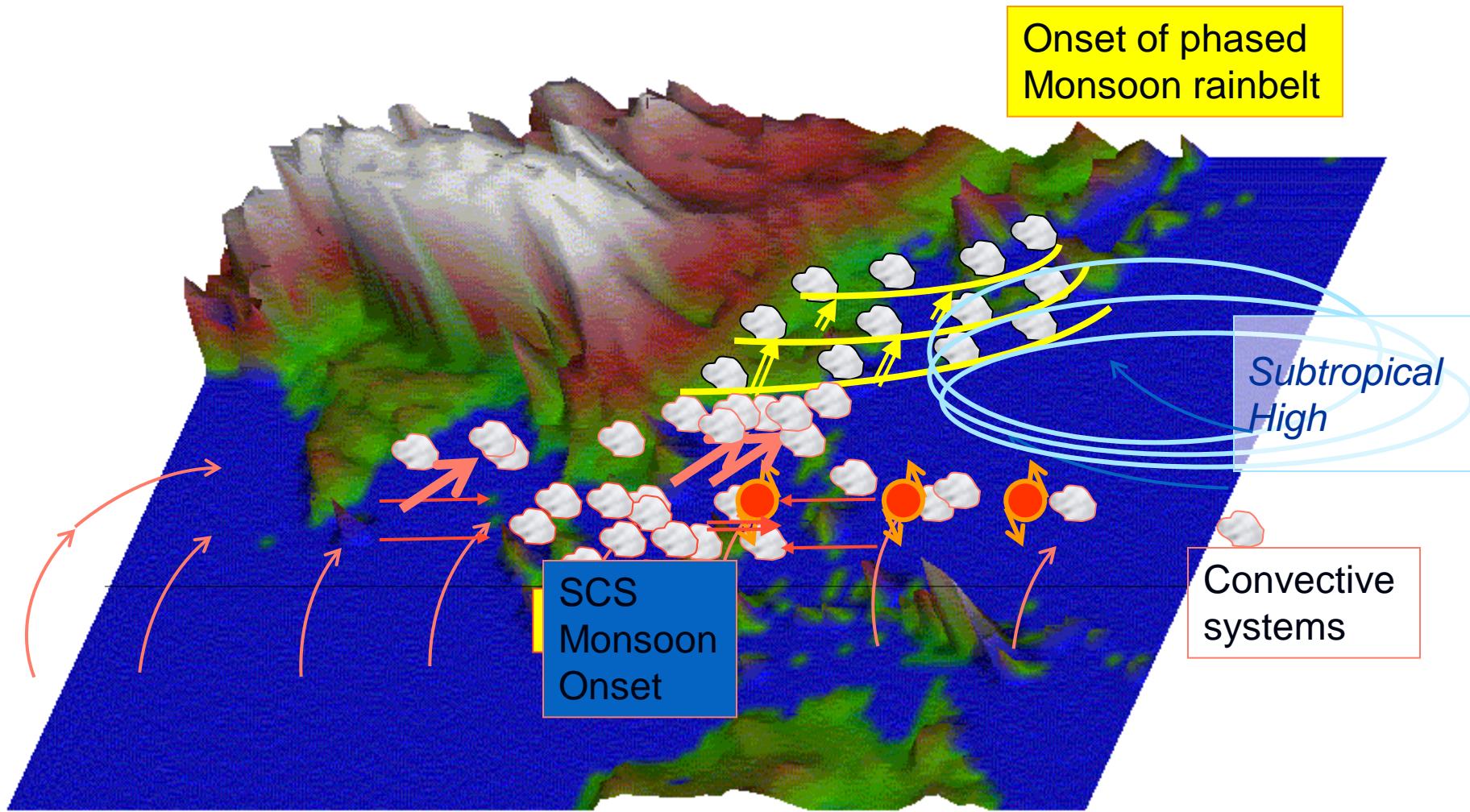
\rightarrow
2 [m/s]

(d) CMIP5 rcp45 (40) DJFM



shade: [mm/day/K]

East Asian Summer Monsoon

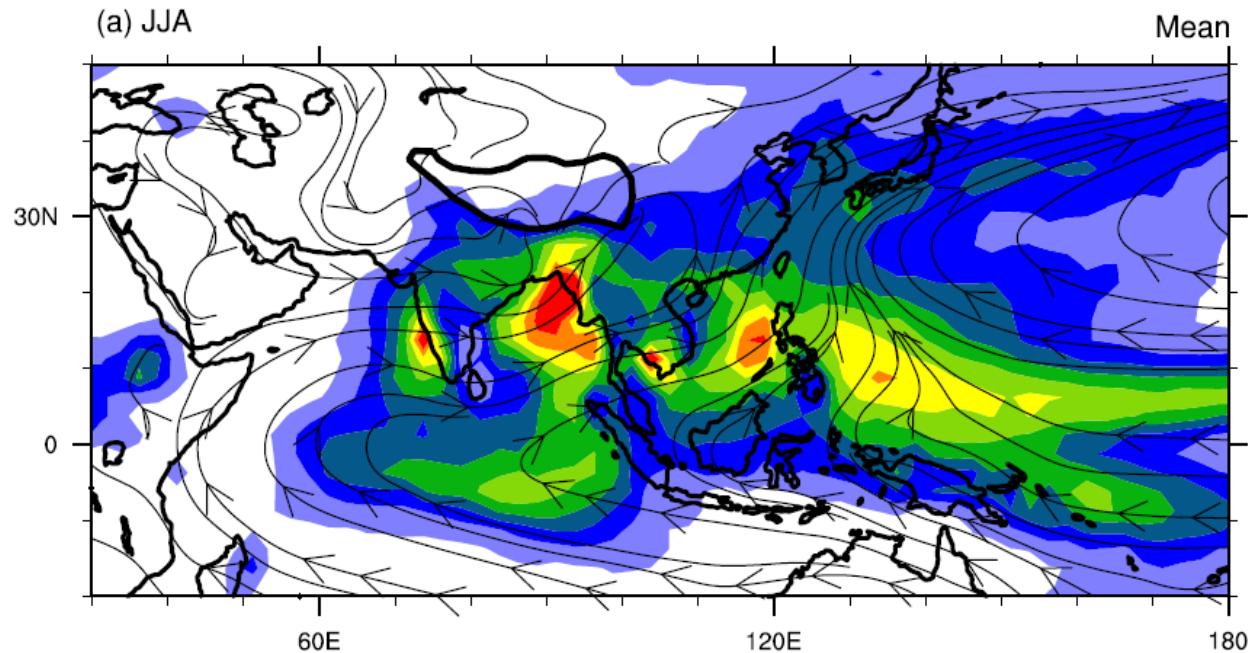


courtesy of C.-P. Chang

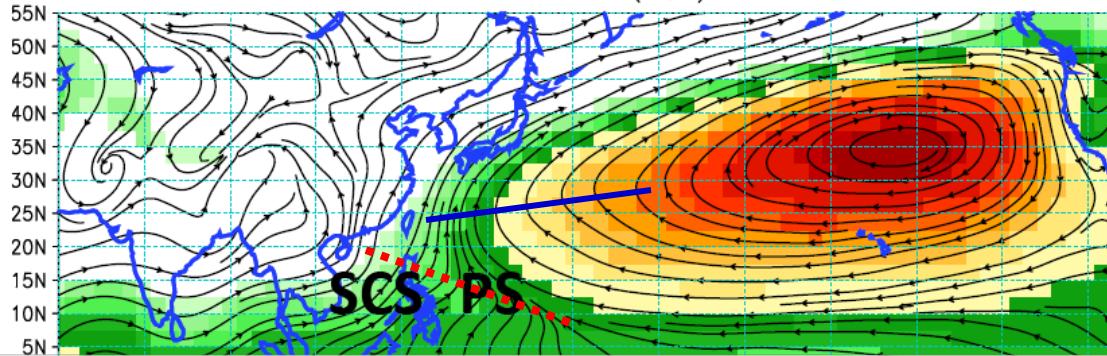
Summer Climate Variability

Multiscale Nature; Convection-Circulation coupling

ERA40 U&V (850 hPa) and CMAP Rainfall 1979-2001

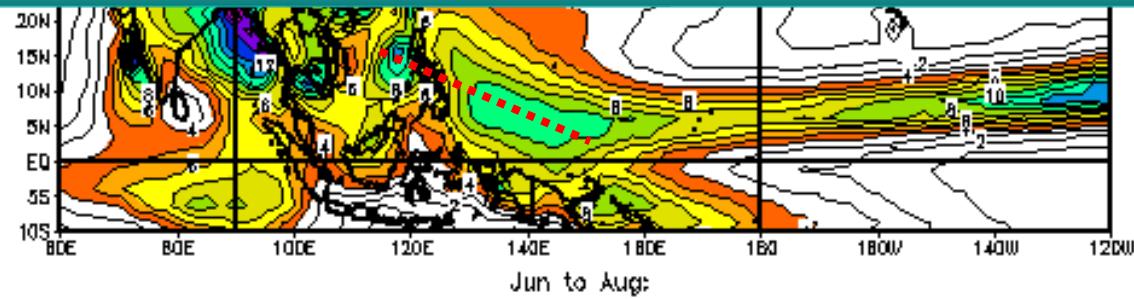


9804JJA 850mb (U,V) & HGT



One of key factors affecting weather and climate in the tropical Western North Pacific -

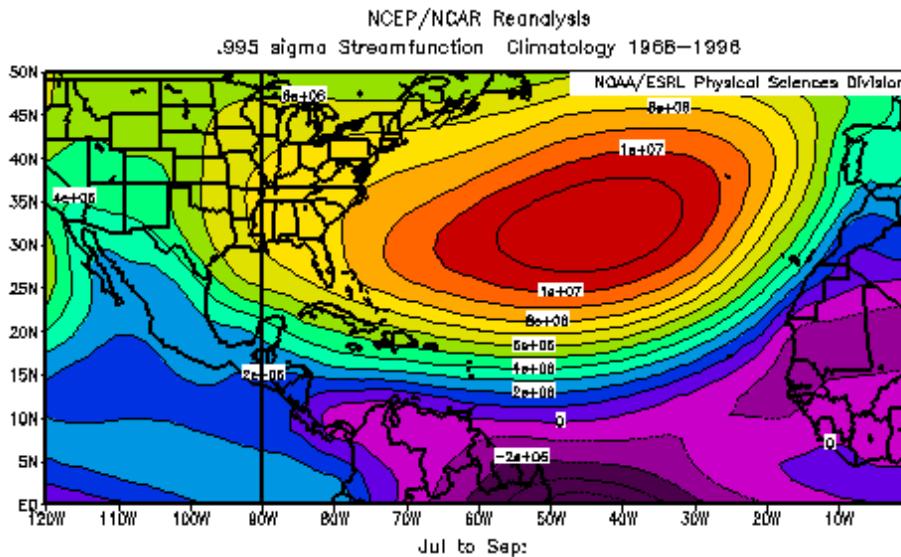
Fluctuation of anticyclonic ridge (ACR) and monsoon trough (MT)



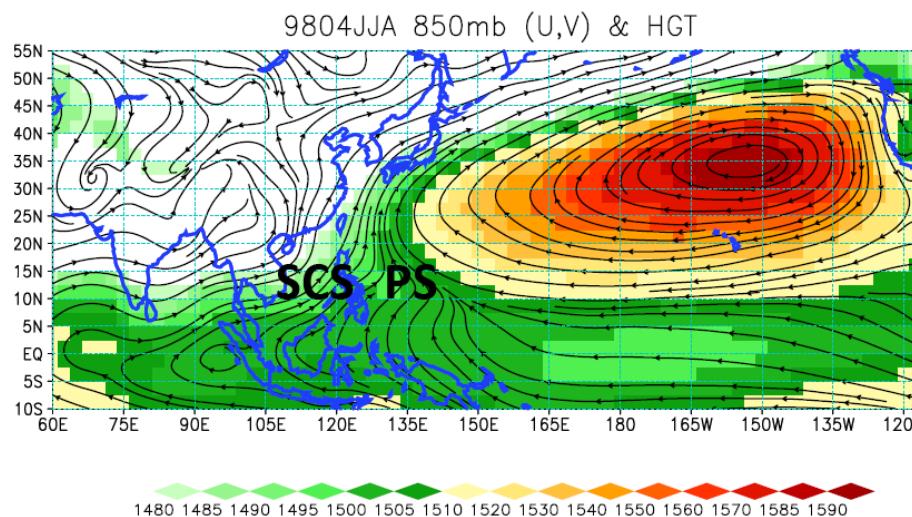
Interesting Contrast bet. Mean Circulation in North Atlantic/North Amer. and WNP/Asia

- No monsoon trough in the NA/NA region

North Atlantic/
North American



Western North
Pacific

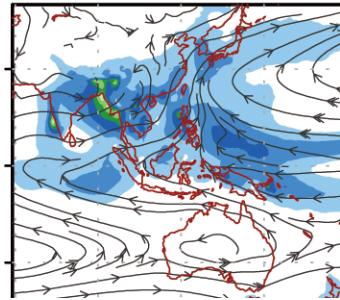




Topographic effect enhances summer monsoon flow (trough) and precipitation

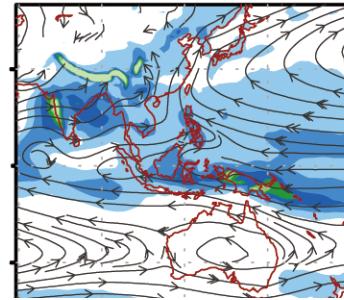
observation

OBS(prec,uv850 JJA)



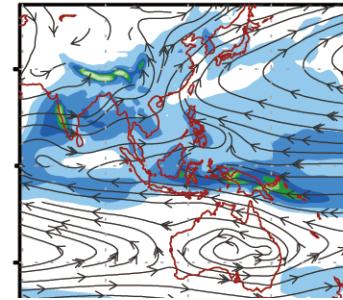
TP12

TP12(prec,uv850 JJA)



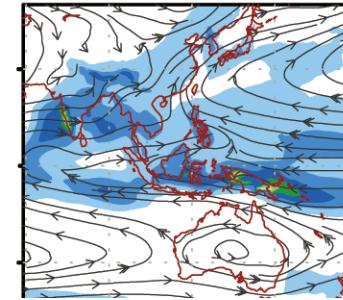
TP07

TP07(prec,uv850 JJA)



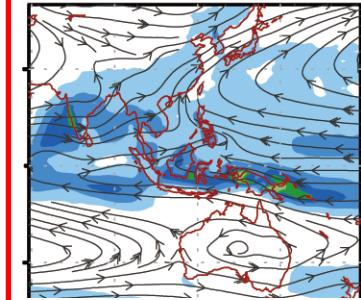
TP04

TP04(prec,uv850 JJA)

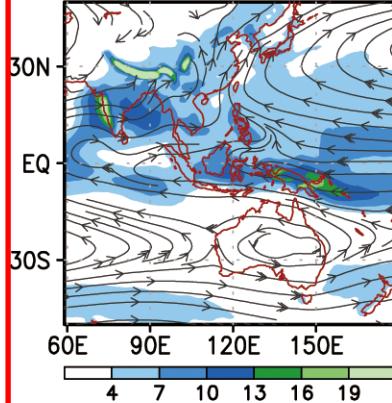


TP01

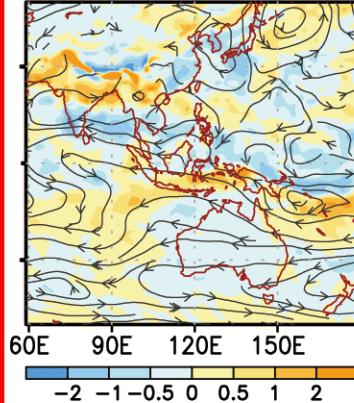
TP01(prec,uv850 JJA)



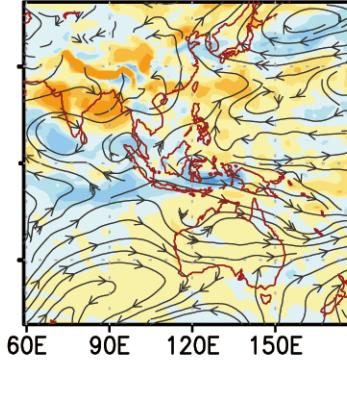
TP10



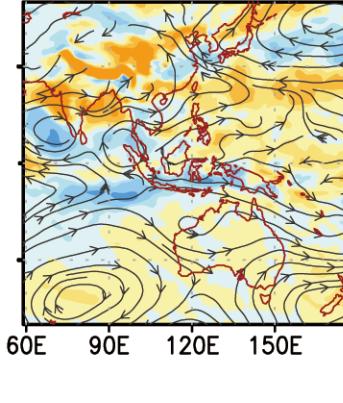
TP12 – TP10



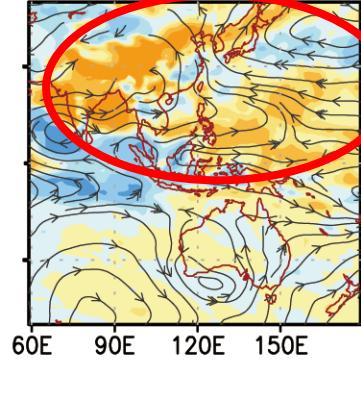
TP10 – TP07



TP10 – TP04



TP10 – TP01

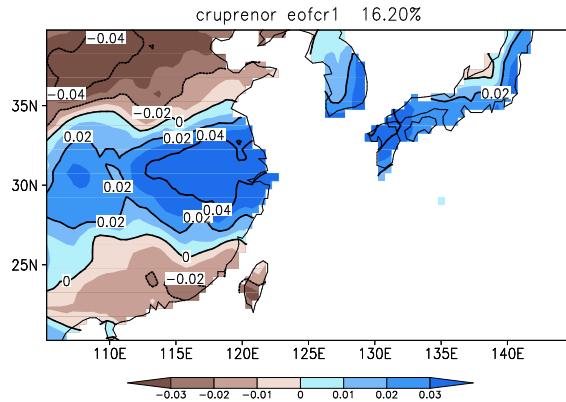


JJA (Precipitation, 850hPa winds)

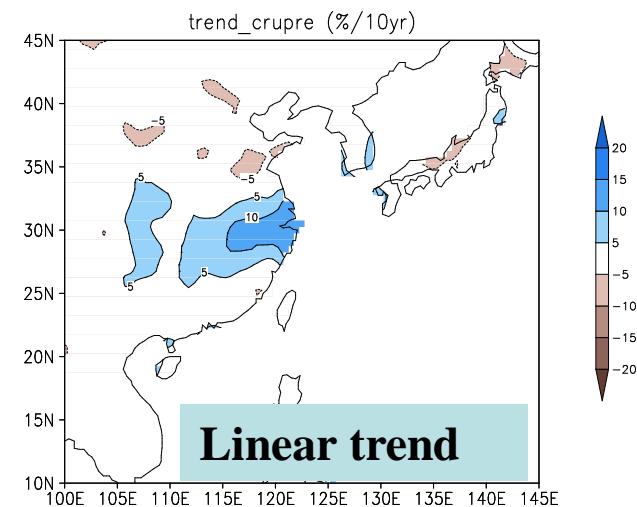
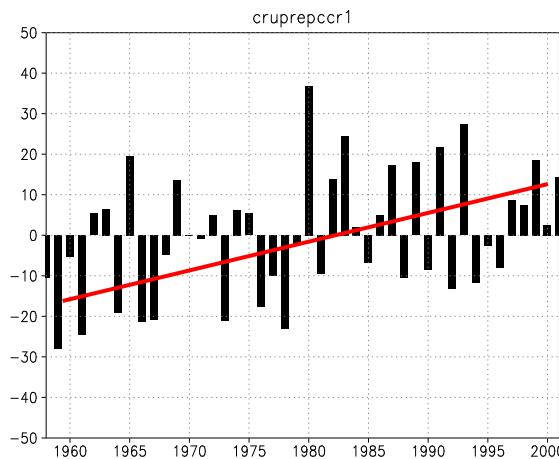
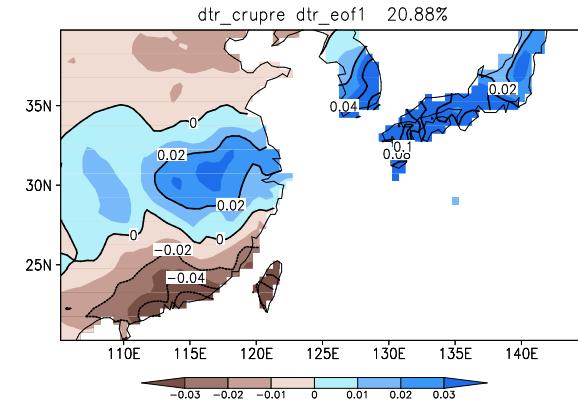
Tri polar Rainfall Pattern (CRU rainfall) (Pacific-Japan Pattern)

multiple (cross) time scales and mechanism

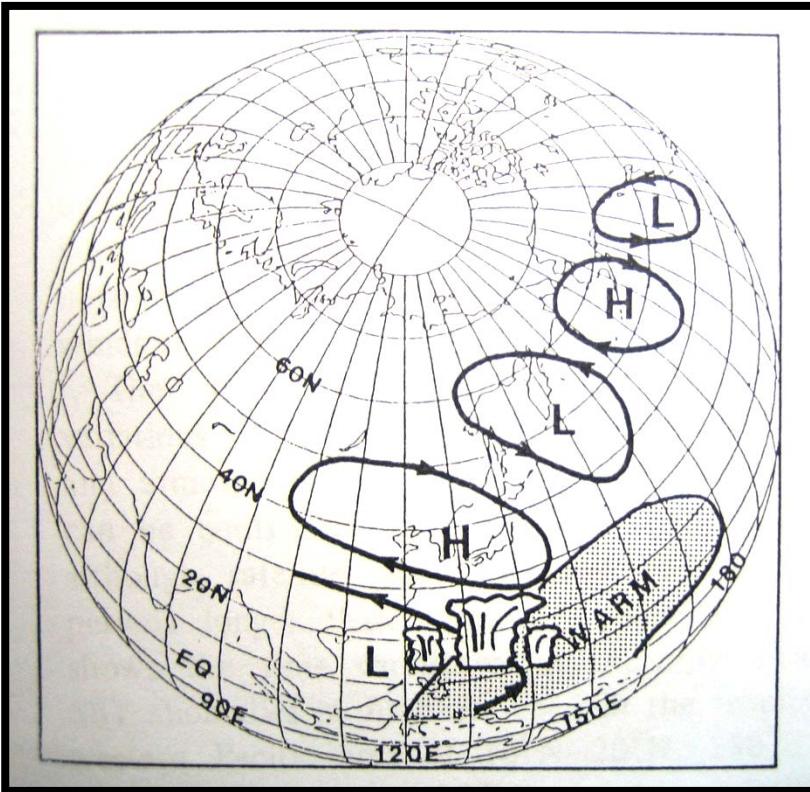
decadal+interannual



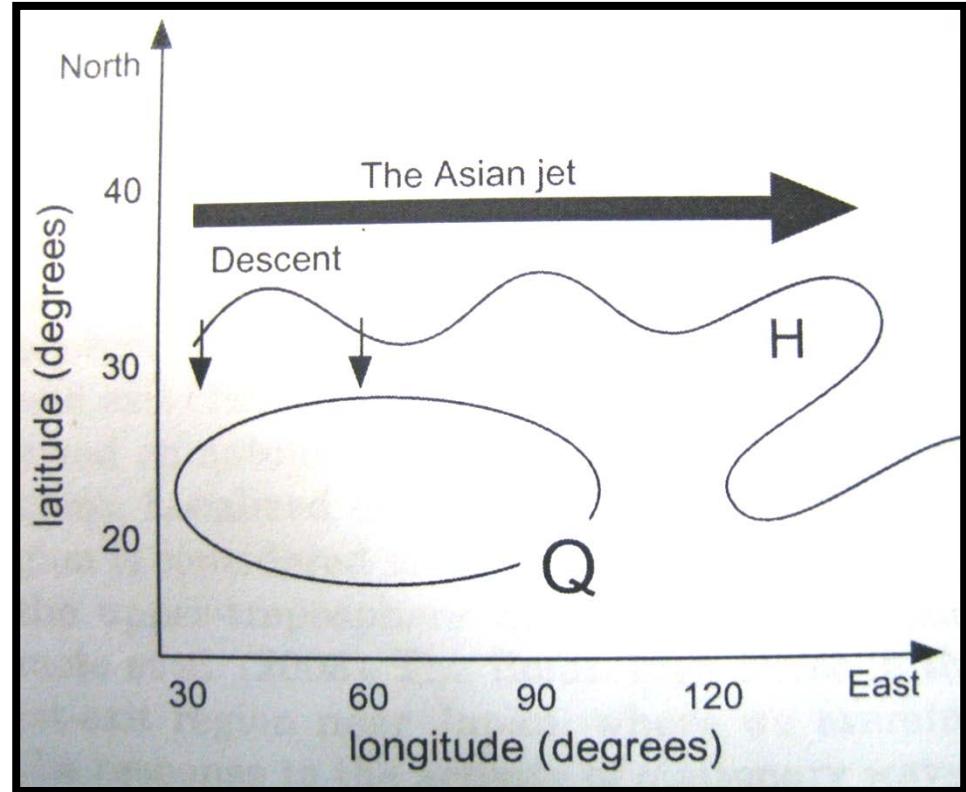
interannual



Nitta (1987) ...
Pacific Japan Pattern

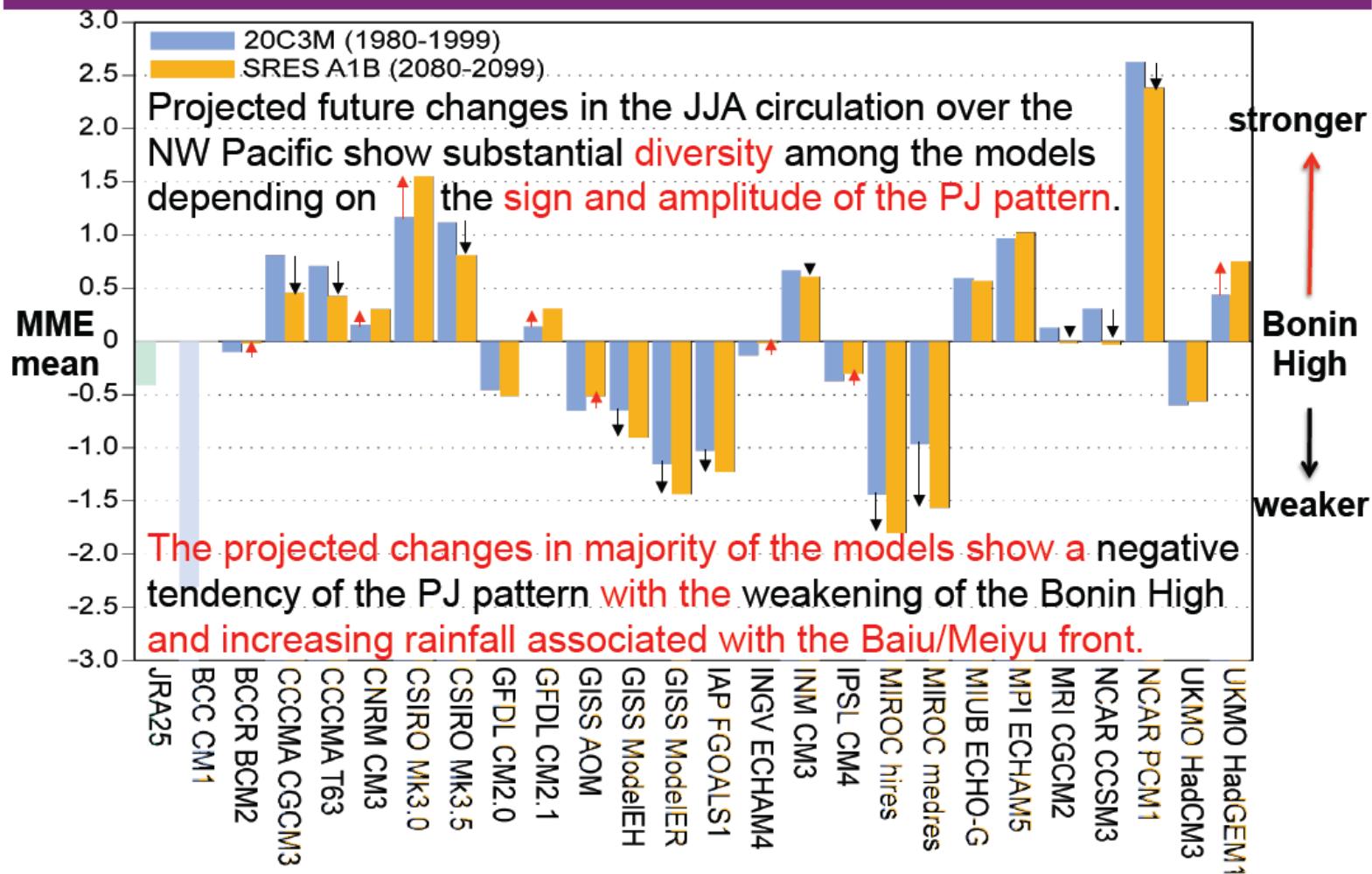


Enomoto et al. (2003) ...
Silk Road Pattern



- Wave-like perturbations forced by tropical WP or Tibetan heating
- ENSO effect
- Downstream propagation from Eurasian continent
 - Can be easily induced by various perturbation and forcing
 - Intrinsic mode? (Kosaka and Nakamura, Hsu et al., Hirota and Takahashi, ...)

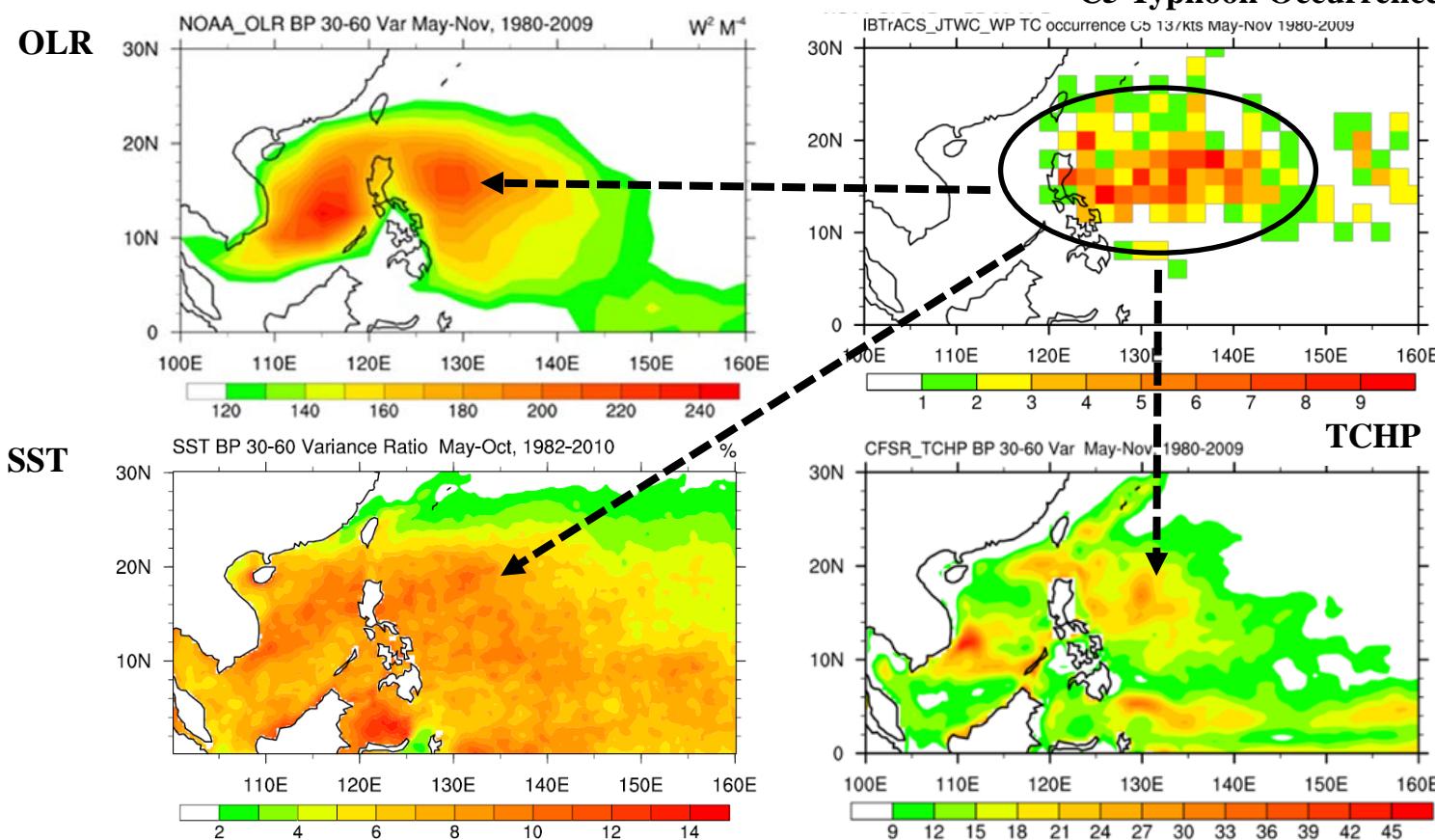
Projection of model bias and future projection (both from 20C MME climatology) in JJA 850-mb vorticity onto inter-model EOF1 (PJ-signal)
 Kosaka, Nakamura (2010, submitted to JC)



courtesy of Nakamura

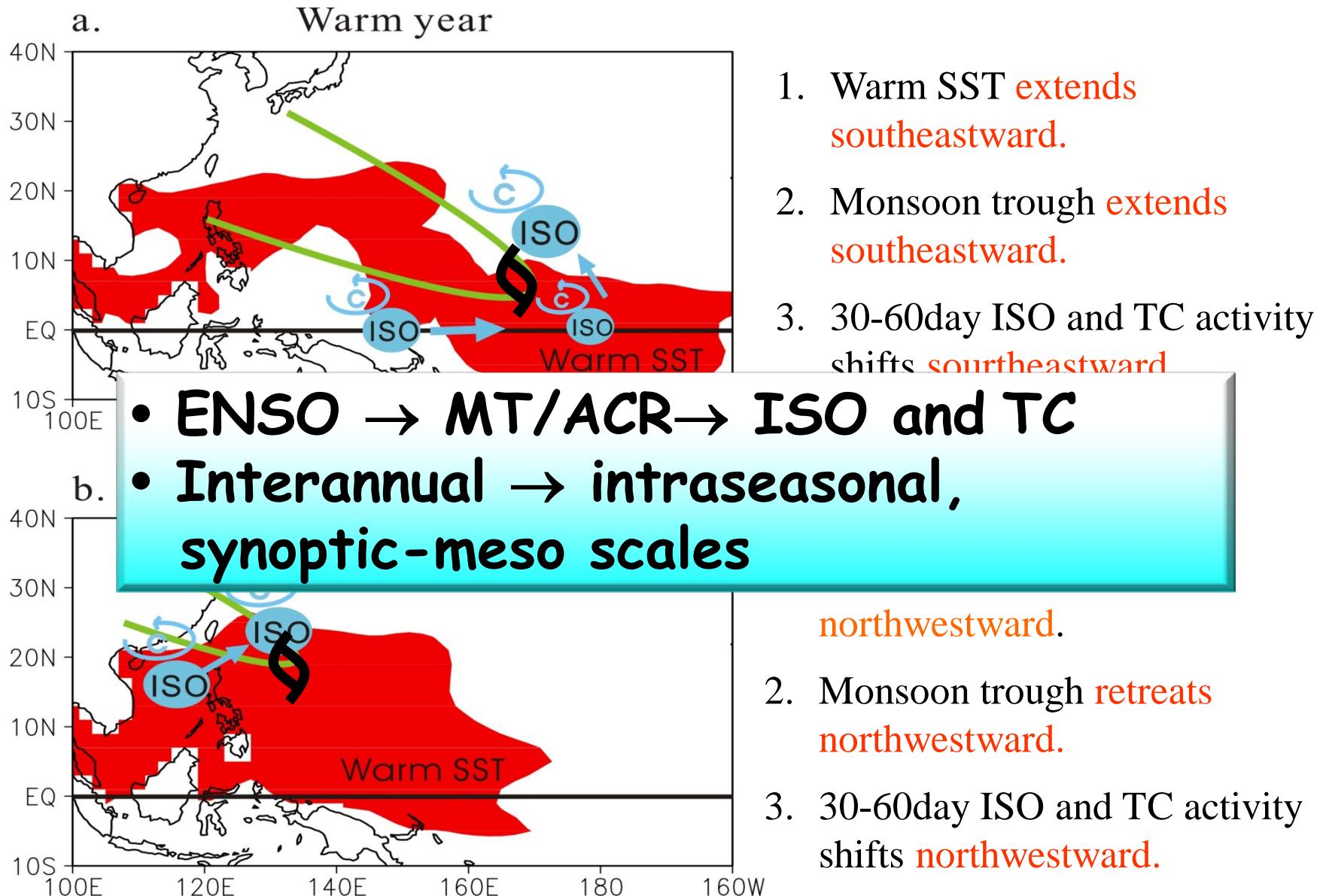
Multiscale Interaction in the Tropical Western North Pacific

Hot Spots of A-O coupled Intraseasonal Oscillation (30-60 day) and TC during May-October (Hsu, 2005, 2010)



C5 Typhoon Occurrence collocated with intraseasonal convection and TCHP variation

Hsu and Weng, 2001,
J. Climate; Weng and
Hsu, 2017, *GRL*



鄒治華(Tsou)、徐邦琪(Hsu)

TC contribution to Climate Variability

44-year mean

original

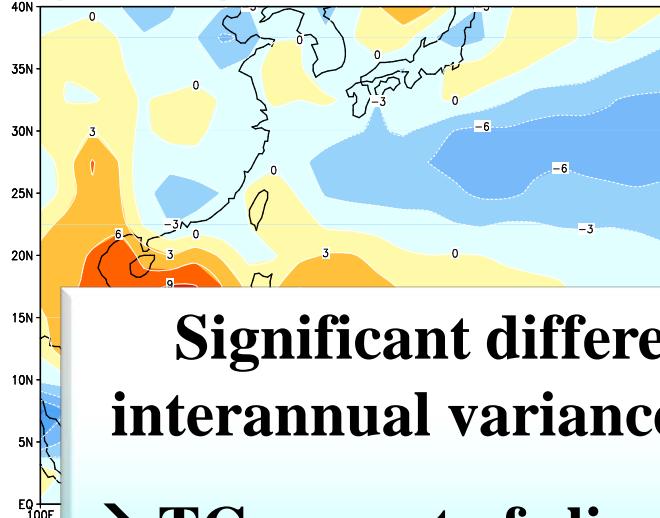
1958-2001 JJASO

850 hPa vorticity

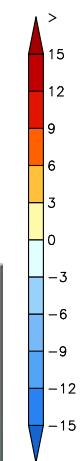
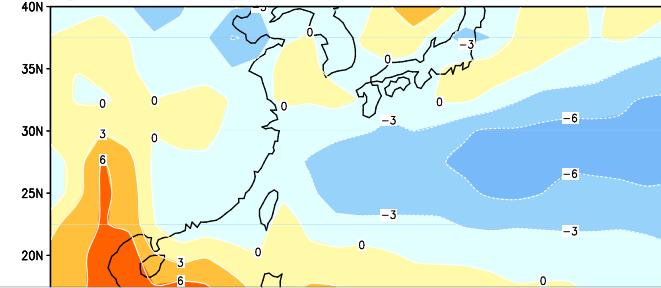
TC-removed

Hsu et al. 2008

58~01 6/16~10/16 850hPa era40_daily vorticity mean
[E-6 1/s**2]



58~01 6/16~10/16 850hPa tyfil_daily vorticity mean
[E-6 1/s**2]

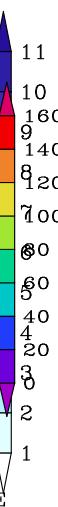
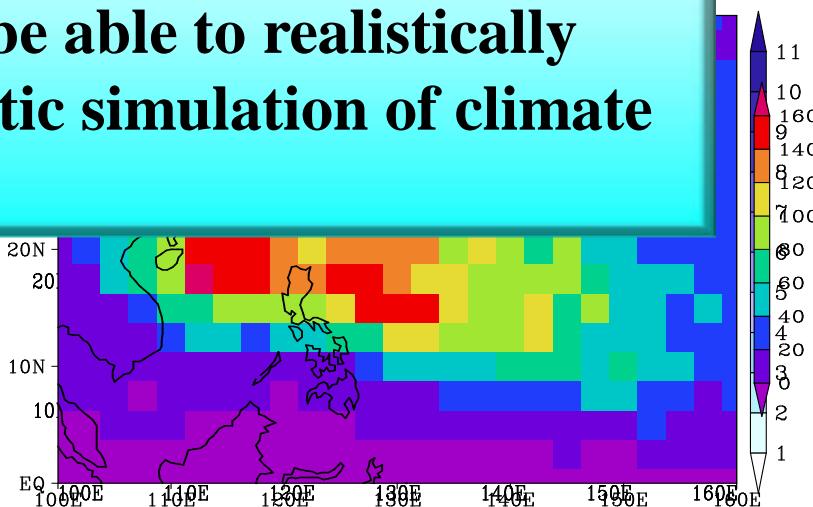
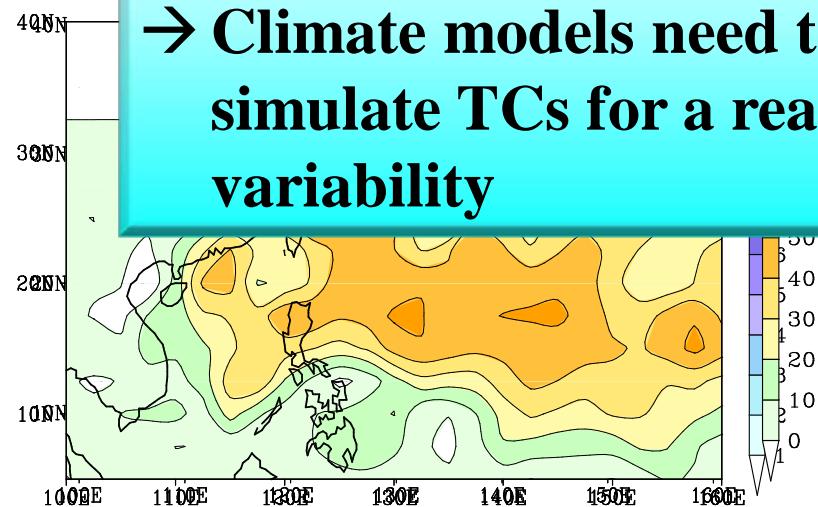


Significant differences in long-term mean and interannual variance (along TC tracks and more?)

→ TC as part of climate variability

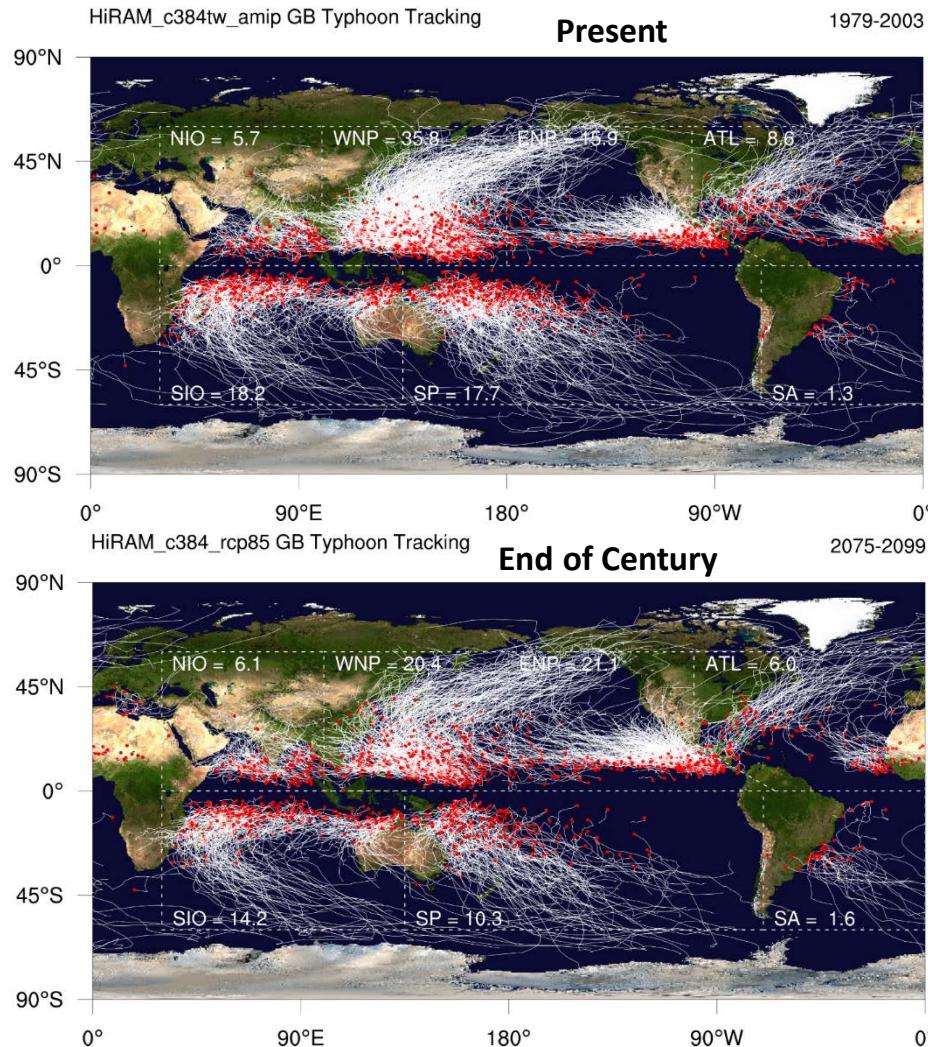
→ Climate models need to be able to realistically simulate TCs for a realistic simulation of climate variability

Interannu



Future Changes in TC Activity

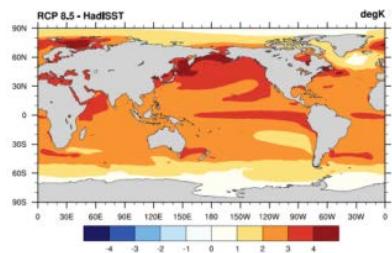
(HiRAM)



TC Projection (C384)

similar in C192

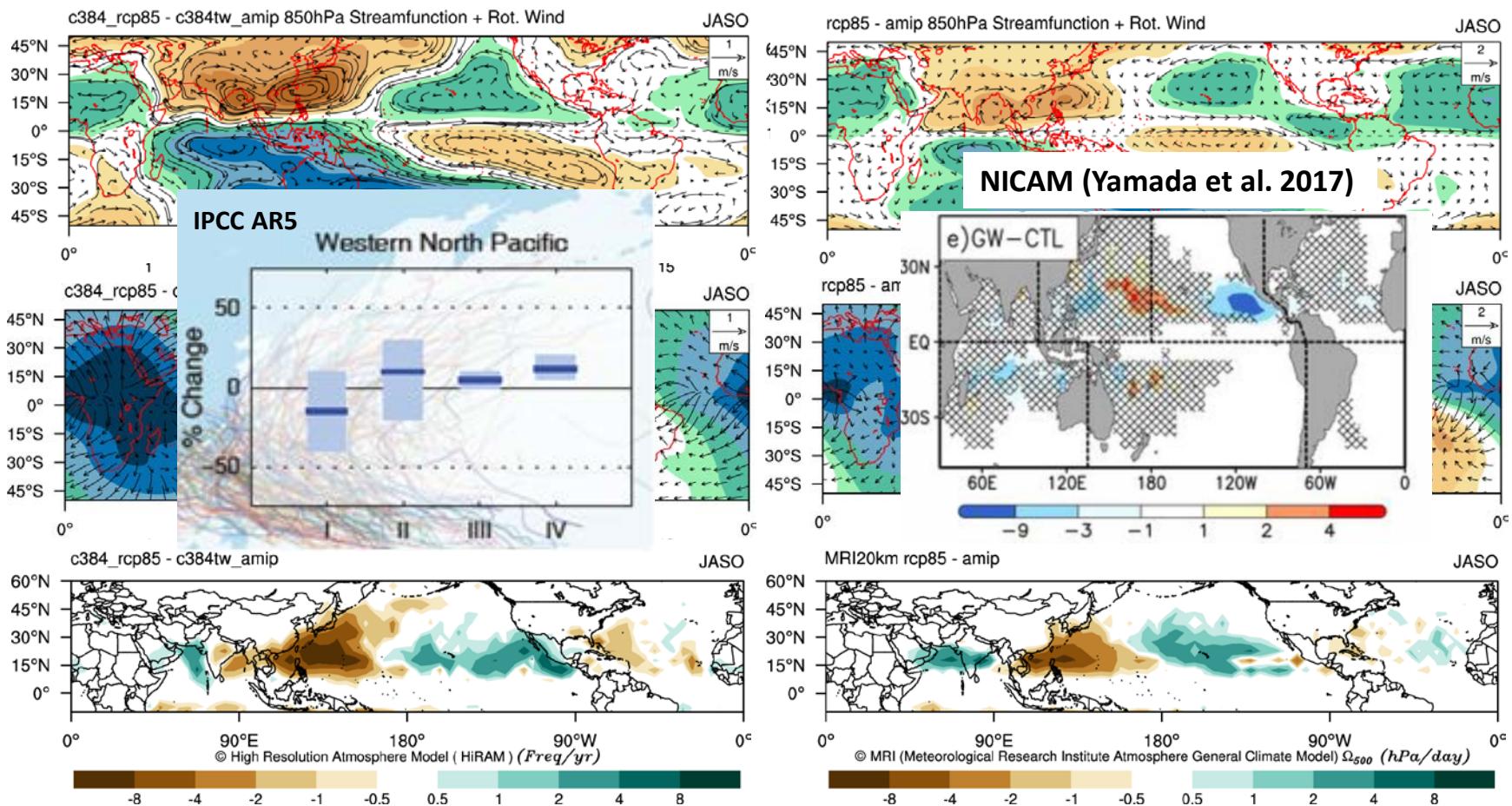
TC Counts	AMIP	RCP8.5
NIO	5.7	6.1
WNP	35.8	20.4
ENP	15.9	21.1
ATL	8.6	6.0
SIO	18.2	14.2
SP	17.7	10.3
SA	1.3	1.6



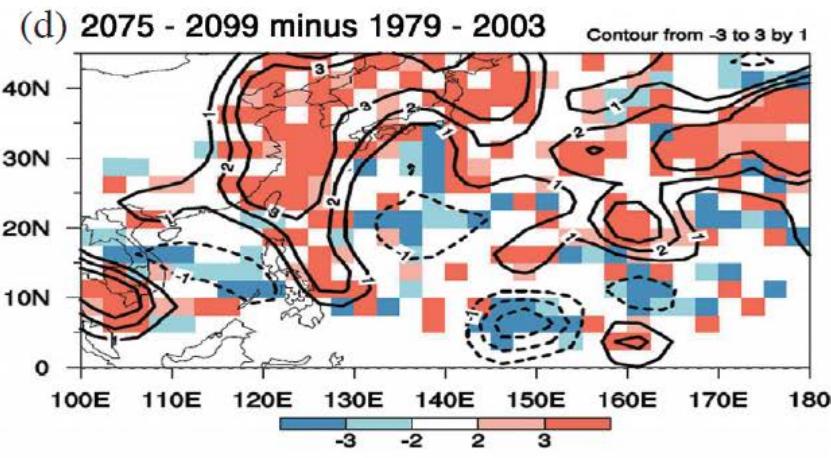
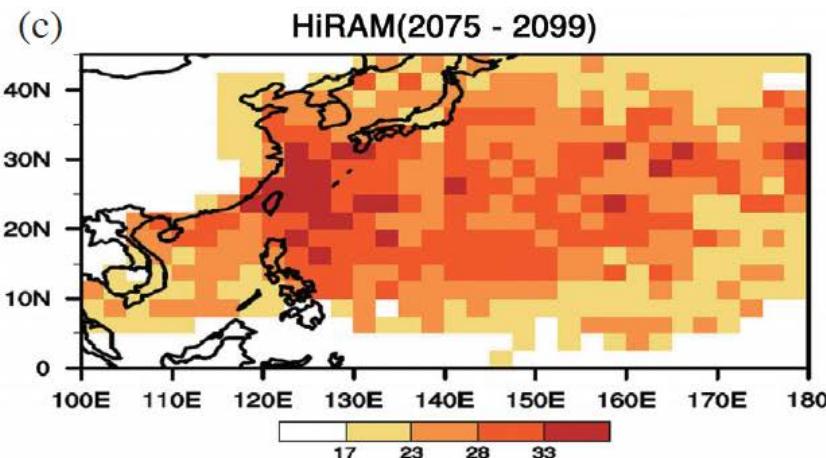
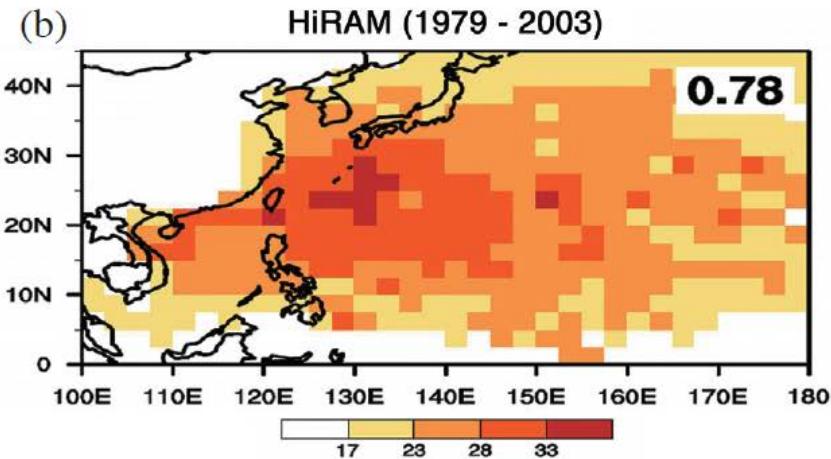
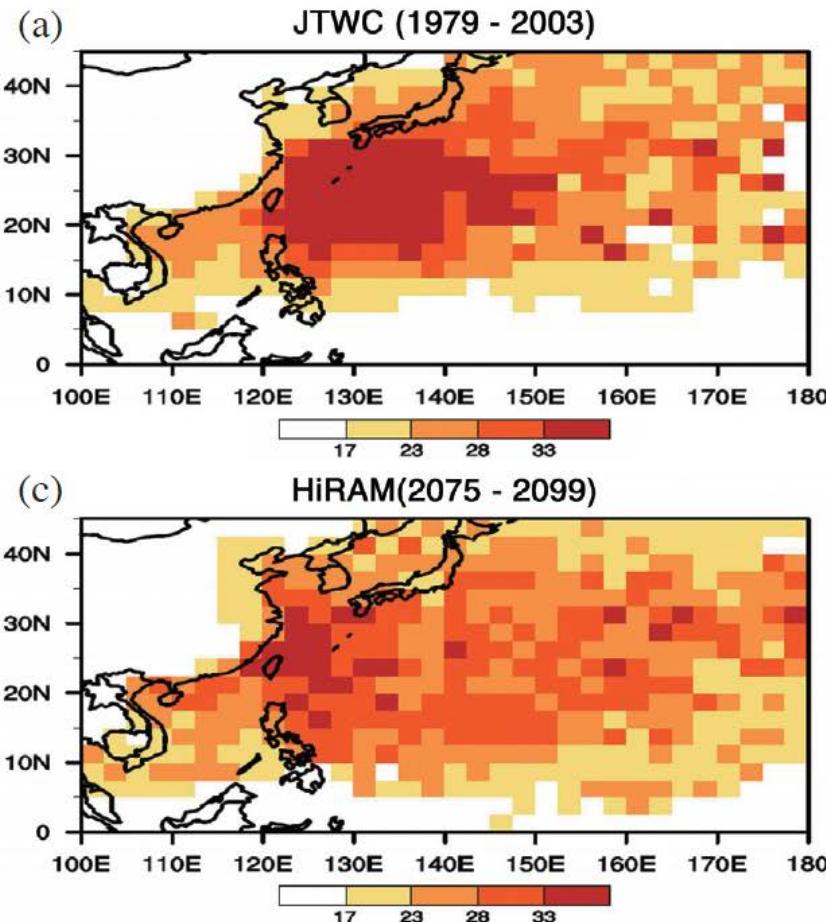
HiRAM C384

vs.

MRI AGCM20km

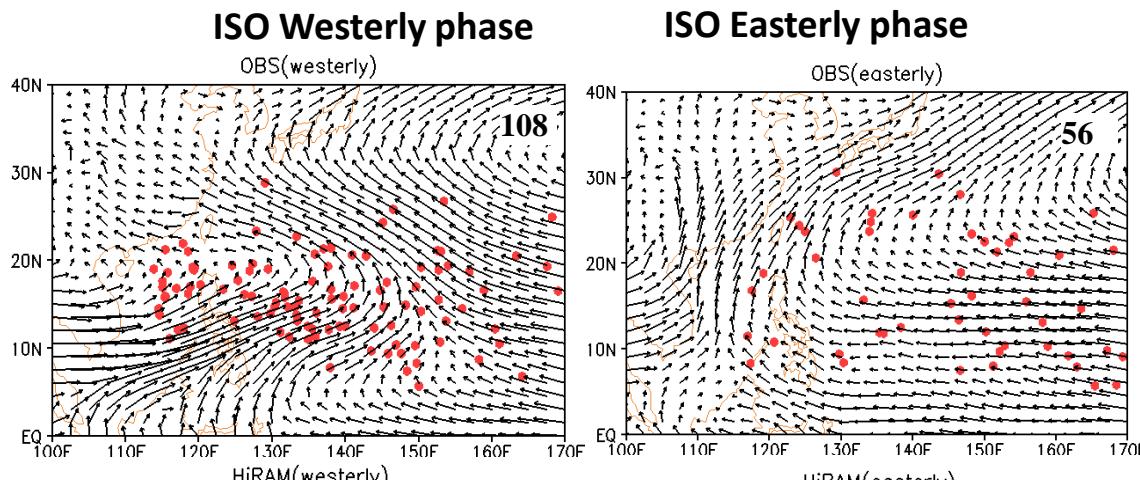


TS intensity distribution (JJASON)

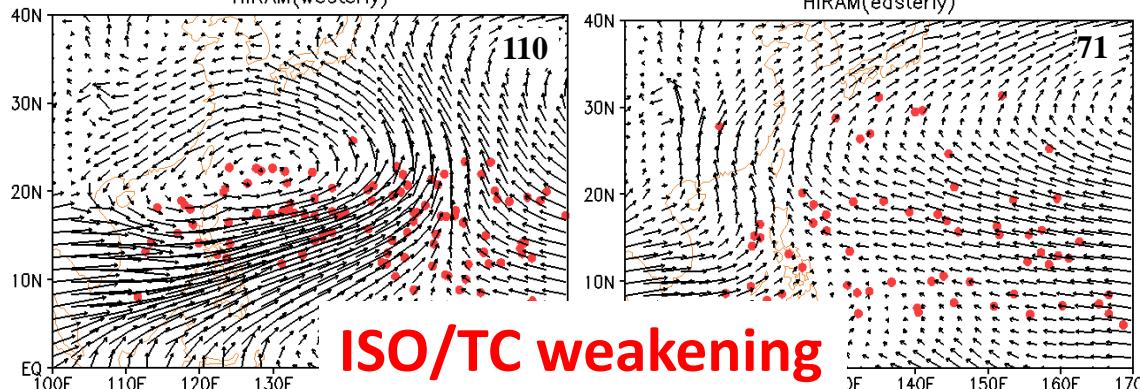


Changes in ISO and TC genesis in the western North Pacific (Jul.-Sept.)

JTWC

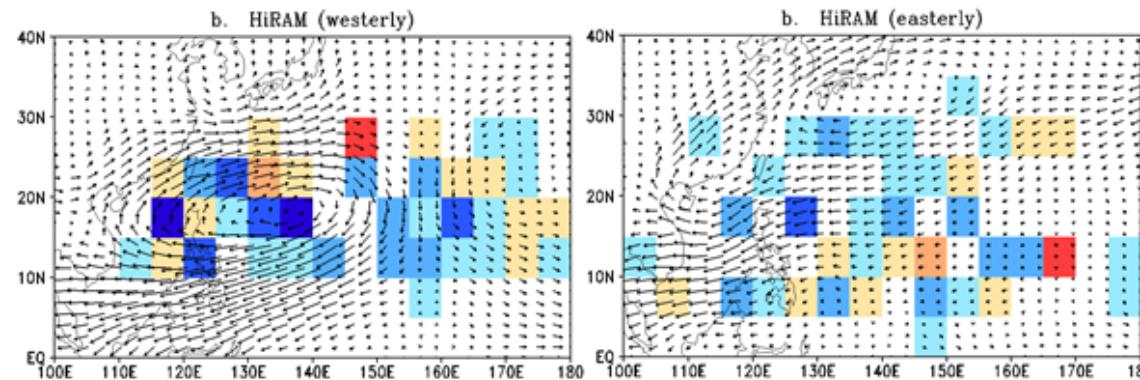


HiRAM



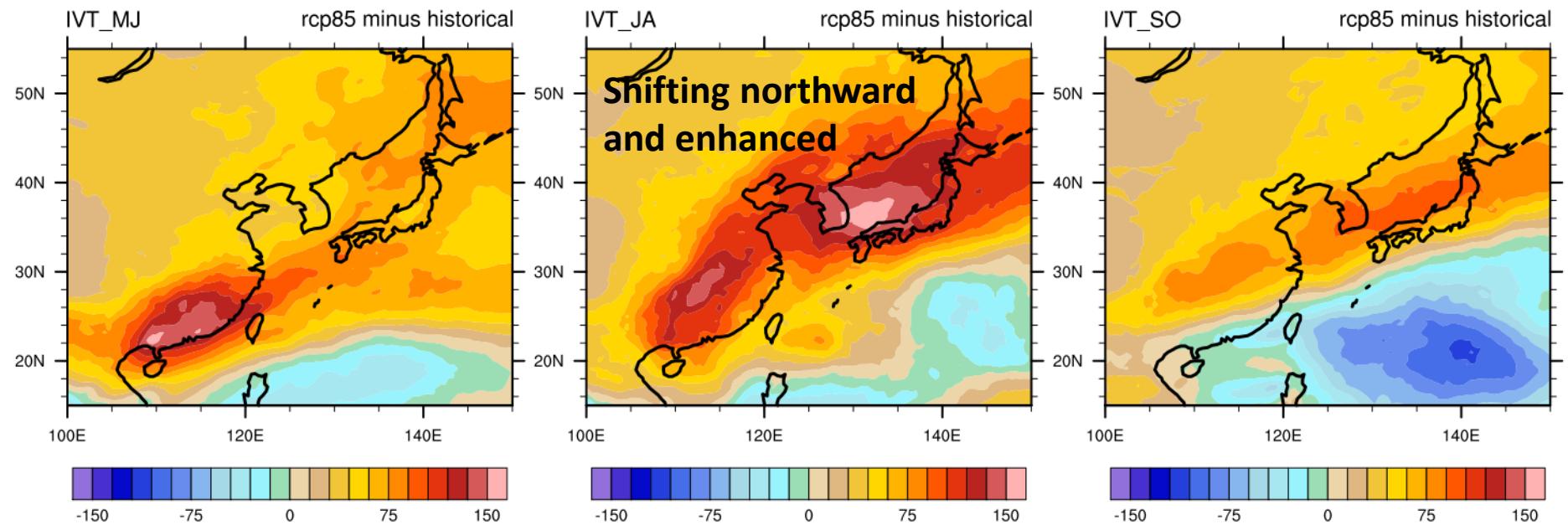
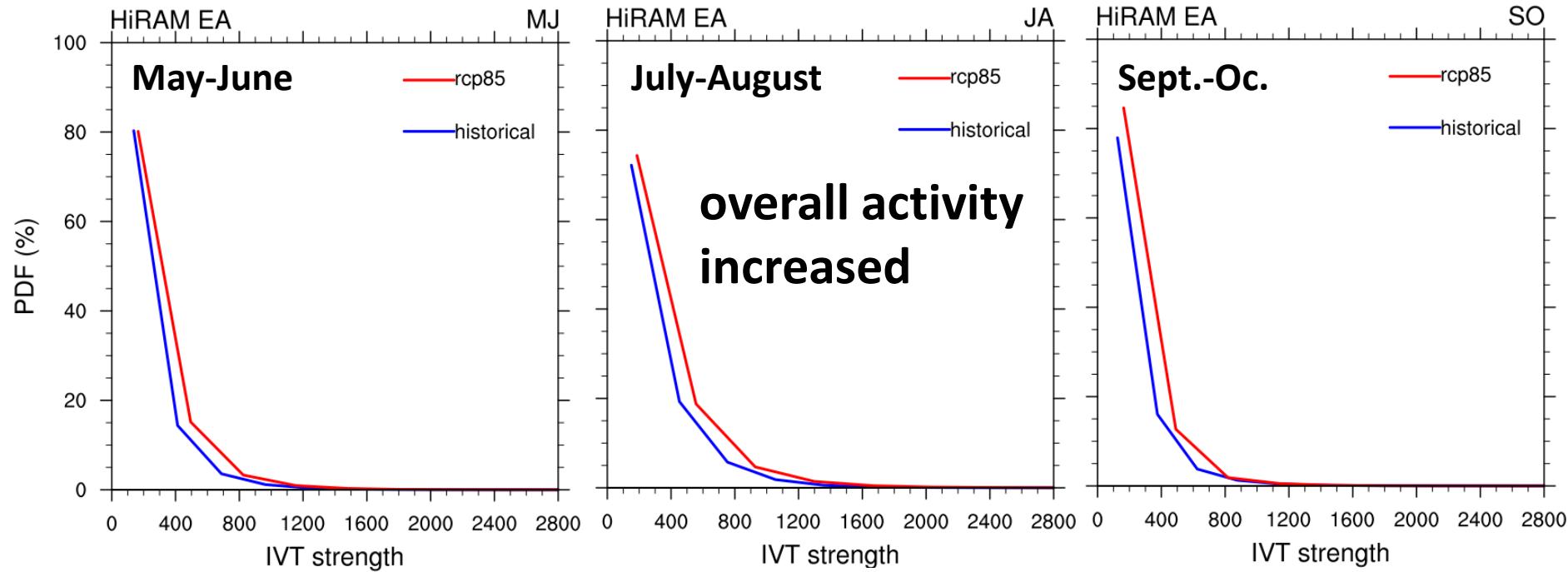
ISO/TC weakening

2075-2099
Minus
1979-2003

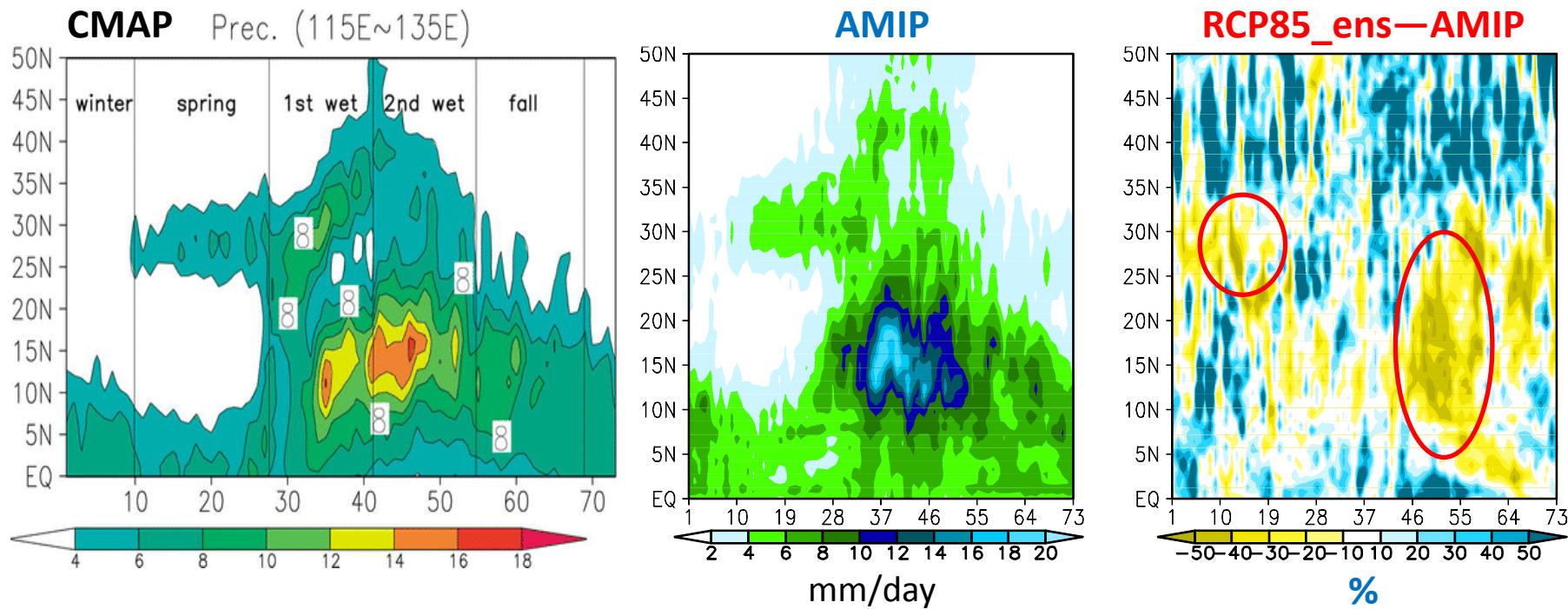


(C.-H. Tsou)

Changes in Atmospheric River Activity

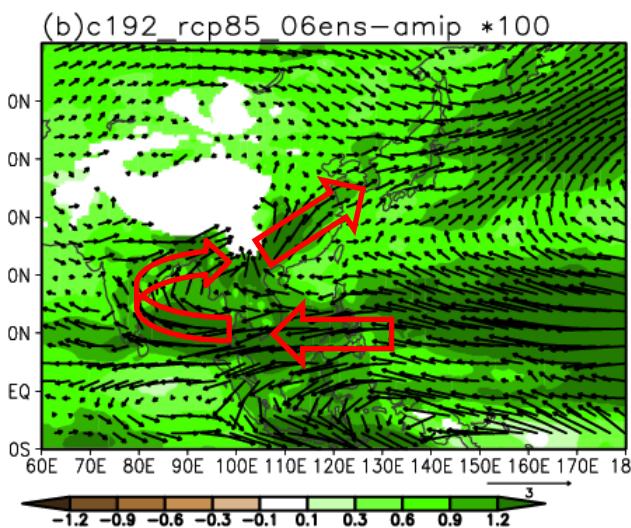
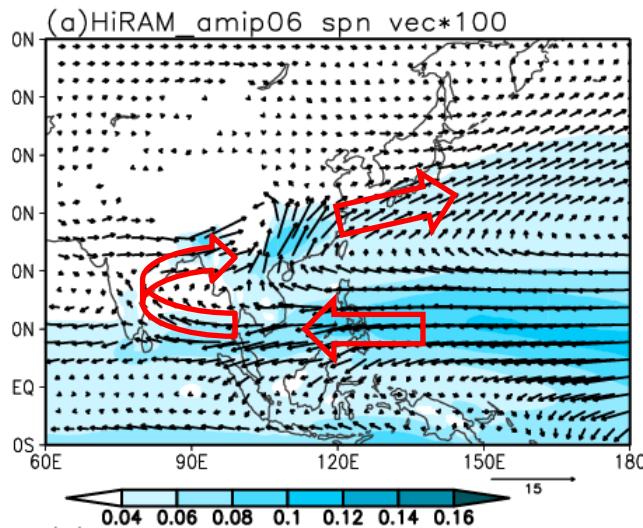


Changes in Annual Precipitation Cycle Projected by HiRAM

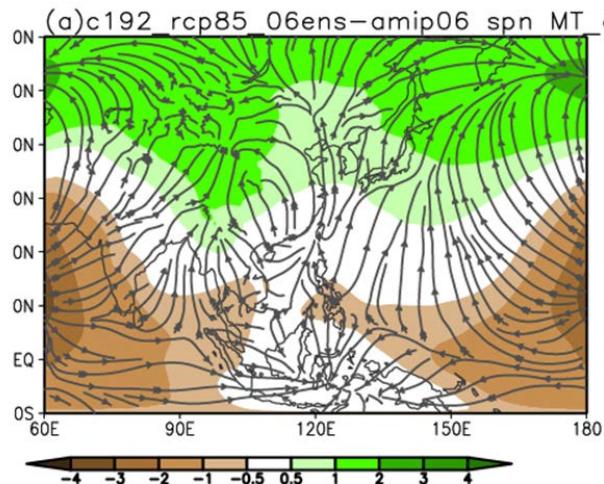


Spring 2/15-5/15

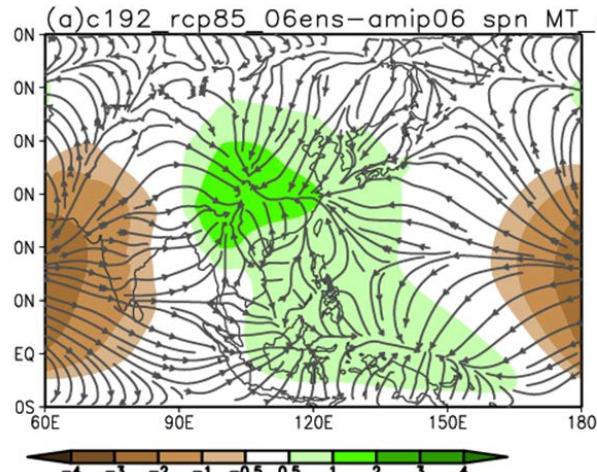
Enhanced anticyclonic circulation and moisture transport



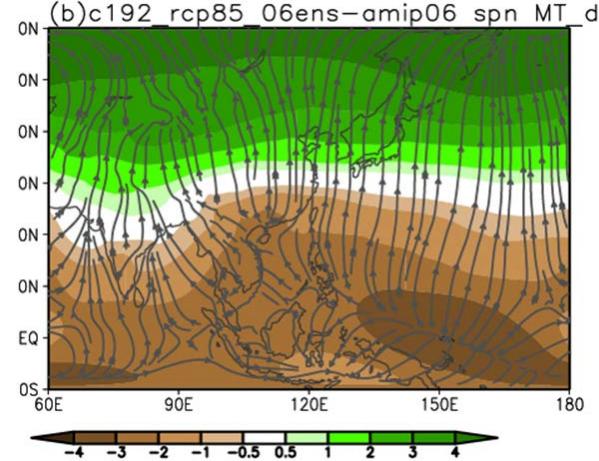
850hPa moisture transport



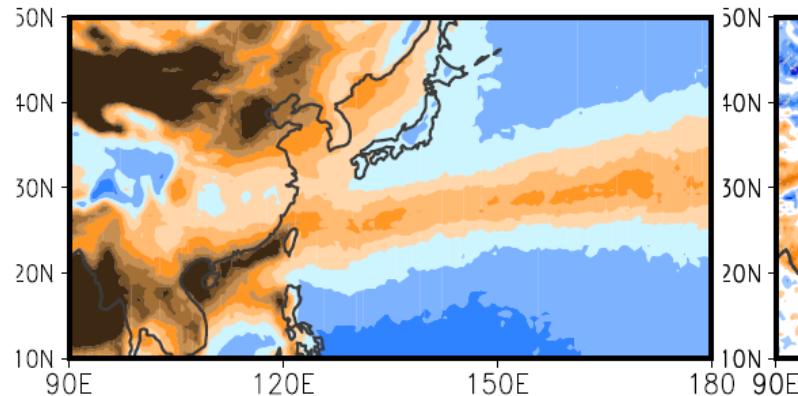
Difference associated with
thermodynamic part (increased q)



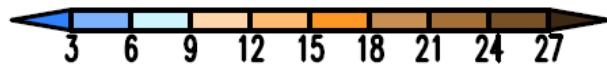
dynamic part (changes in u,v)



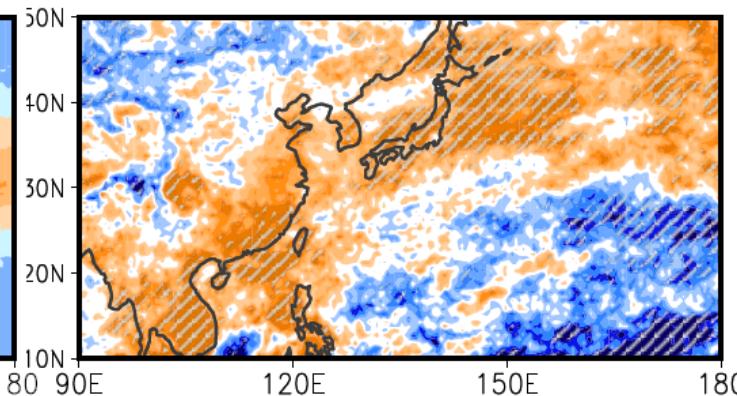
(a)HiRAM_amip06 CDD spn



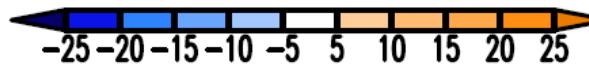
CDD (unit: day)



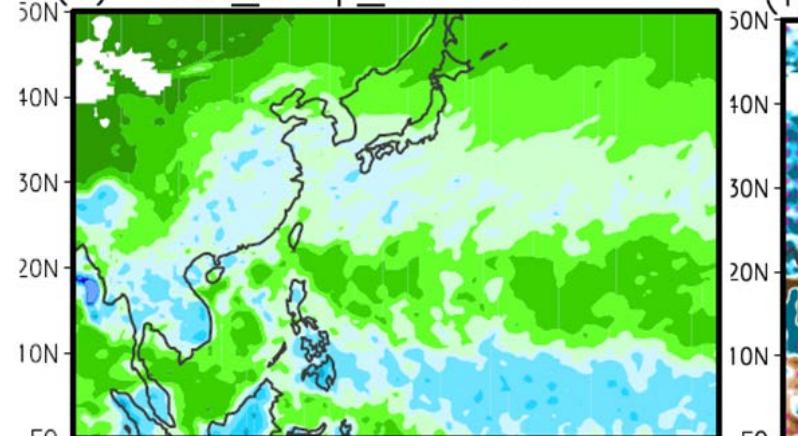
(b)HiRAM_rcp85_06ens-amip06



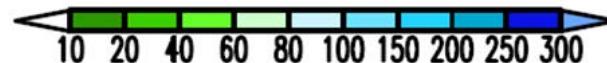
Changes in CDD (unit: %)



(c)HiRAM_amip_06 Pr99

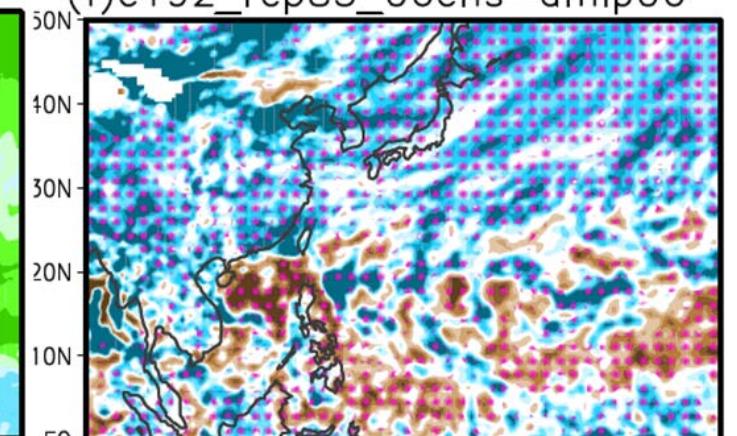


mm/day

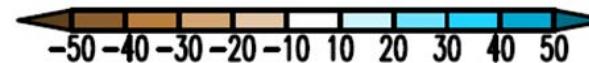


**Spring
CDD**

(f)c192_rcp85_06ens-amip06

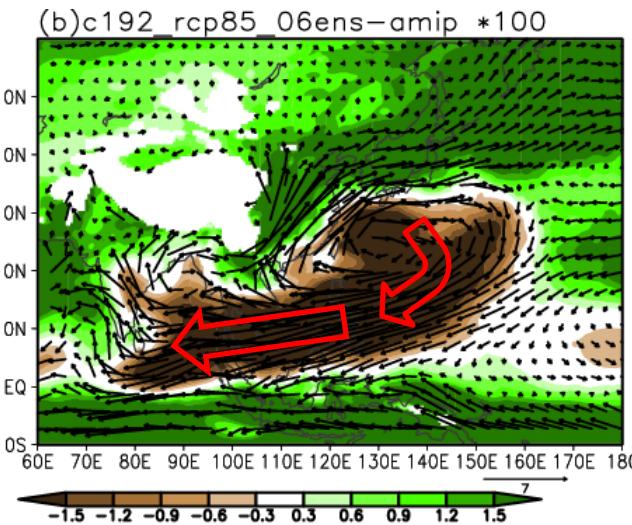
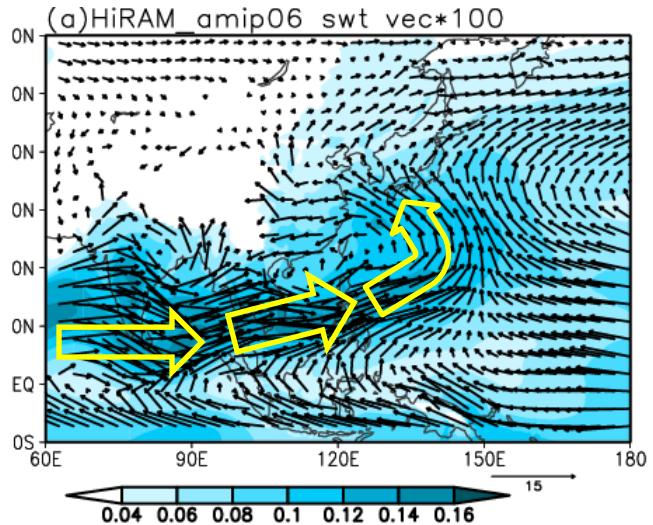


%



**Extreme
Rainfall
(Pr99)**

Hatched regions: 5% significance



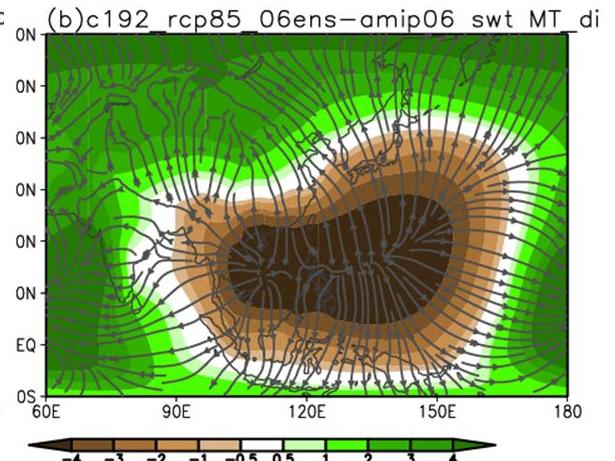
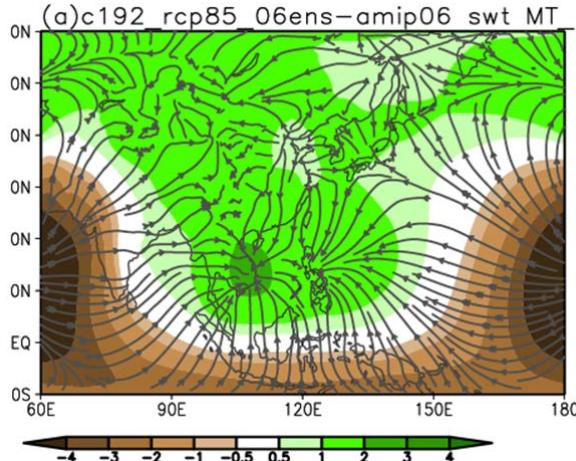
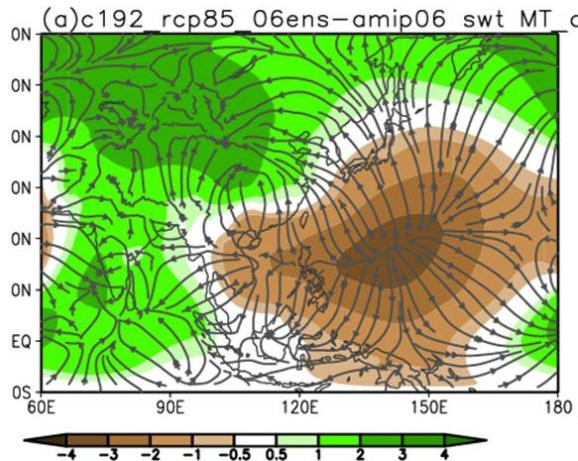
**Typhoon Season
7/25-9/27**

850hPa moisture transport
vector: uq, vq
shading: $(uq^2 + vq^2)^{1/2}$

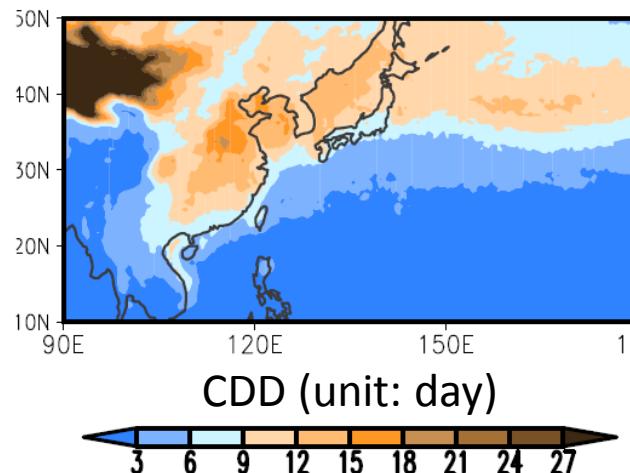
850hPa moisture transport

Difference associated with
thermodynamic part (increased q)

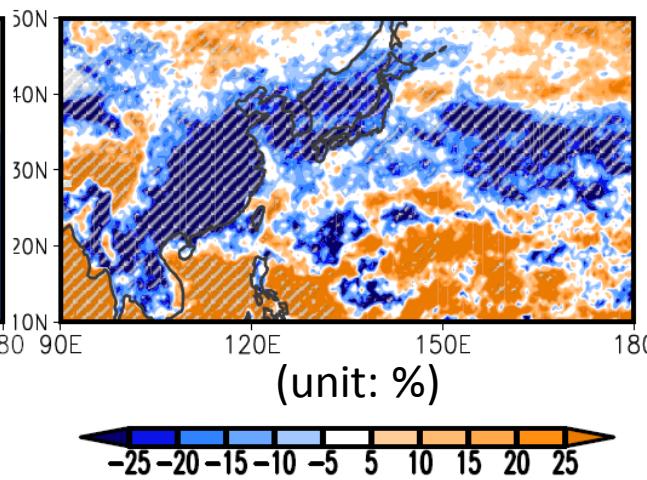
dynamic part (changes in u, v)



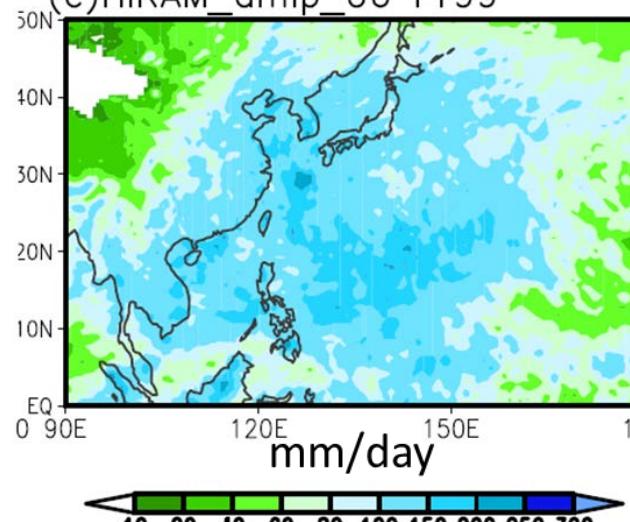
(a)HiRAM_amip06 CDD swt



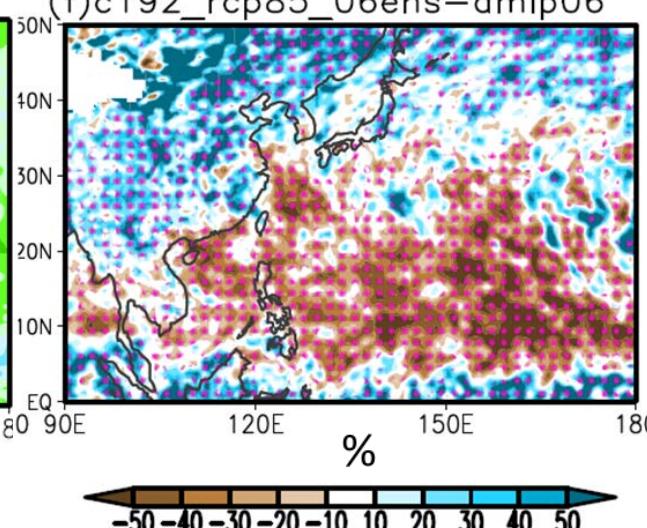
(b)HiRAM_rcp85_06ens-amip06



(c)HiRAM_amip_06 Pr99



(f)c192_rcp85_06ens-amip06



**Typhoon
Season
CDD**

**Extreme
Rainfall
(Pr99)**

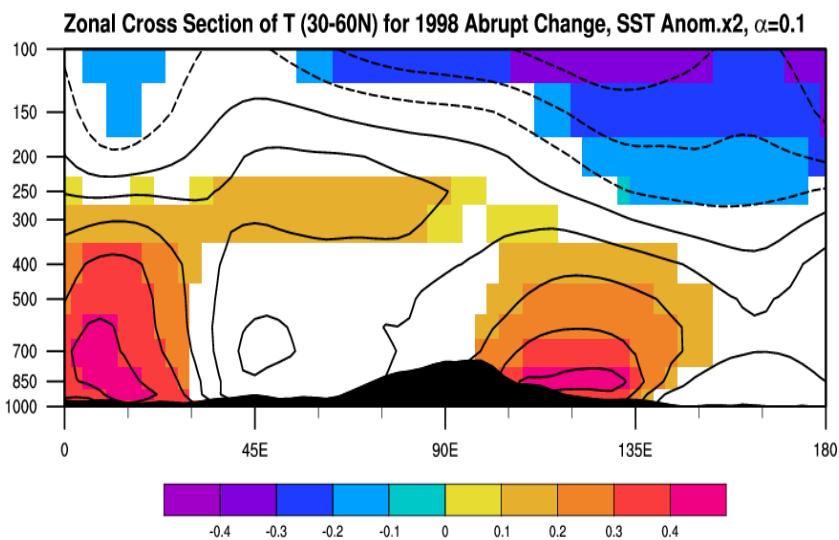
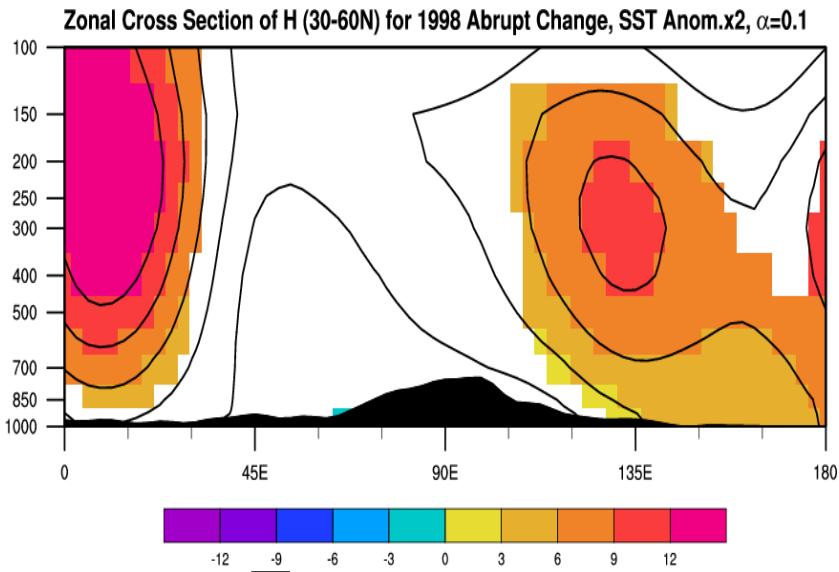
Hatched regions: 5% significance

Thank you for Your Attention Questions and Comments?



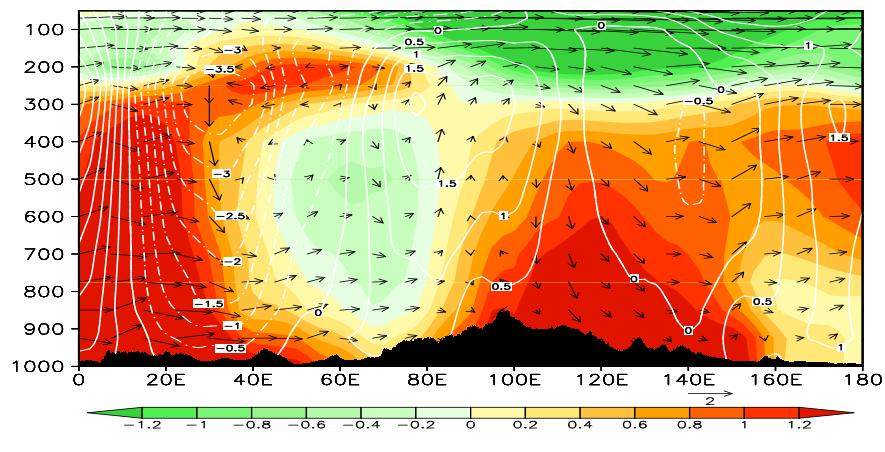
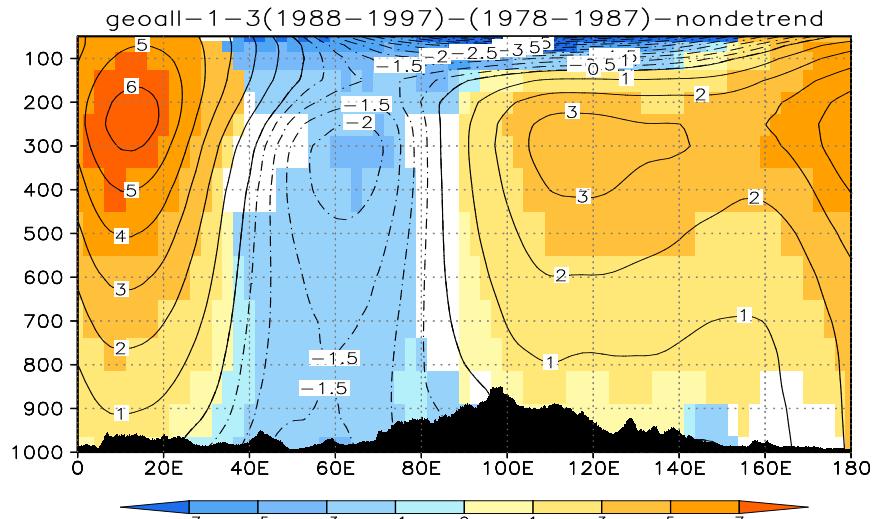
SSTA \times 2 simulation

Geopotential height

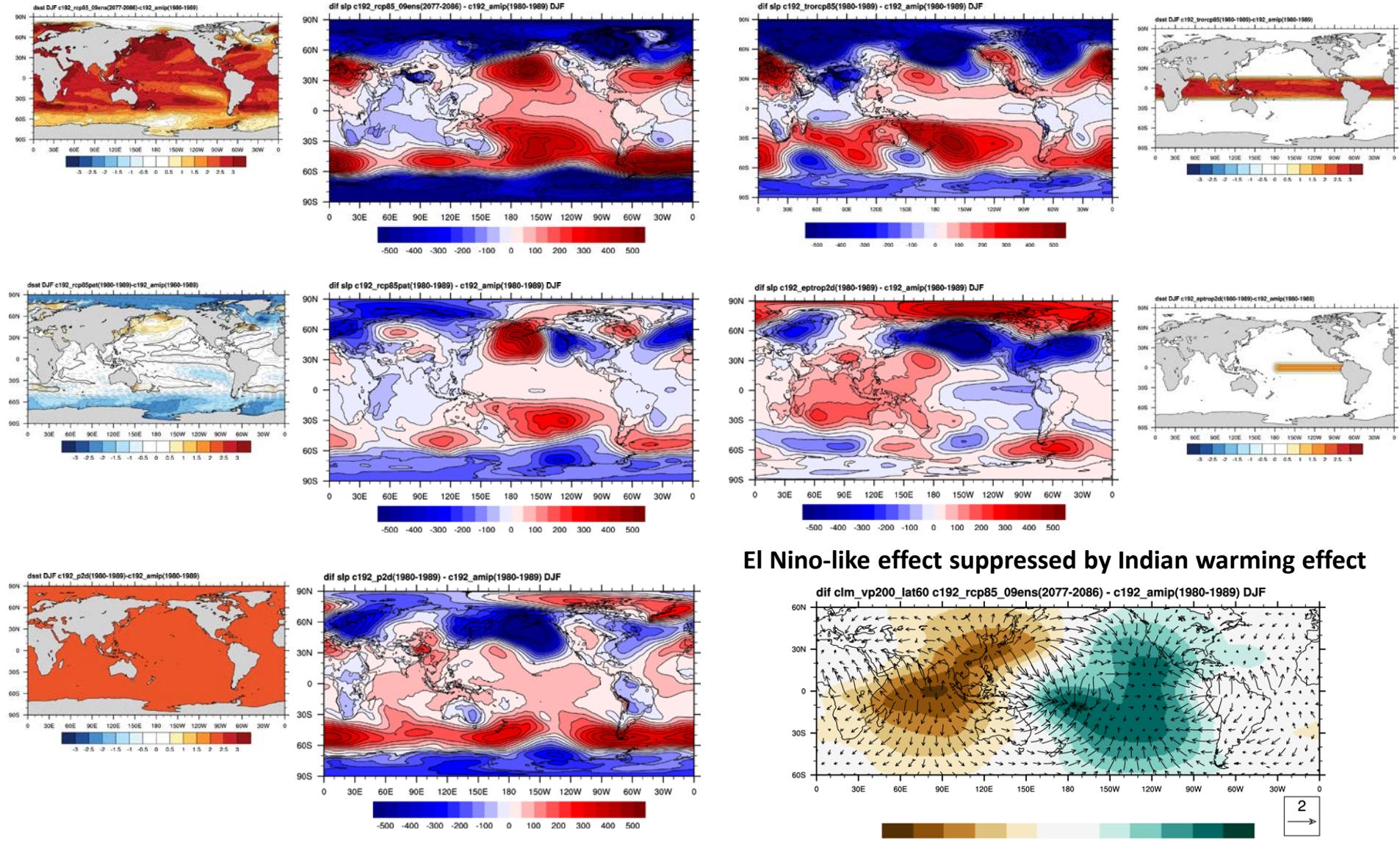


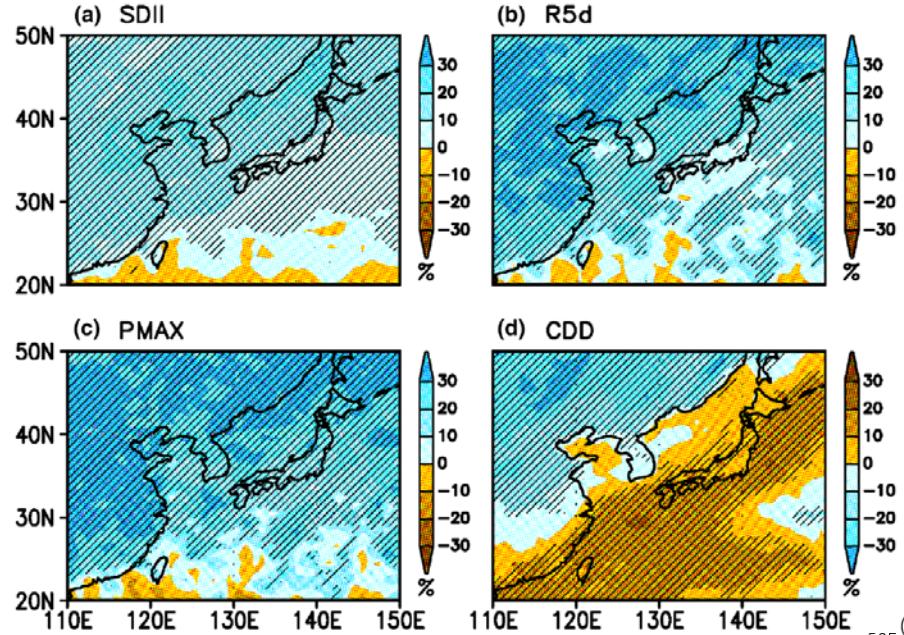
Observation

Geopotential height



Relative Contribution of SST Anomaly Pattern



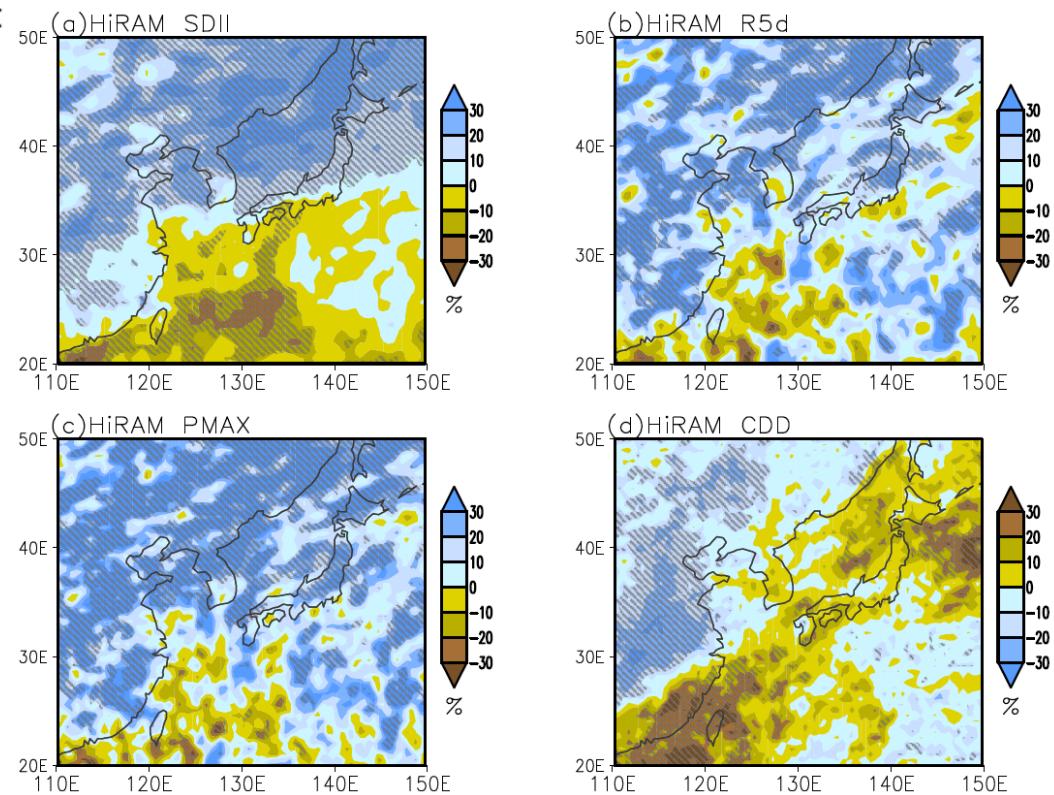


MRI(Kusunoki 2017)

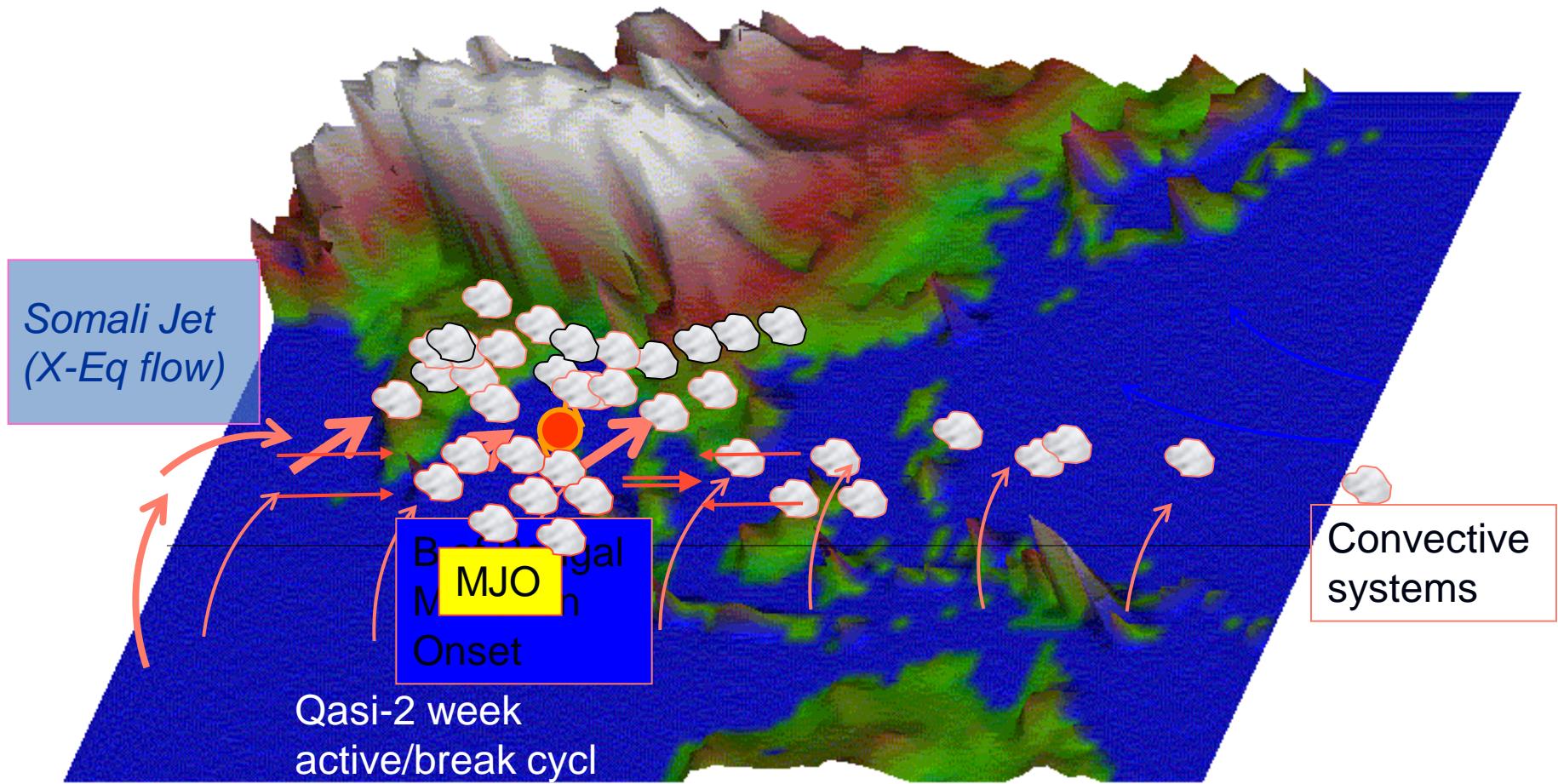
Fig. 9 Future changes in extreme precipitation events (annual statistics) projected by the all simulations. Unit is %. *Hatched regions* show changes above the 95% significance

Extreme Precipitation Comparison between MRI and HiRAM

HiRAM_C192



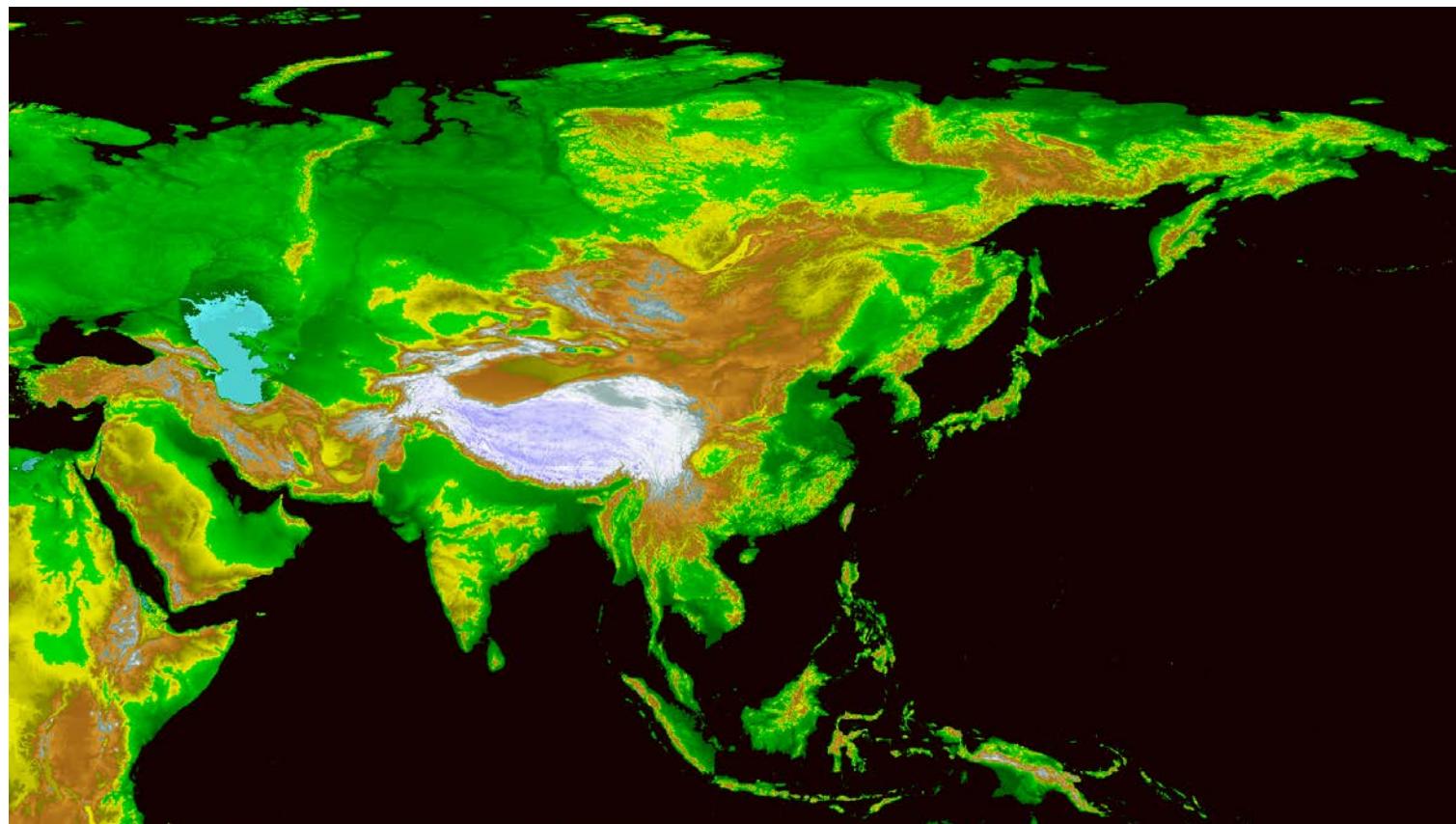
South Asian Summer Monsoon



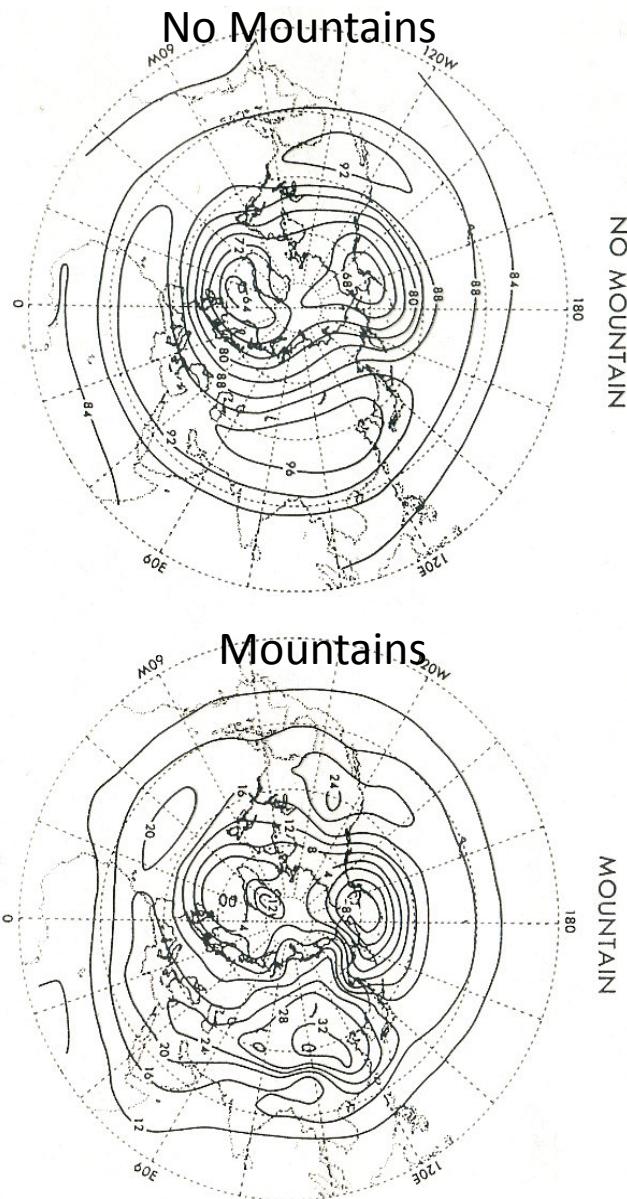
courtesy of C.-P. Chang

Mean Climatology

Role of Topography



Sea Level Pressure



The sea-level pressure for the mountain and no-mountain calculations, xx mb should be read 10xx mb for the mountain model and 9xx mb for the no-mountain model.

Topographic Effect - Boreal Winter

Mountains minus No Mountains

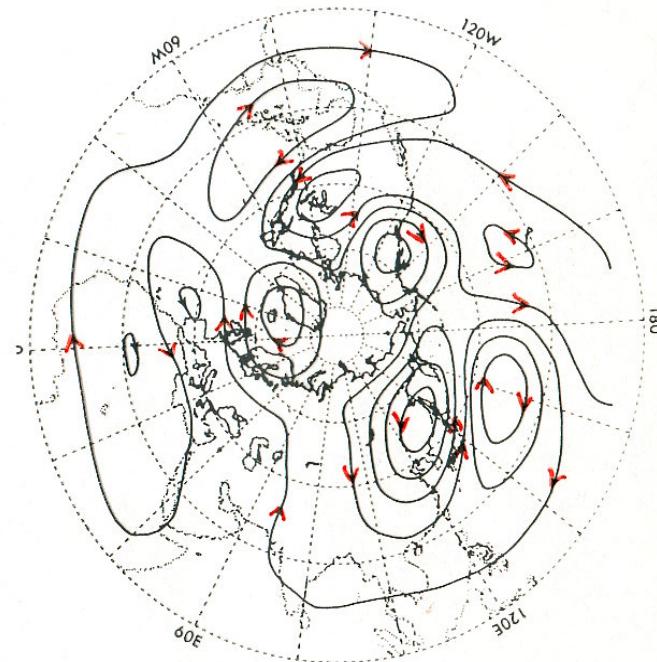
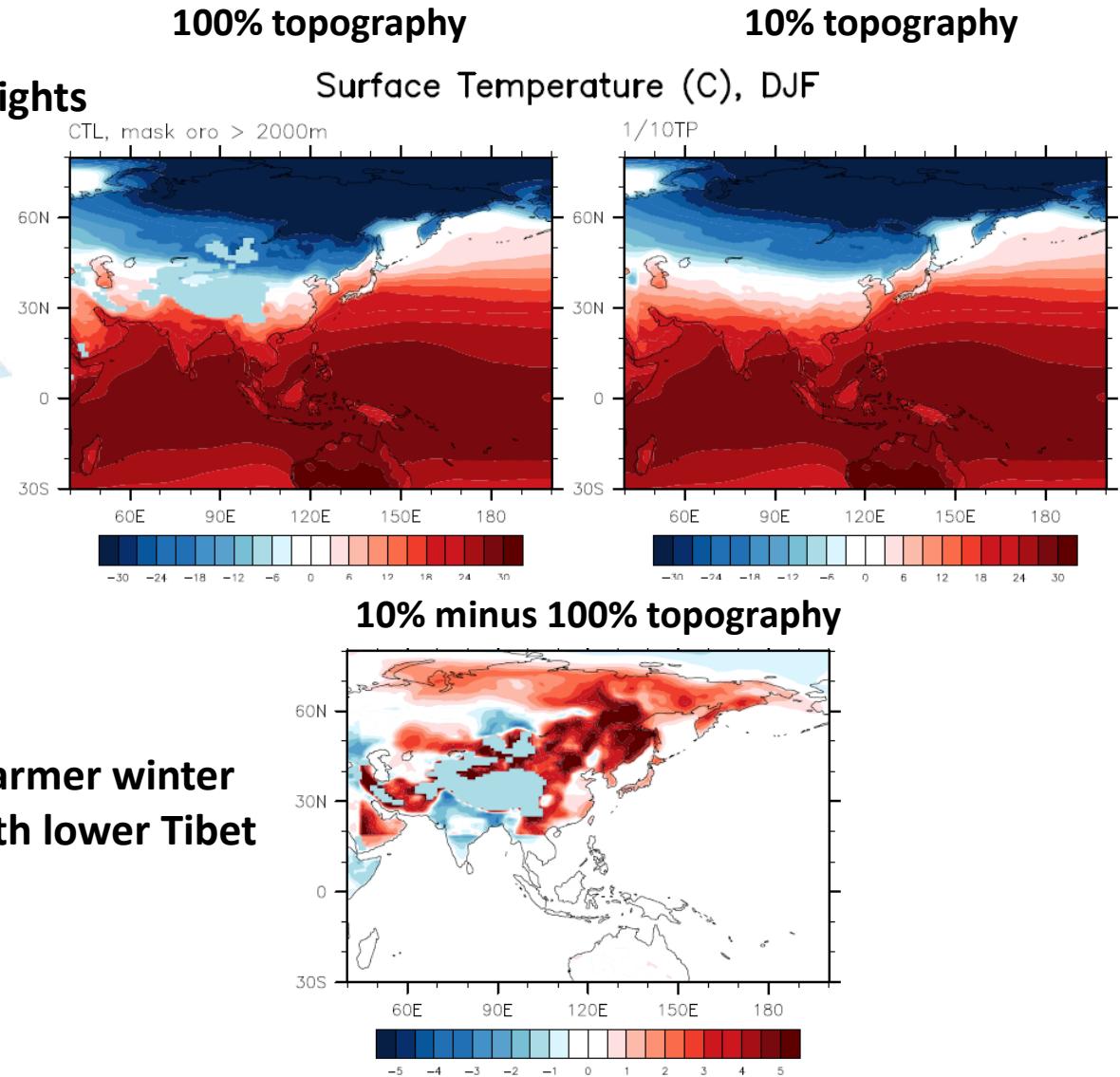
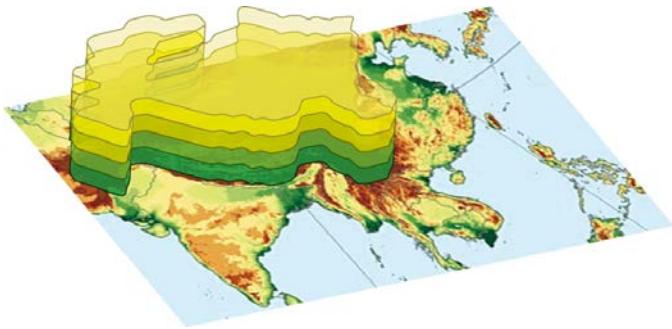


Fig. 6.23. The streamfunction for the mountain minus no-mountain flow at 300 mb. Contour interval is $4 \times 10^5 \text{ m}^2 \text{ s}^{-1}$.

Held (1983)

Effect of the Tibetan Plateau

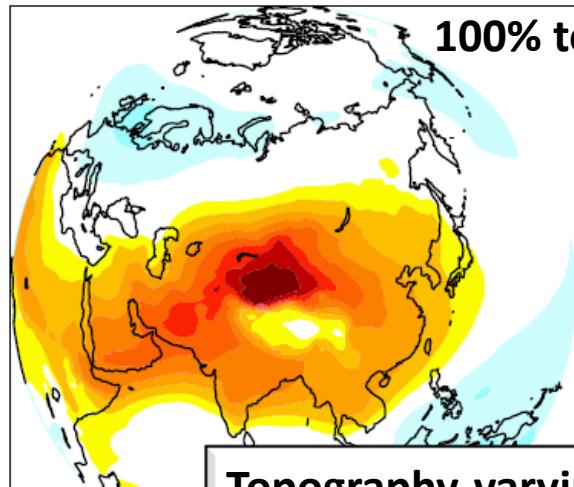
- Simulations with different heights



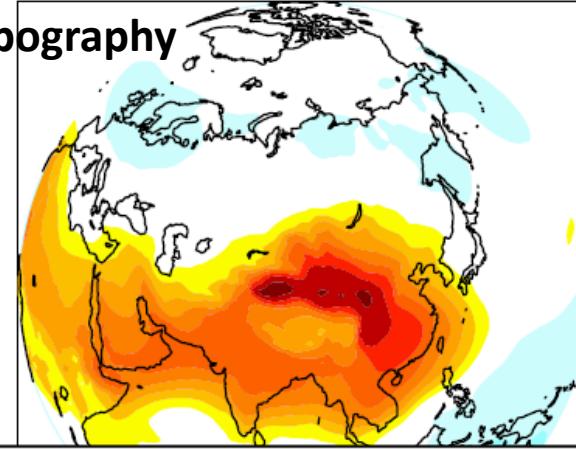
Warmer winter
with lower Tibet

Low-level synoptic fluctuation propagating around the Tibetan Plateau

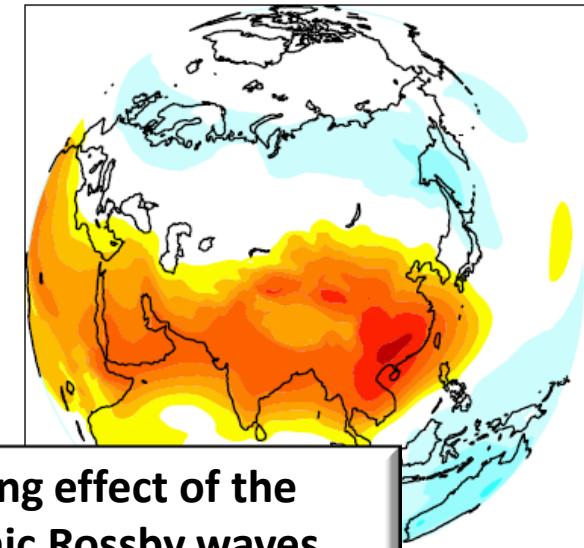
-1 day, unfiltered



+1 day, unfiltered

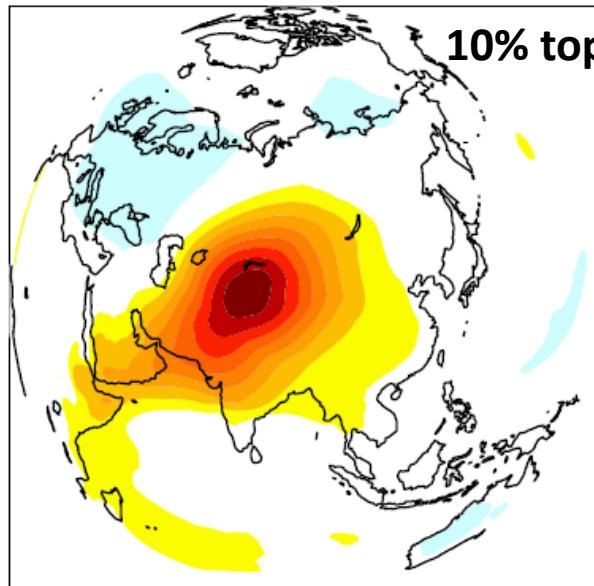


+2 day, unfiltered

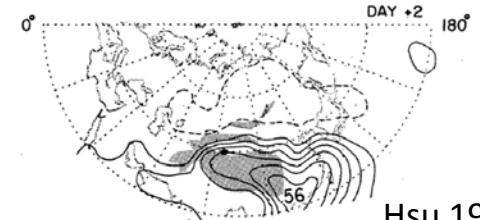
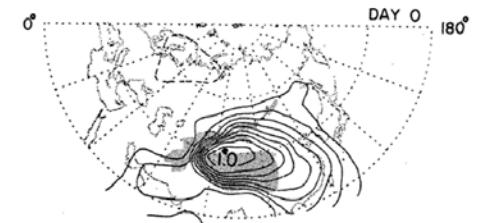
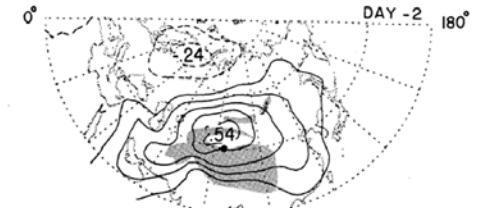
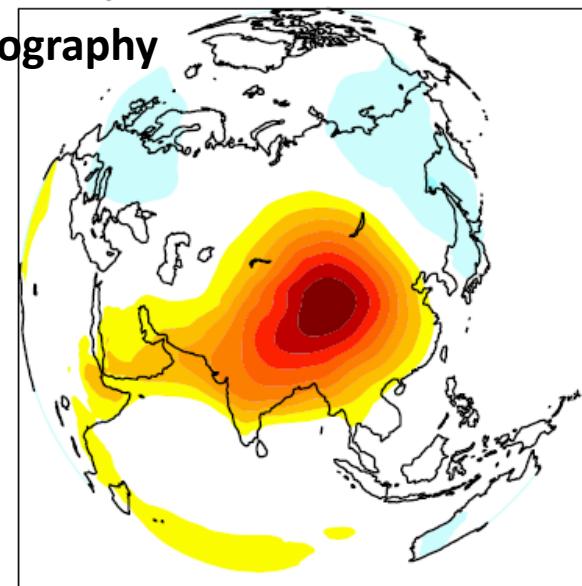


Topography-varying experiments verify the steering effect of the Tibet Plateau on circulation by exciting topographic Rossby waves.

-1 day, unfiltered



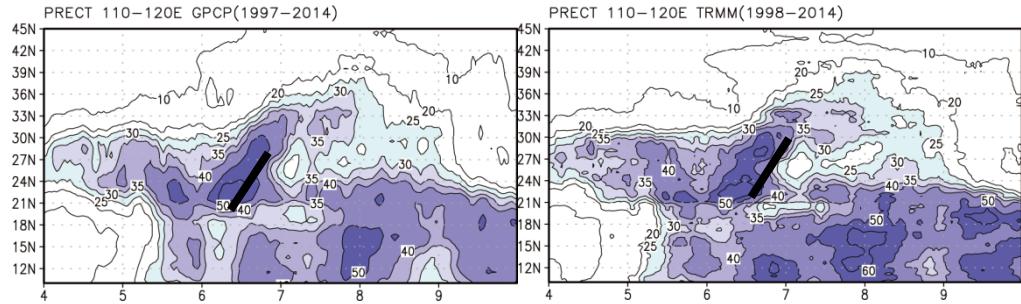
+1 day, unfiltered



Hsu 1985

Observation

GPCP(1997-2014)

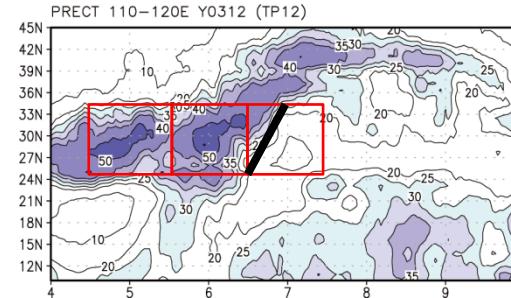


TRMM(1998-2014)

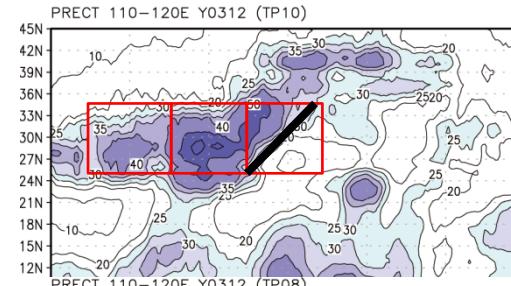
Topography enhances the seasonal *northward march* from spring to summer.
- The higher, the stronger

Simulation

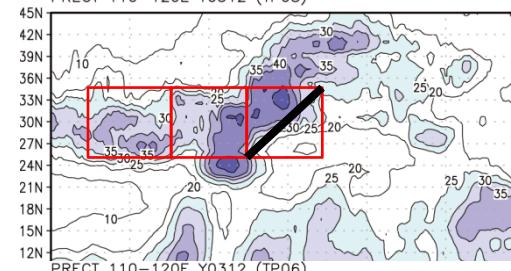
1.2



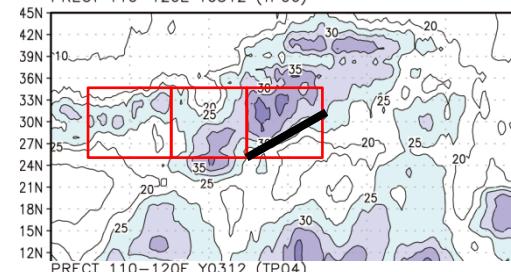
1.0



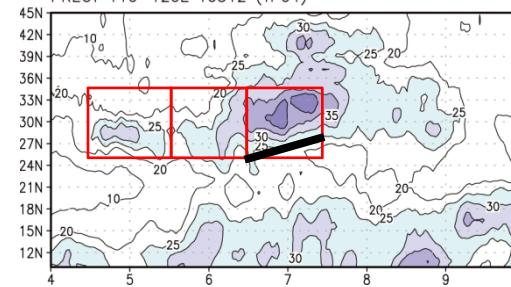
0.8



0.6

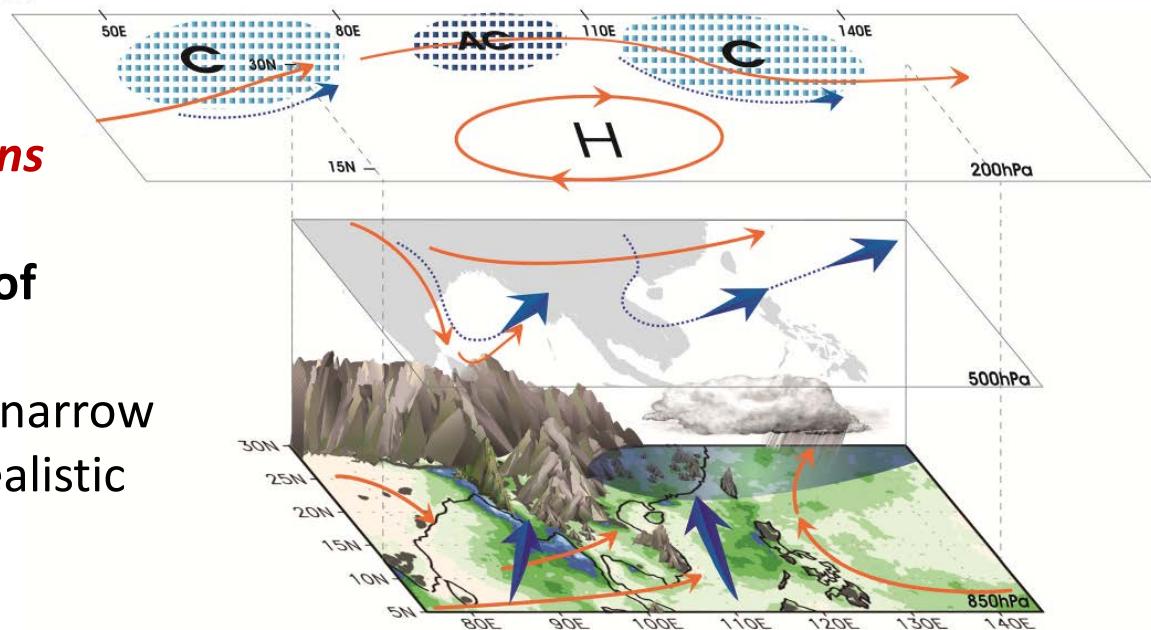


0.4



C.-H. Wu

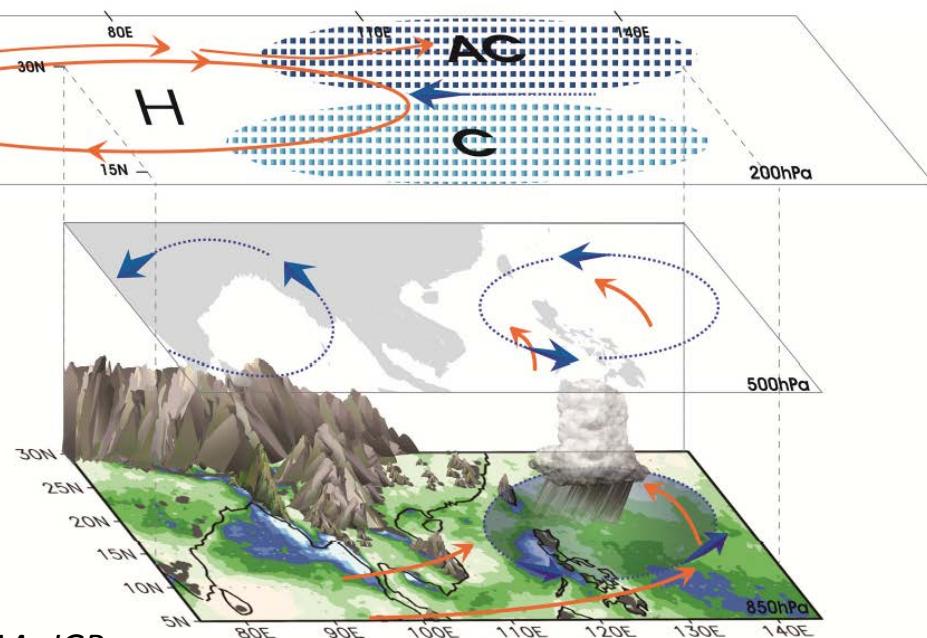
(a) LATE MAY



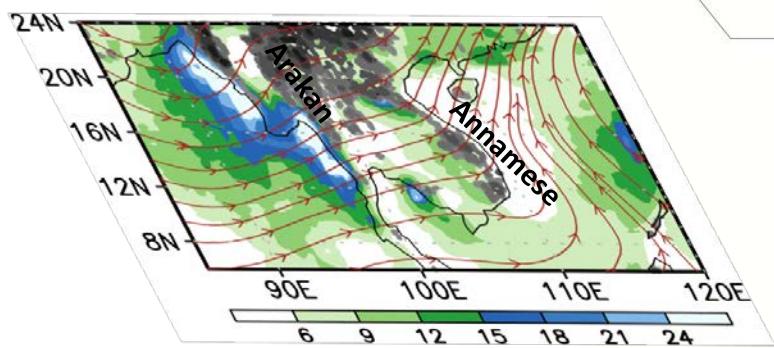
Effects of narrow mountains in the Indochina Peninsula on *subseasonal evolution* of Asian summer monsoon

- Proper representation of narrow mountains is needed for realistic simulations.

(b) LATE JULY



Annamese Mountains
Arakan Mountains



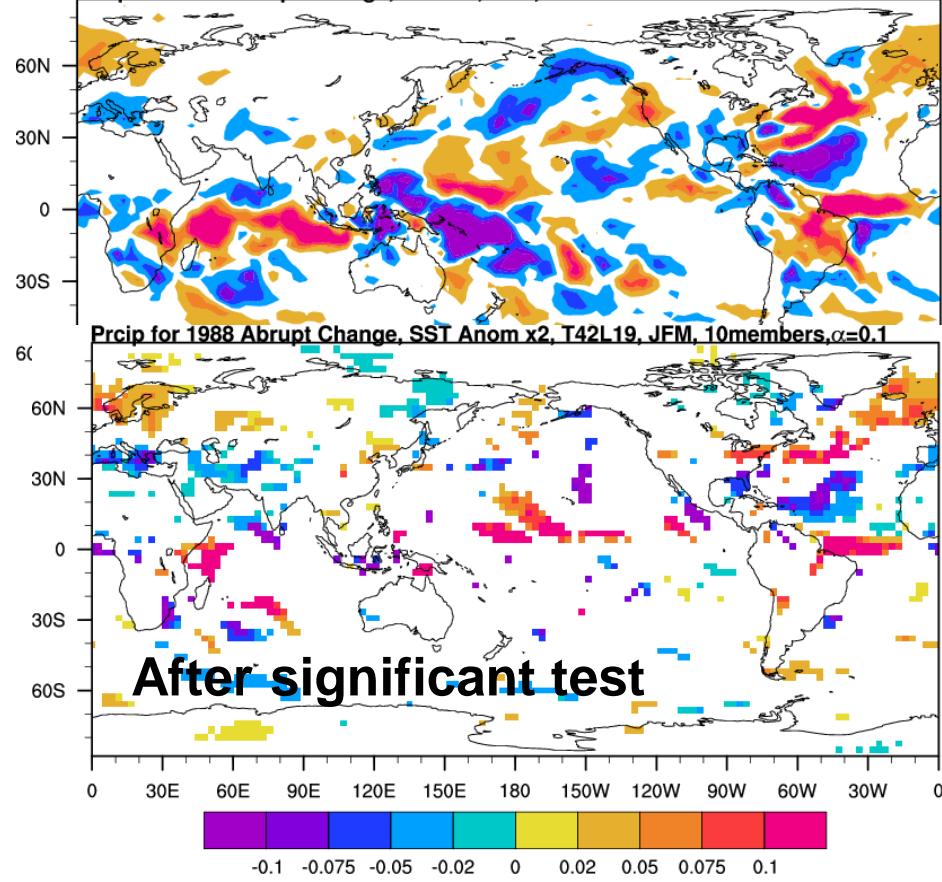
Simulation of late 1980s regime shift

- ❖ Model : ECHAM5, T42L19
- ❖ Anomaly Run SST Pattern : 1988-1997 minus 1978-1987, only over North Atlantic, **amplitude x1, x2 and x4**
- ❖ Control Run = climatological SST
- ❖ **10 (20) members** from 01Jan-20Jan 1988 initial conditions
- ❖ length of each simulation was 10 year and year 4-10 were calculated.
- ❖ Only focus on JFM

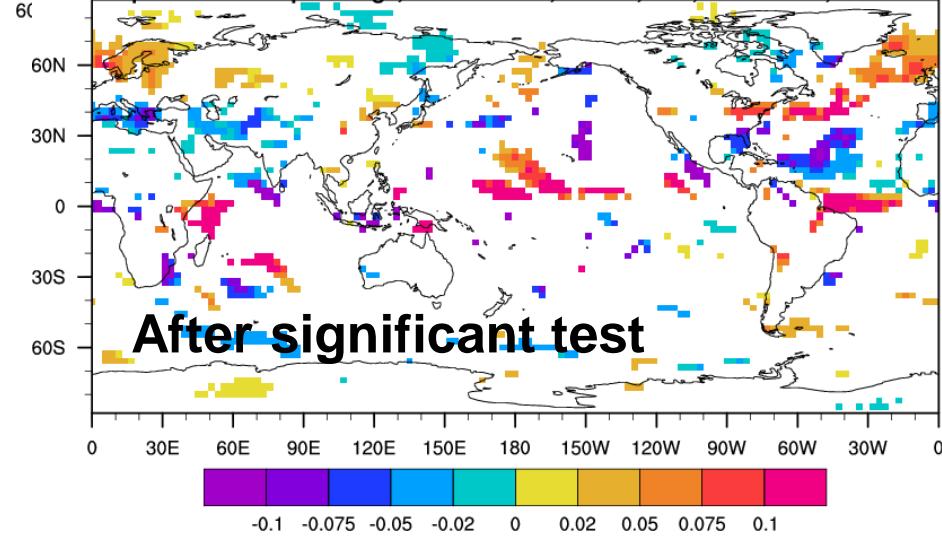
GPCP

Even precipitation?

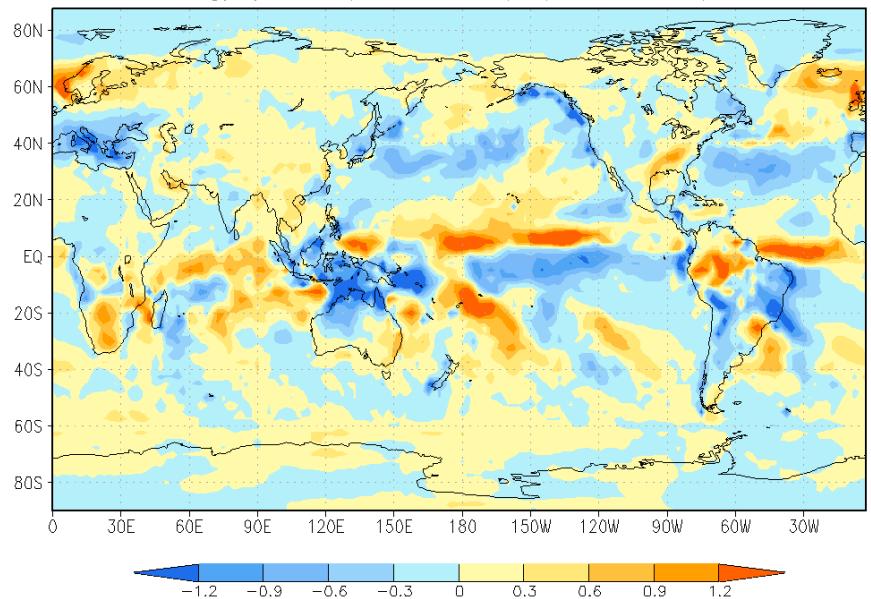
Prcip for 1988 Abrupt Change, T42L19, JFM, 10members



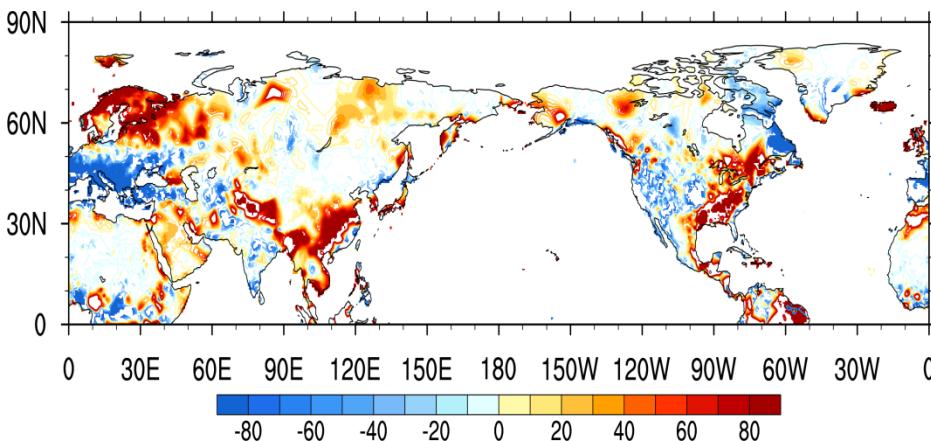
Prcip for 1988 Abrupt Change, SST Anom x2, T42L19, JFM, 10members, $\alpha=0.1$

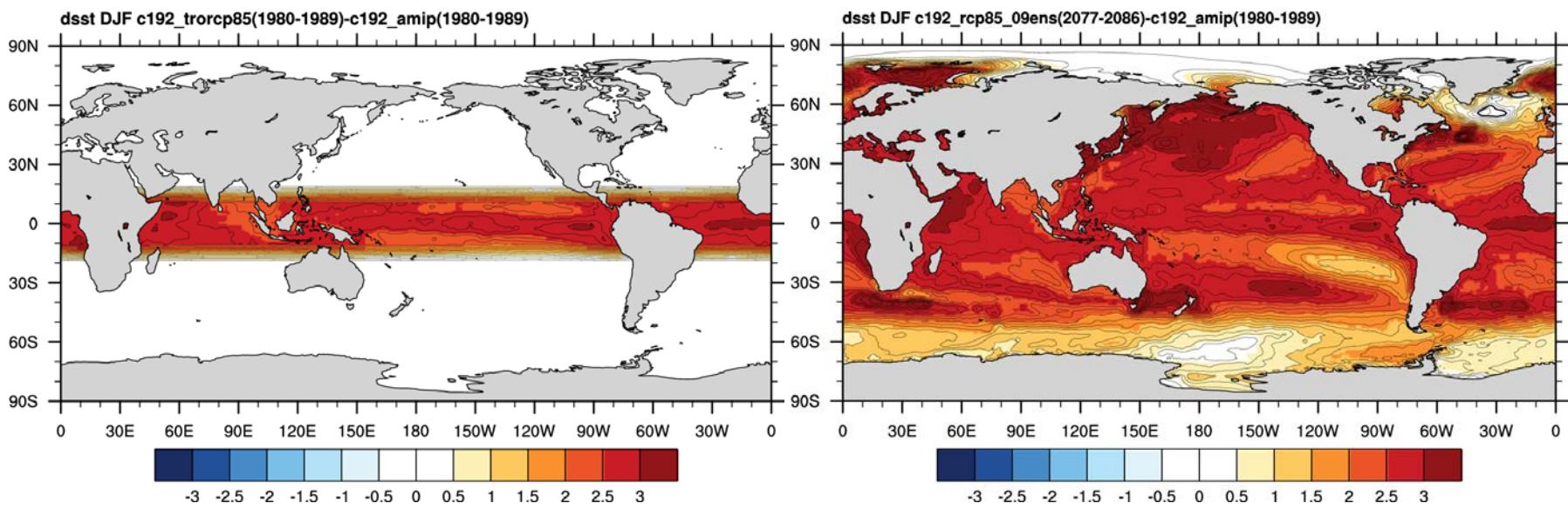
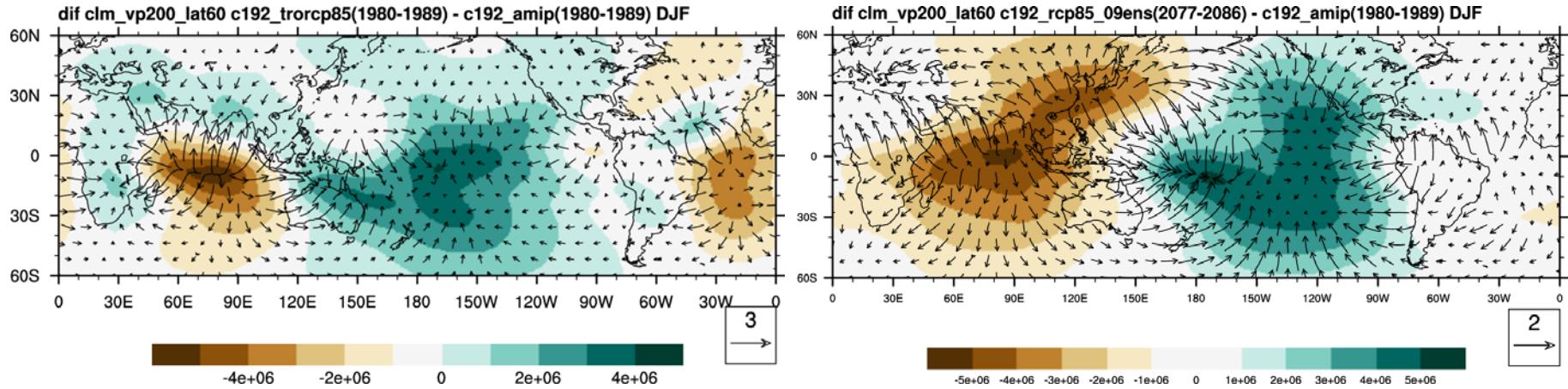


gpcp-1-3(1988-1997)-(1979-1987)

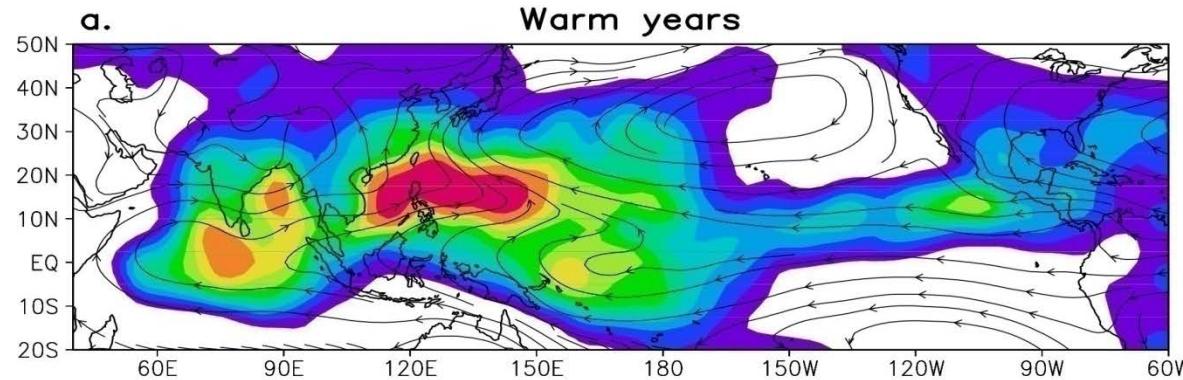


CRU

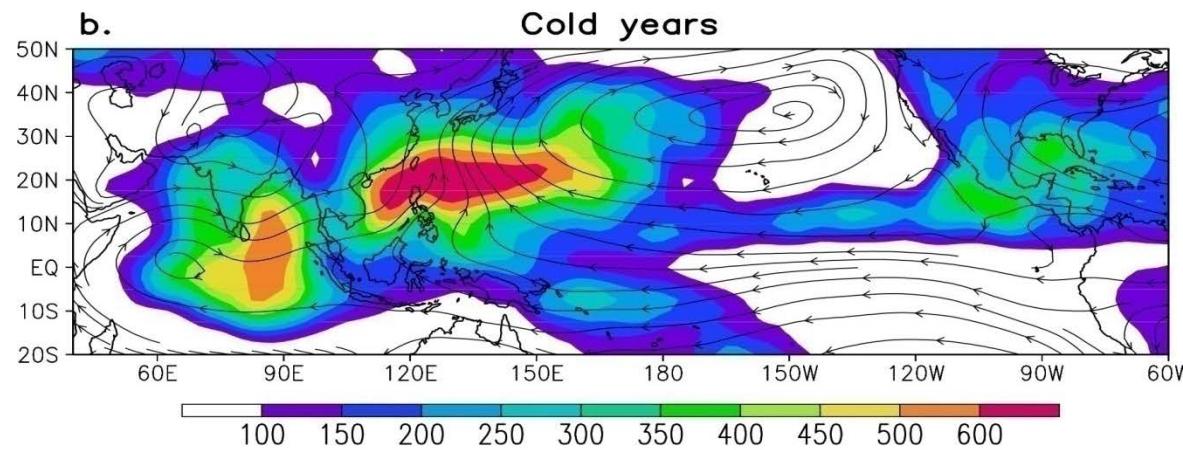




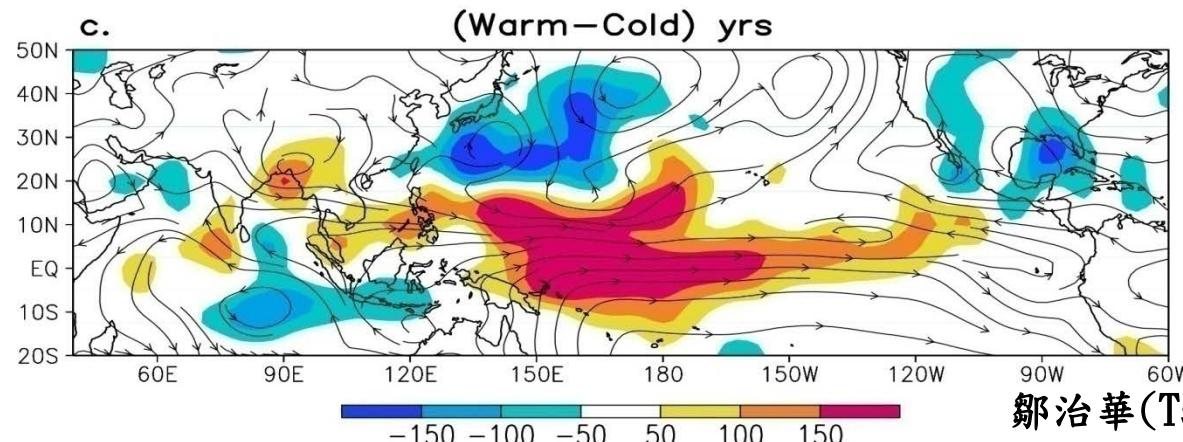
30-60day OLR var. and 850hPa streamline (JAS) composite



El Niño



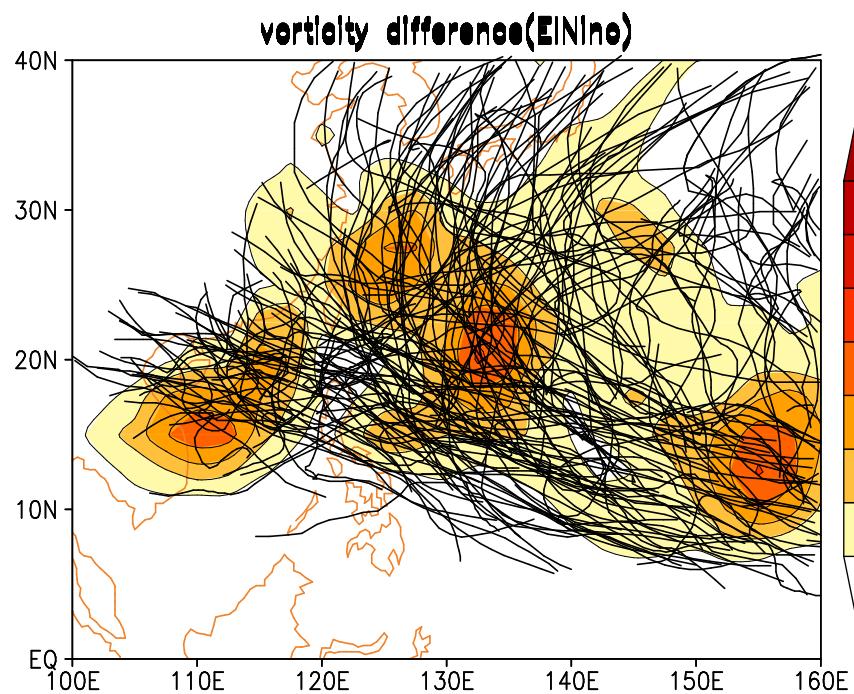
La Niña



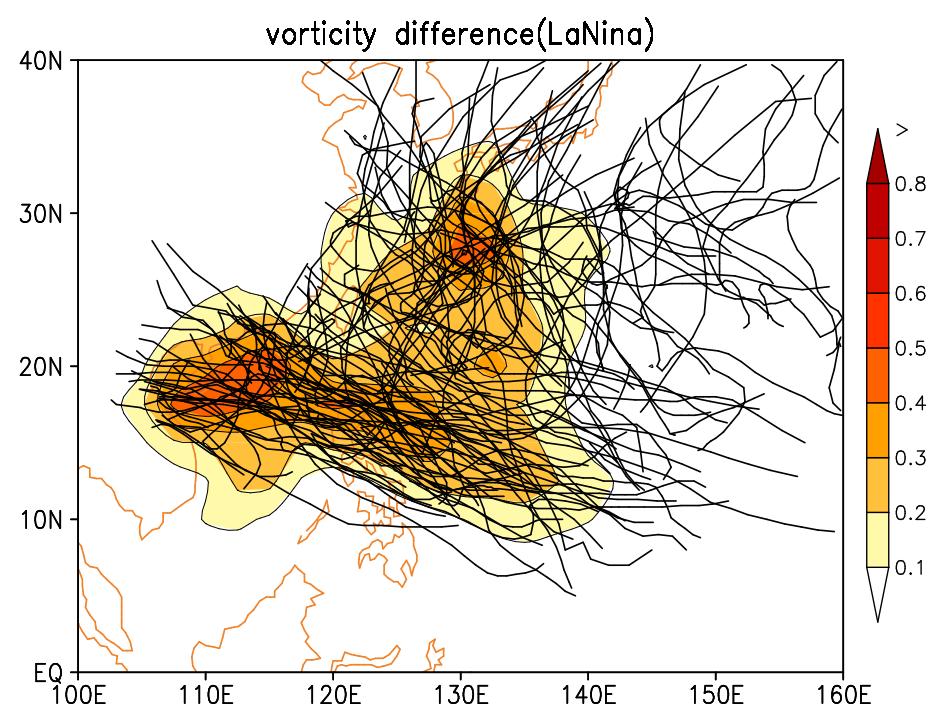
鄒治華(Tsou)、徐邦琪(Hsu)

TC also modulated by ENSO

El Nino



La Nina

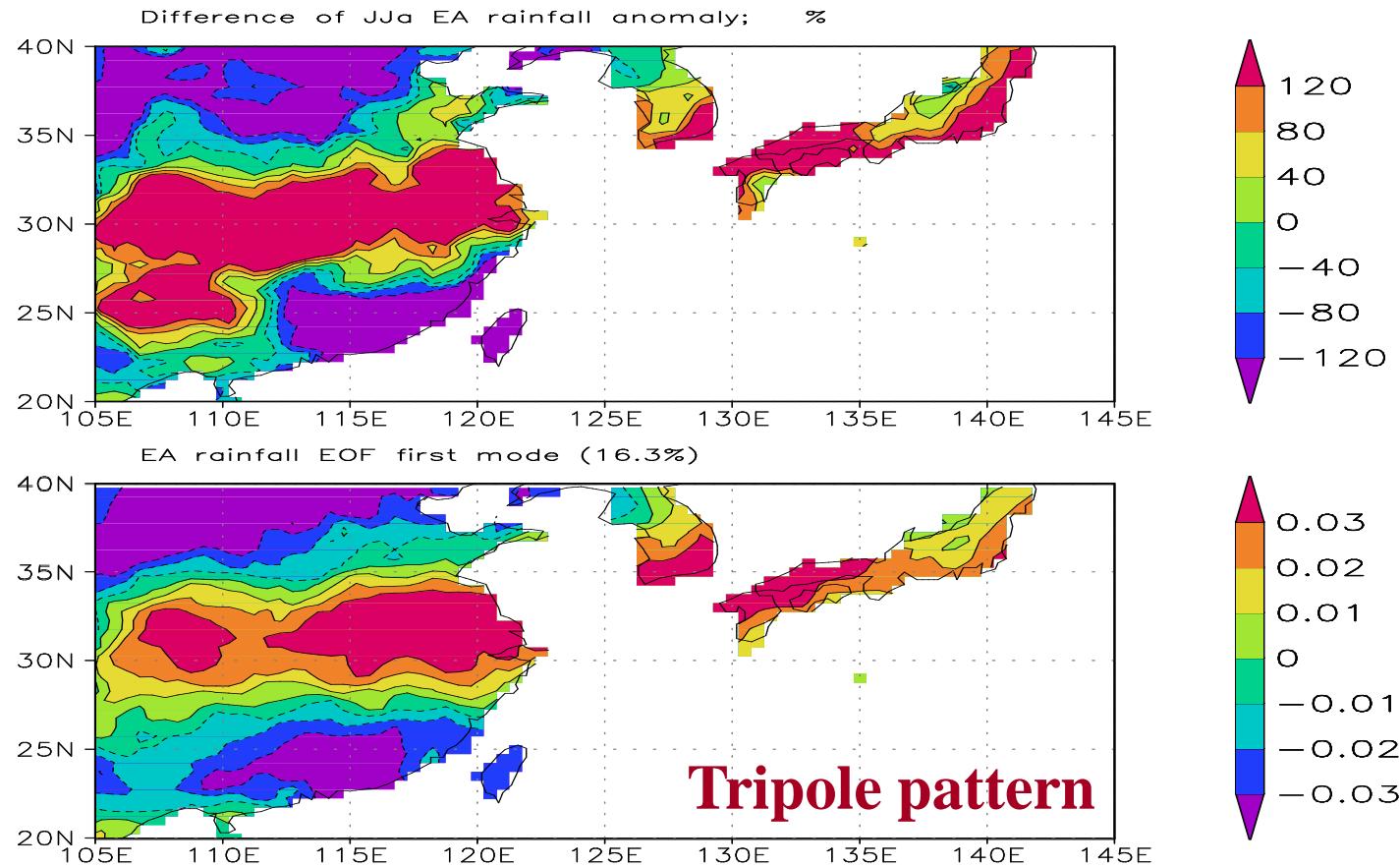


Multiple time scale nature

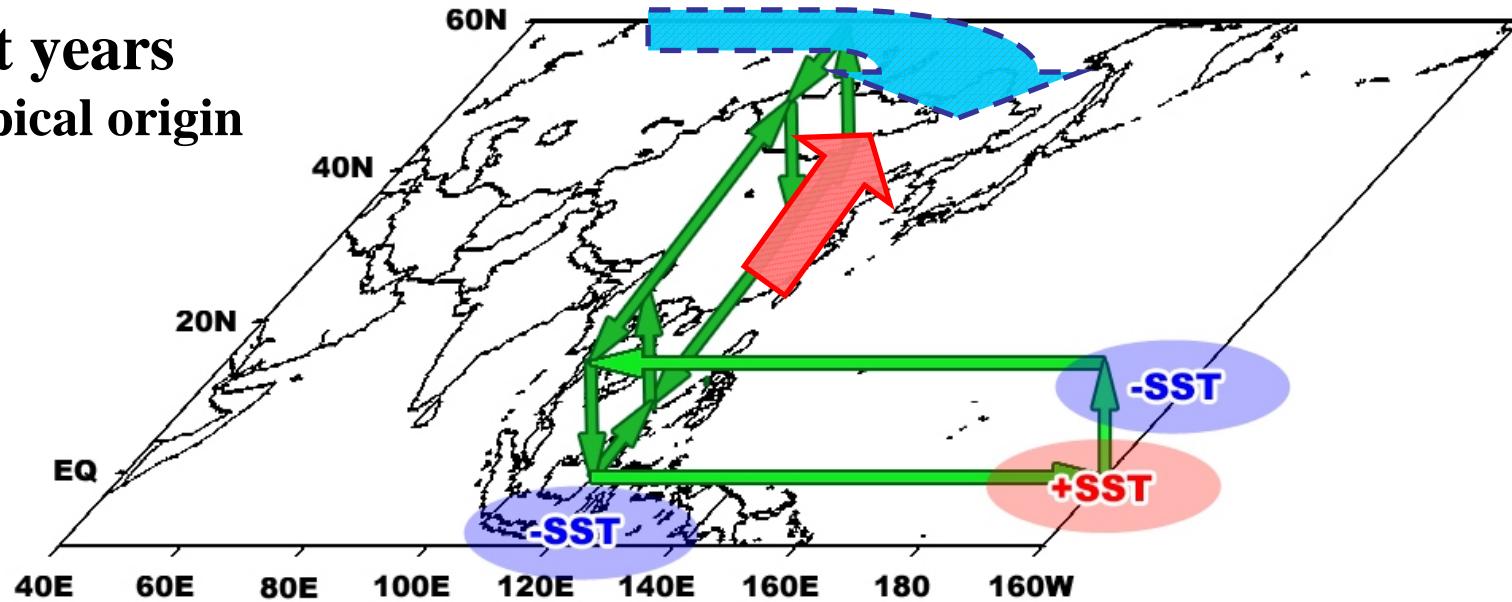
- Secular trend + Interdecadal-decadal variation + Interannual variation + Seasonal evolution + Intraseasonal and sub-seasonal variation + ..
- Example :
 - decadal vs. abrupt change (oscillation vs. shift)
 - decadal+interannual
 - Intraseasonal Oscillation (ISO, MJO, 季內震盪) vs. monsoon onset
 - TC vs. ISO, interannual, decadal-interdecadal

Also related to diabatic heating over Tibet

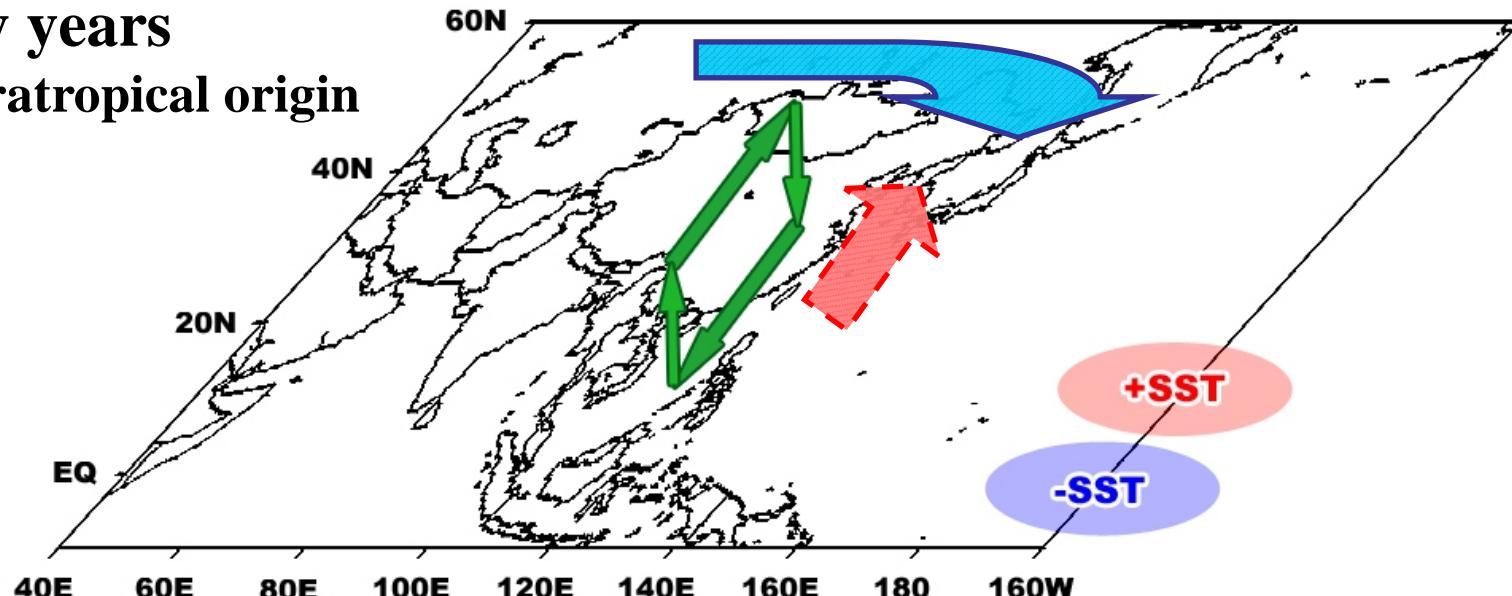
Rainfall Composites (8 strong cases - 8 weak cases)



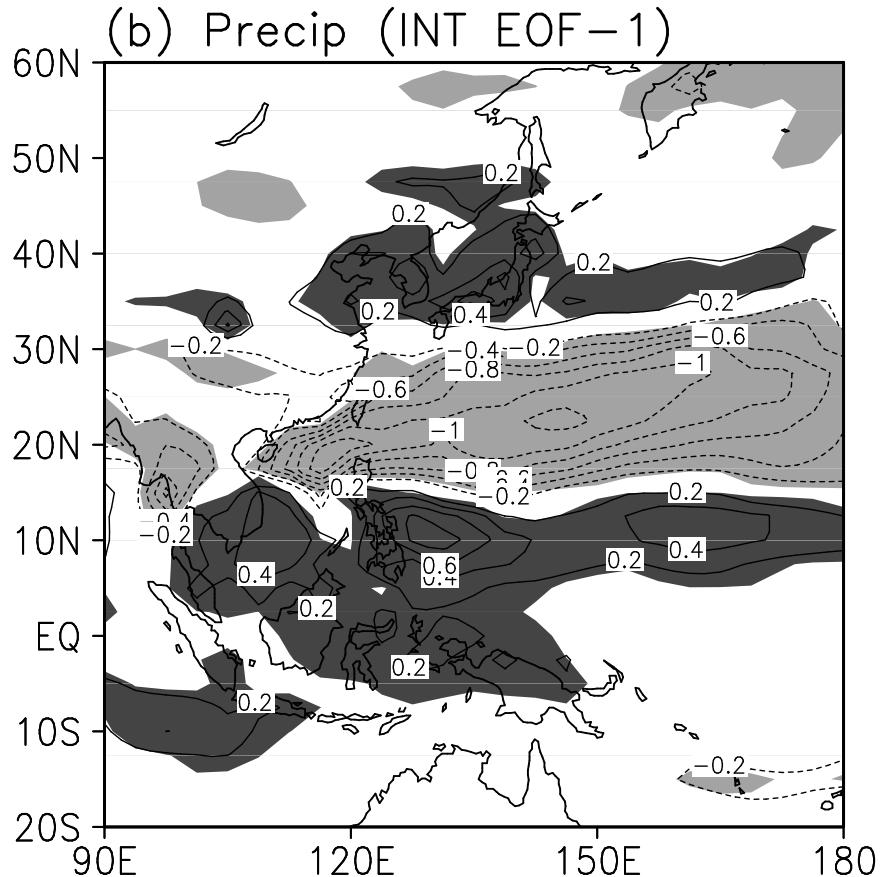
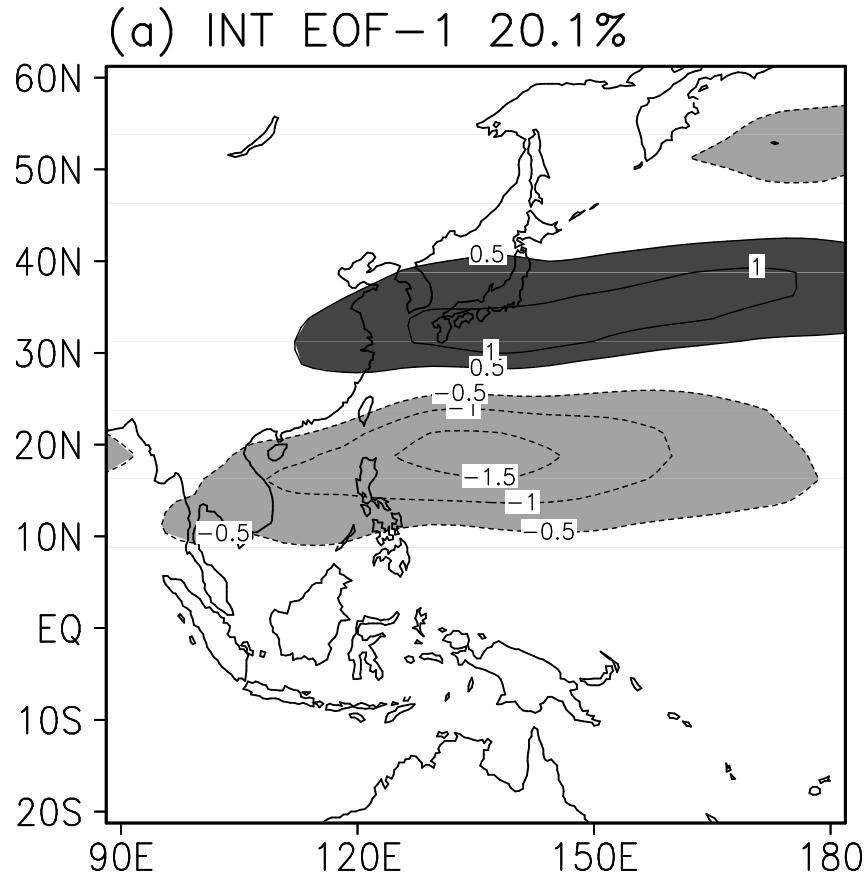
Wet years
Tropical origin



Dry years
Extratropical origin

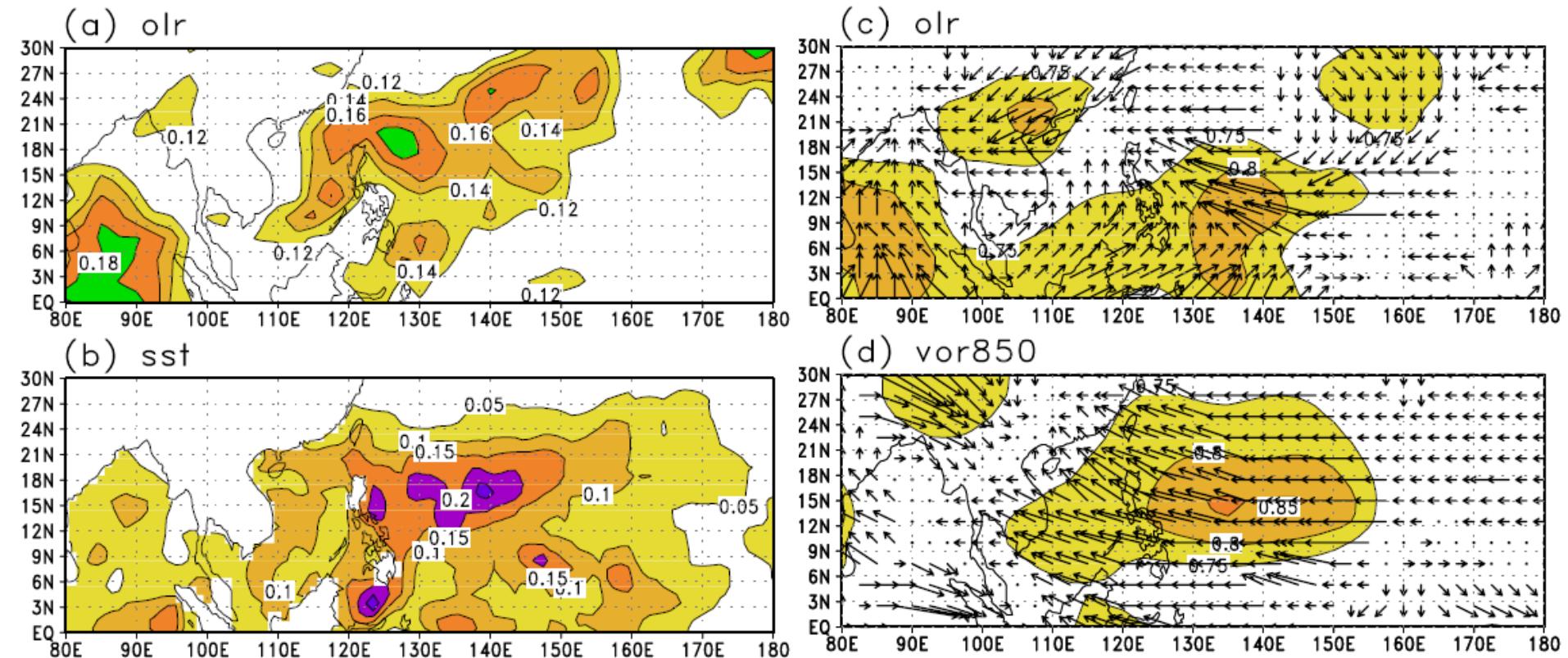


intrinsic mode of mean flow? can be triggered by various forcing?



(Lu et al., 2006, JMSJ)

Multiscale Interaction in the Tropical Western North Pacific Hot Spots of Intraseasonal Oscillation (Hsu, 2005, 2010)

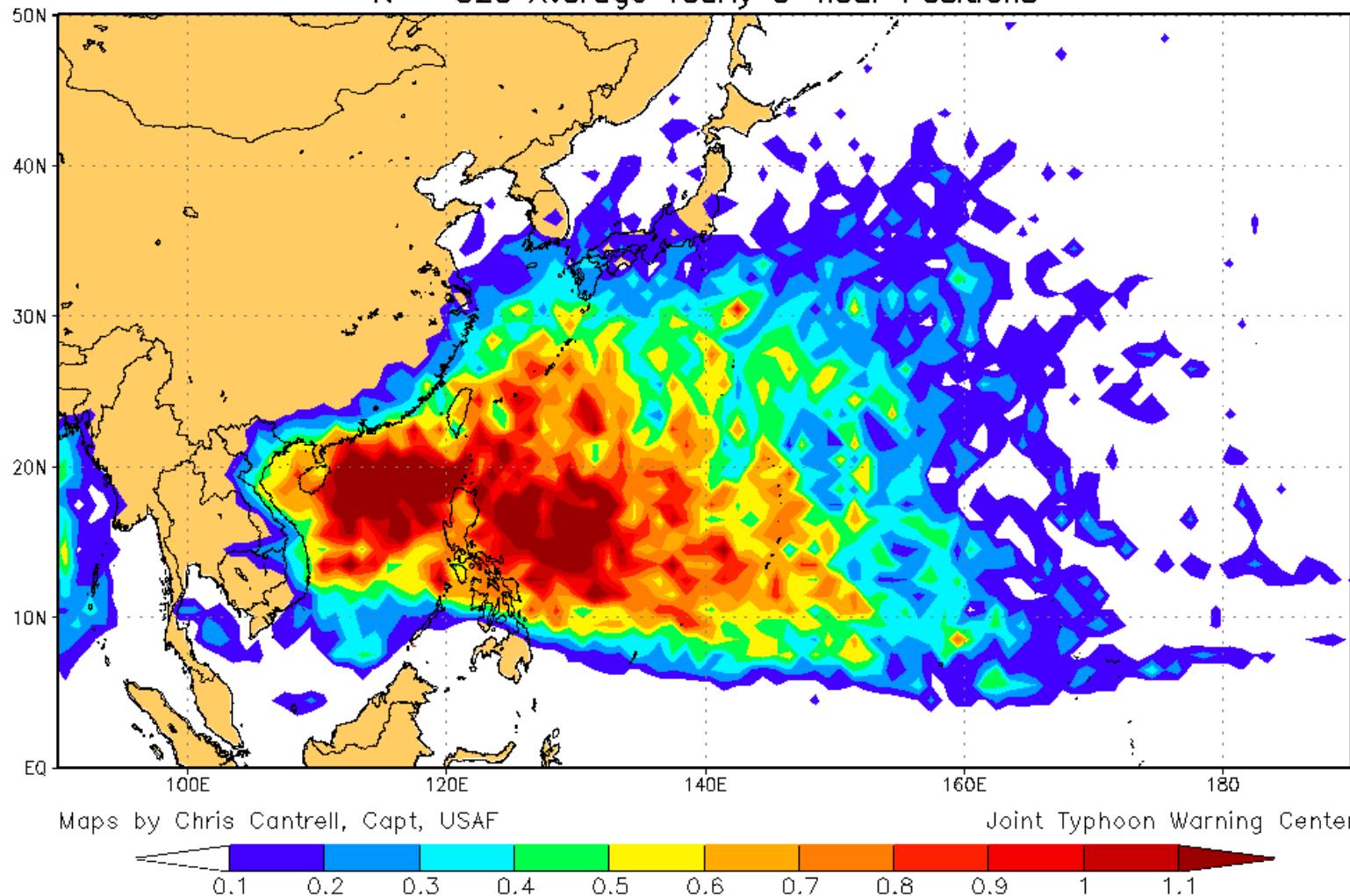


Hsu and Weng, 2001, J. Climate

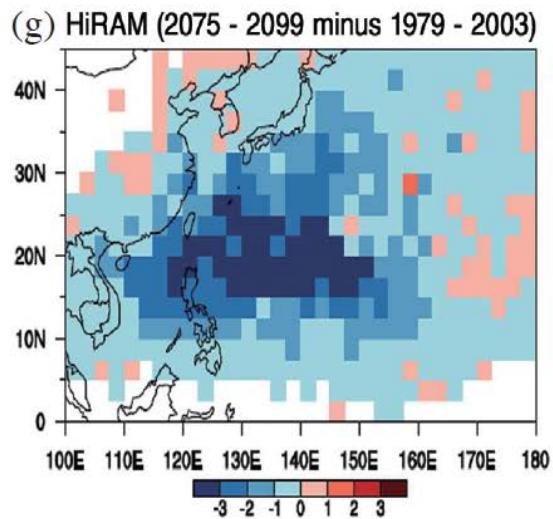
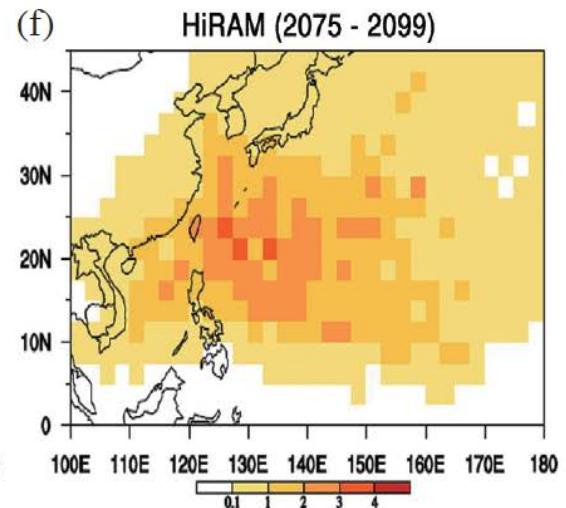
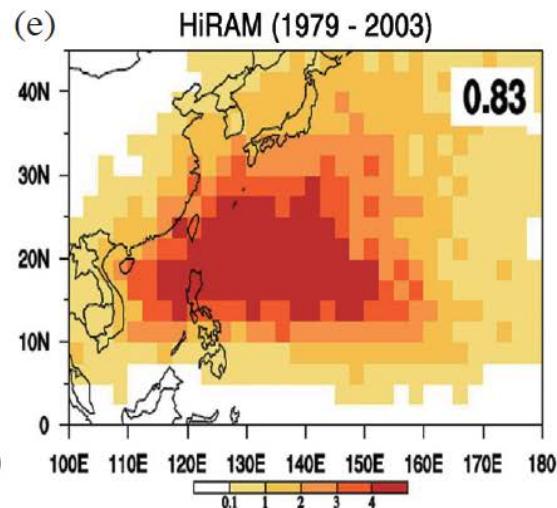
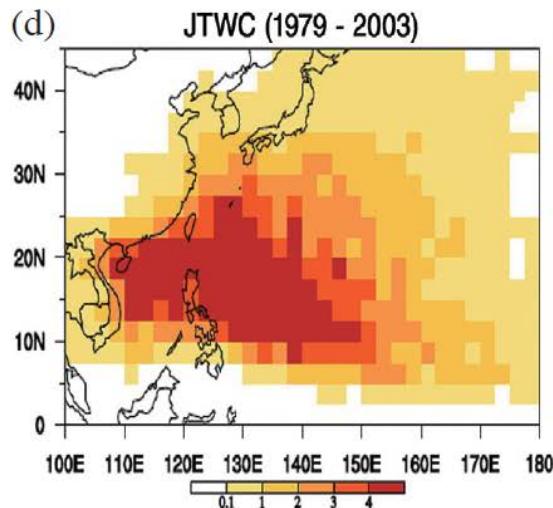
Breeding and Play Ground of TCs

Average Annual TC Occurrence for 1972 to 2001

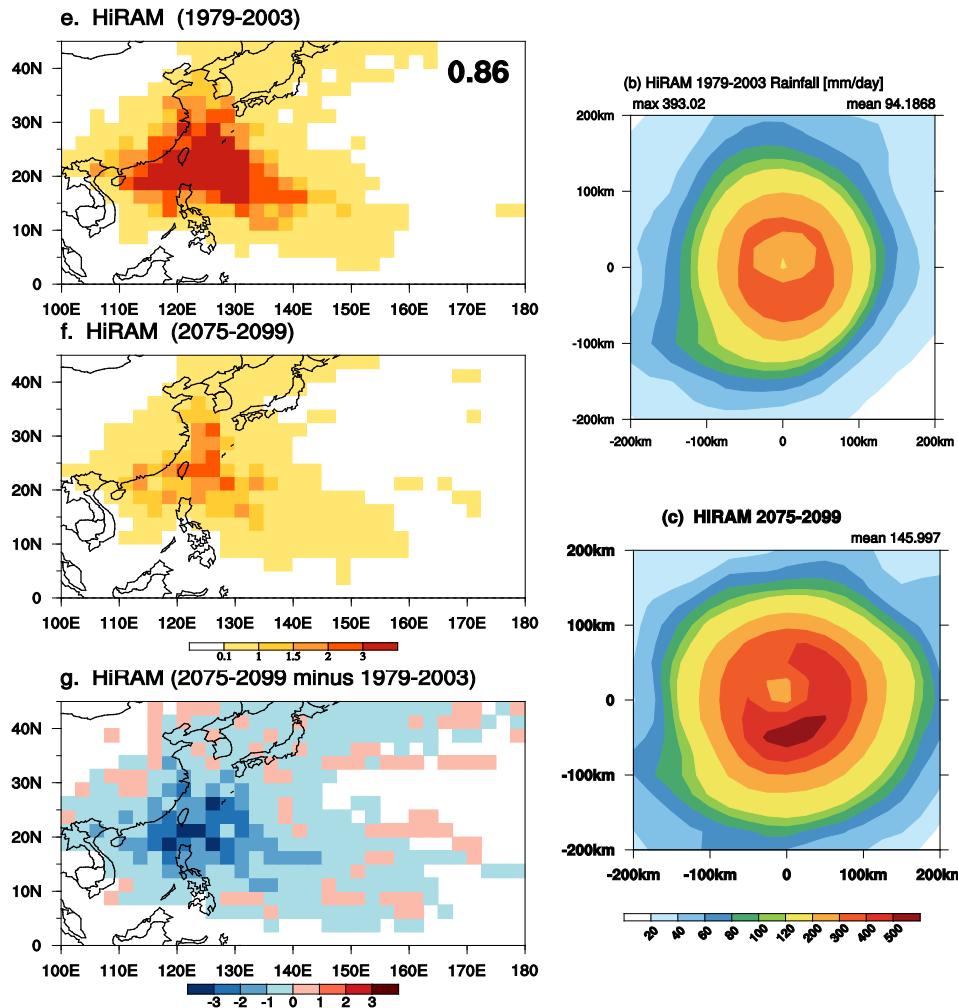
N = 826 Average Yearly 6-hour Positions



TS Frequency (JJASON)



Changes in typhoon (affecting Taiwan) frequency and strength

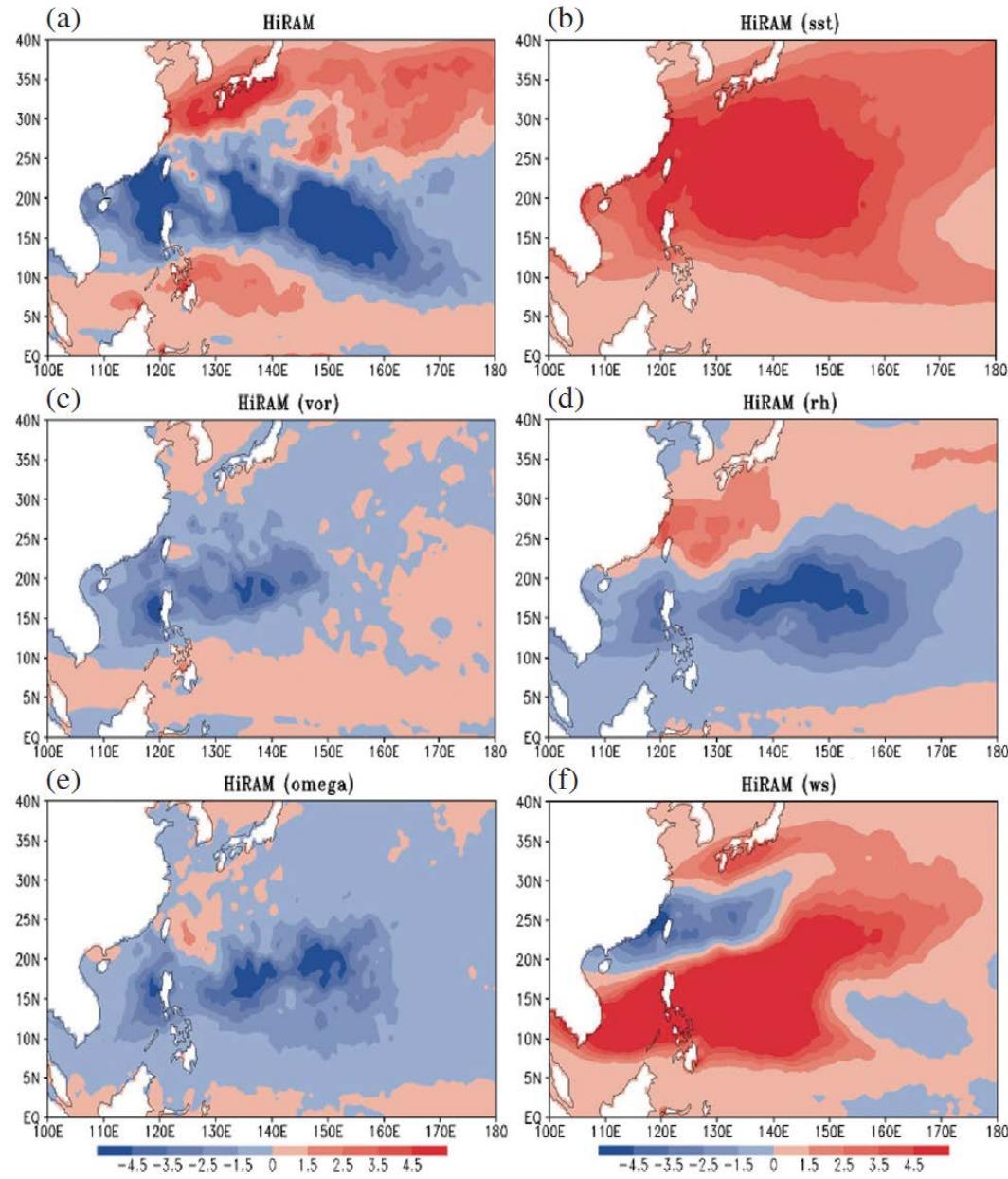


- ◆ Significantly reduced in number but strengthened
- ◆ Rainfall within 200 and 100 km radius increased by 44 and 20 %

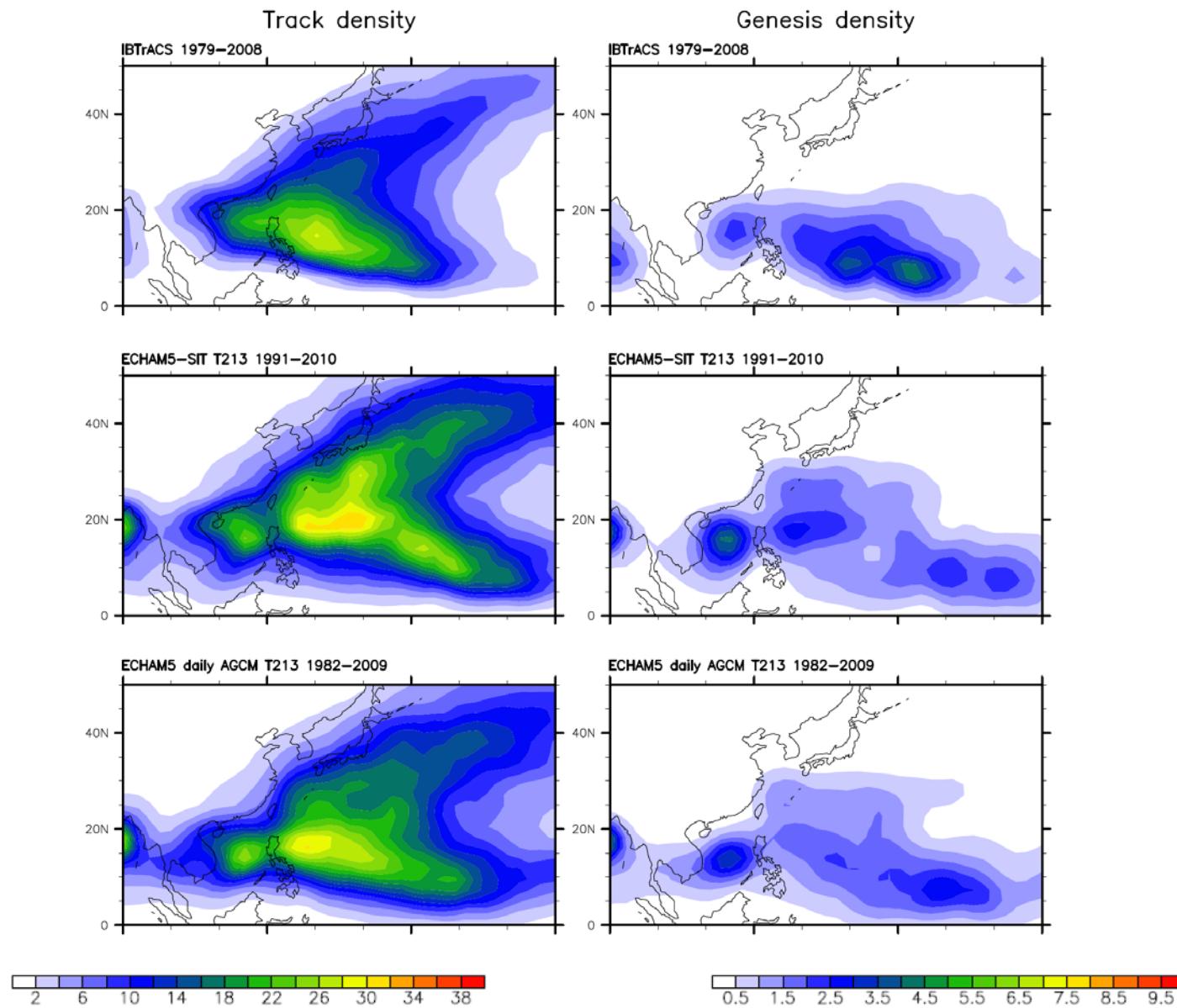
Tsou et al., 2016, TAO

FIG. Distribution of TWCN TS track frequency (shaded, unit: numbers per 2.5° longitude/latitude per season period) in JUN.-NOV.

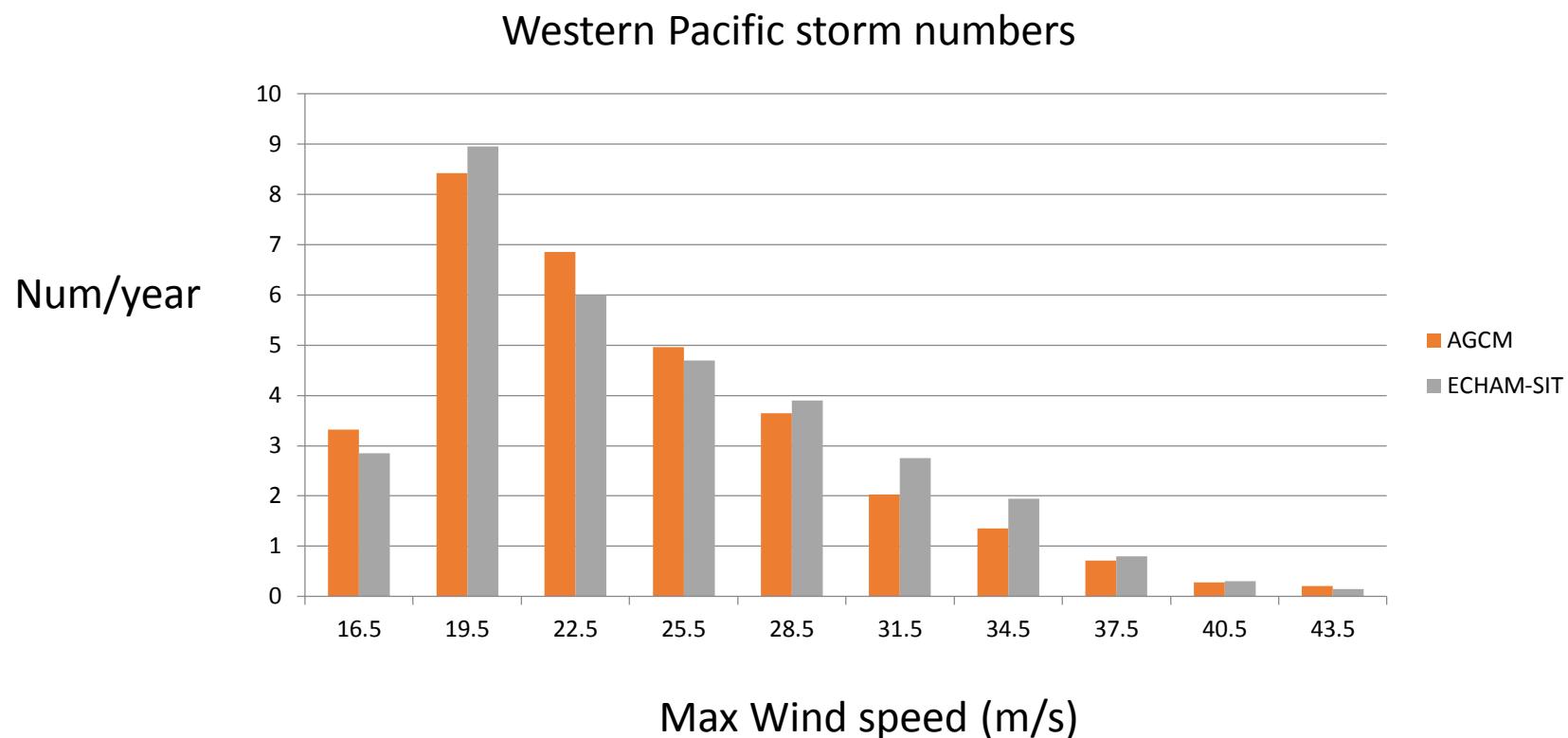
GPI Changes

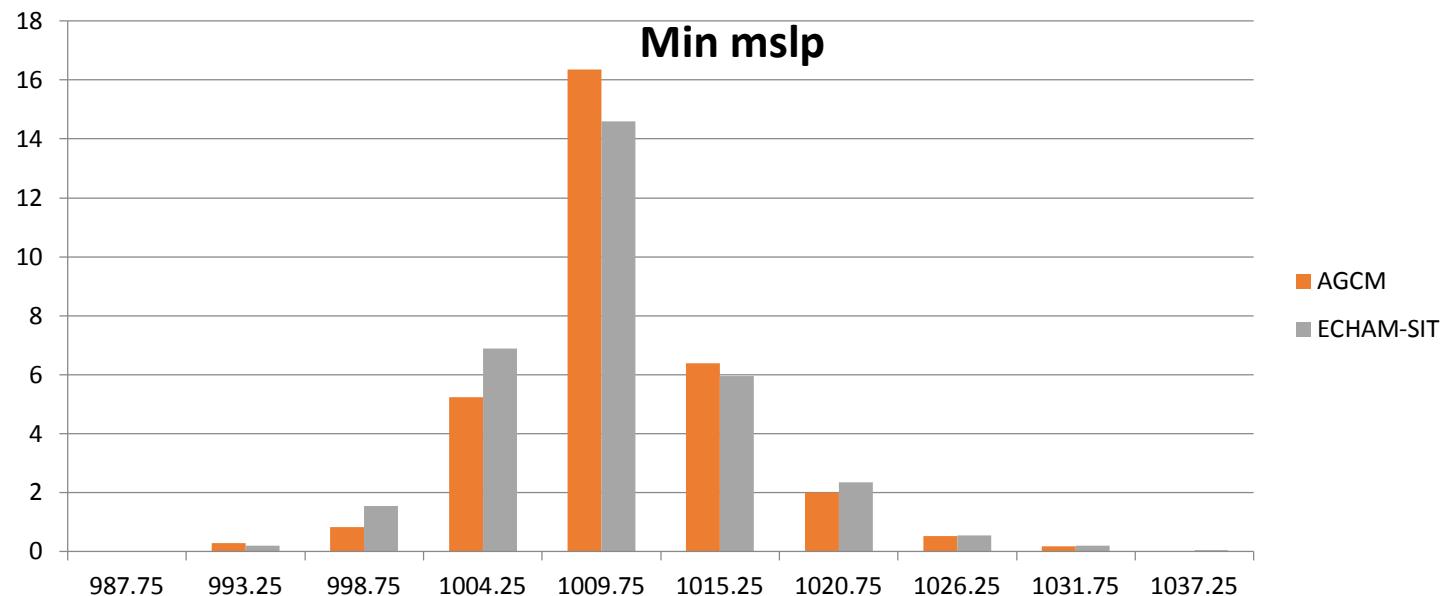
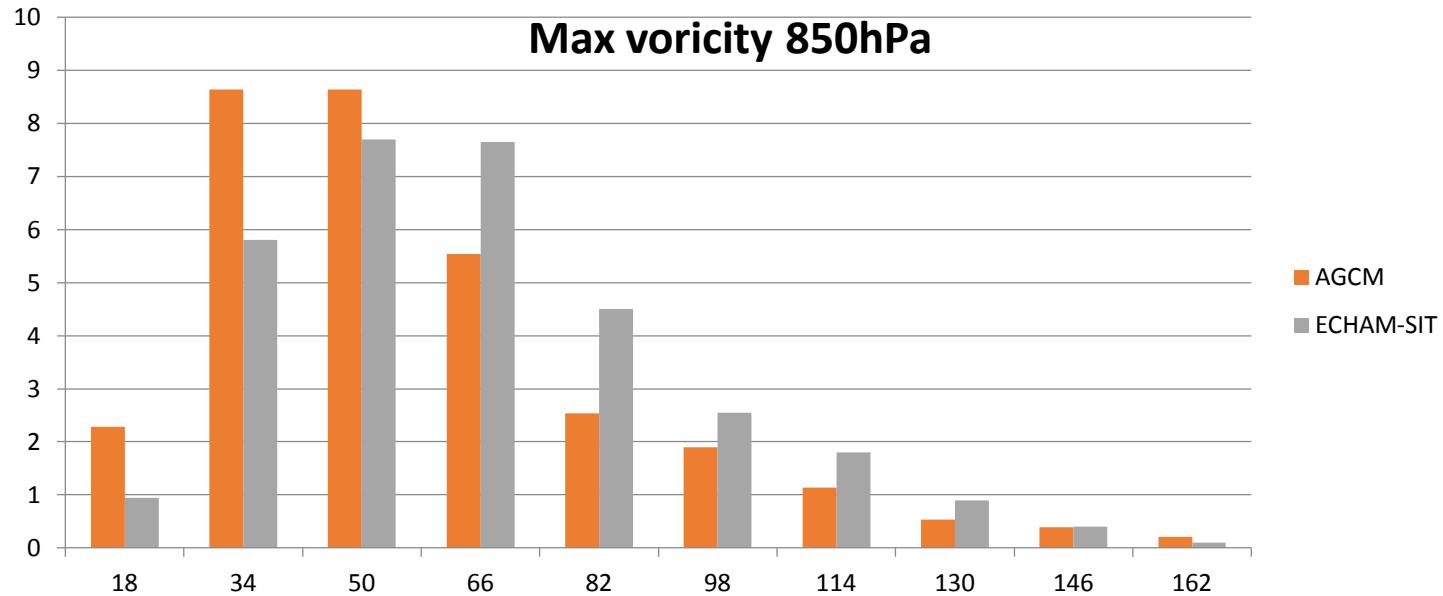


ECHAM5 T213: AGCM (uncoupled) vs. AGCM-SIT (coupled)

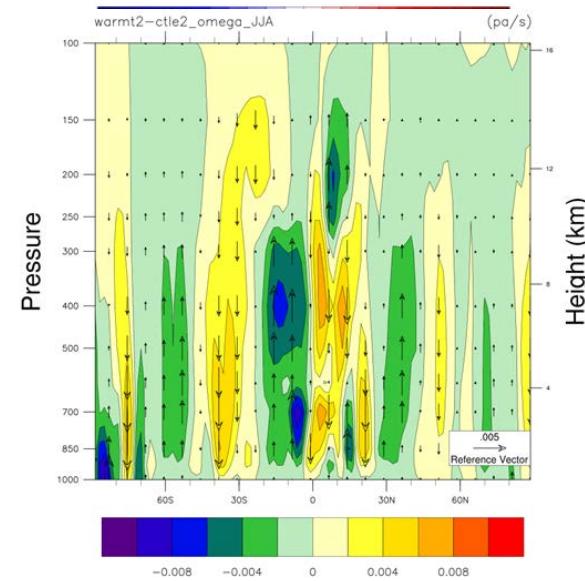
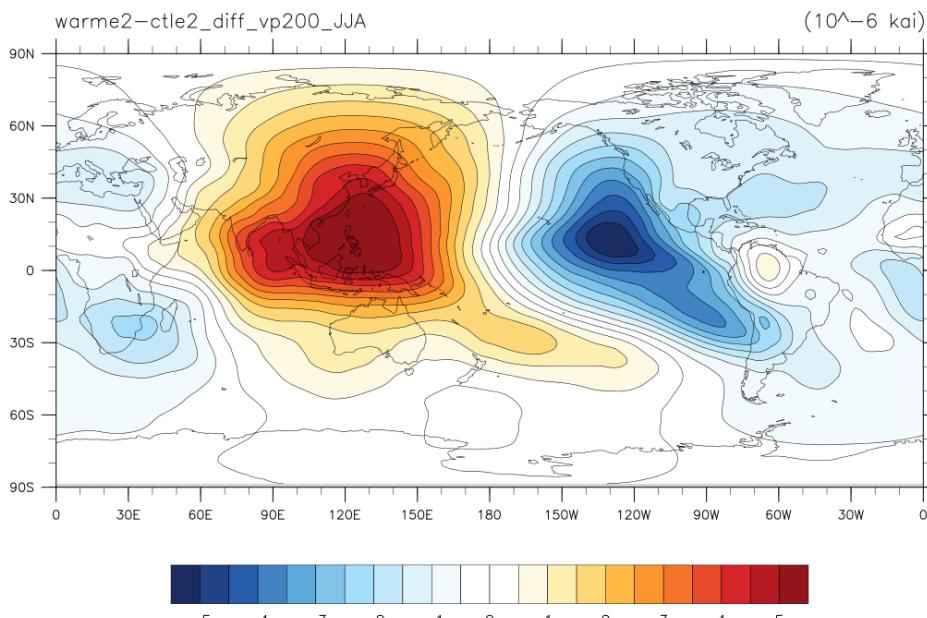
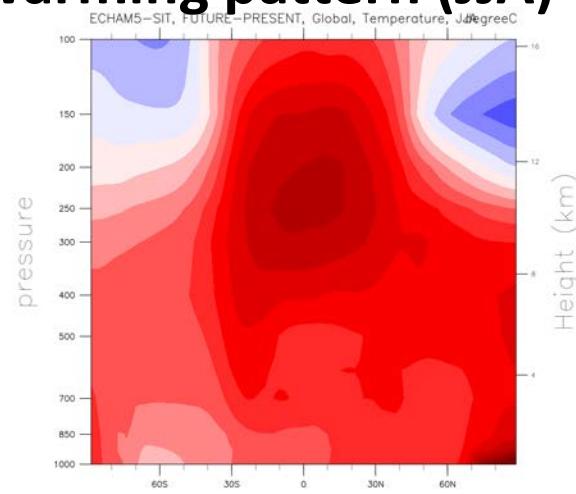
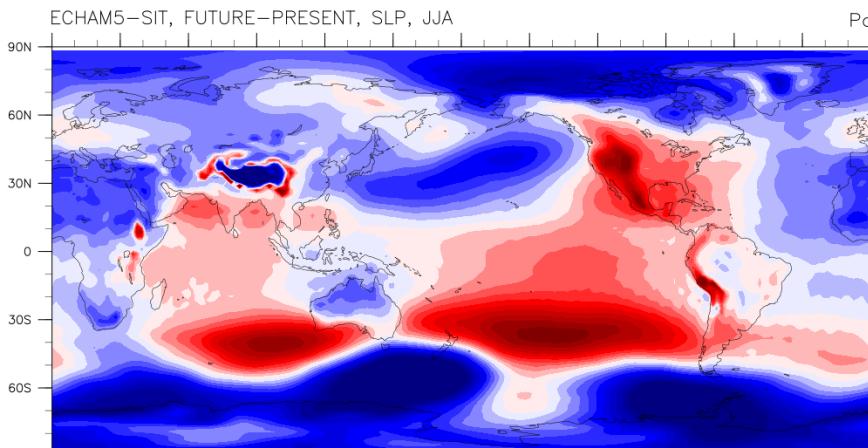


	AGCM (daily OISST) 28 years	ECHAM5-SIT 20 years
NH	74.5	77.1
WP	31.8	32.4

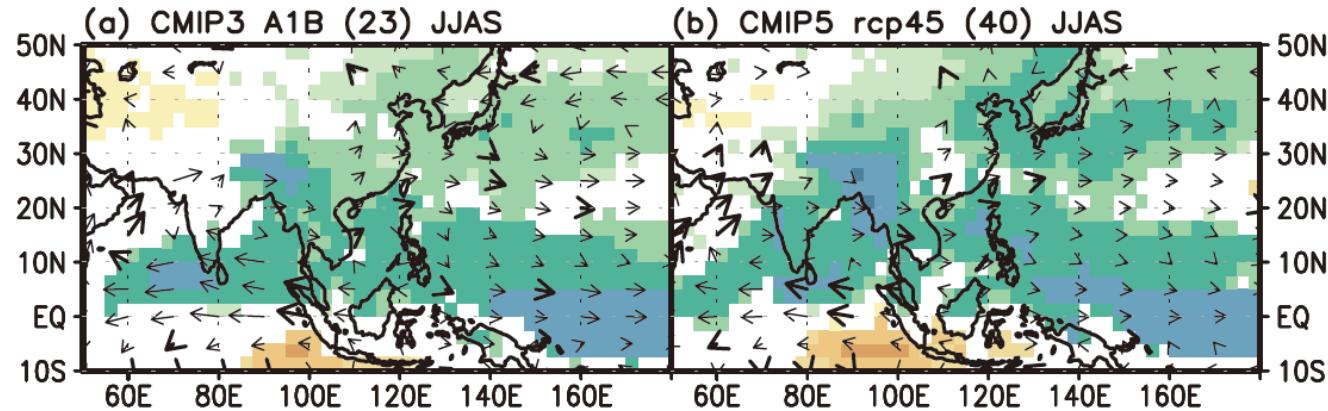




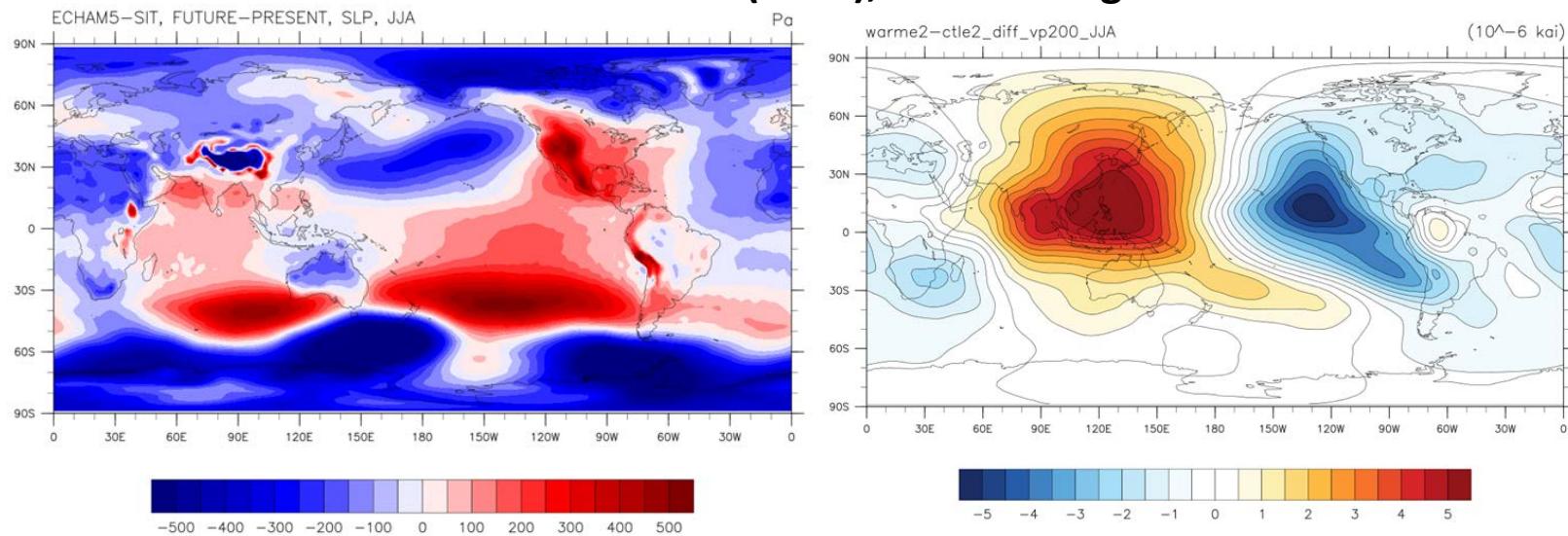
Similar response in another RCP8.5 warming simulation using ECHAM5-SIT with different SST warming pattern (JJA)

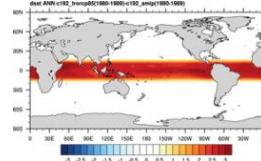
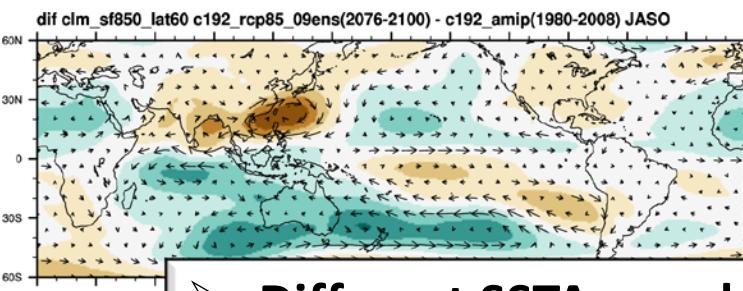
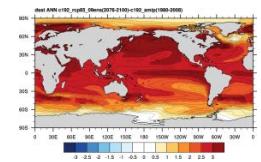


Comparison with other models (not resolving TCs)

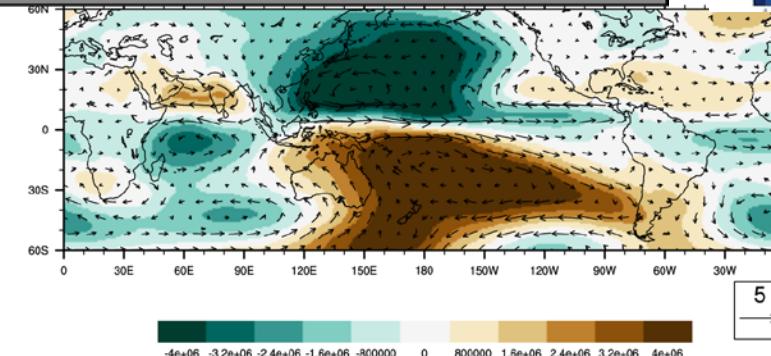
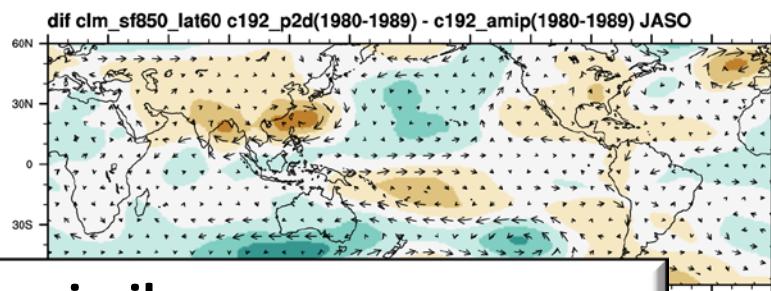
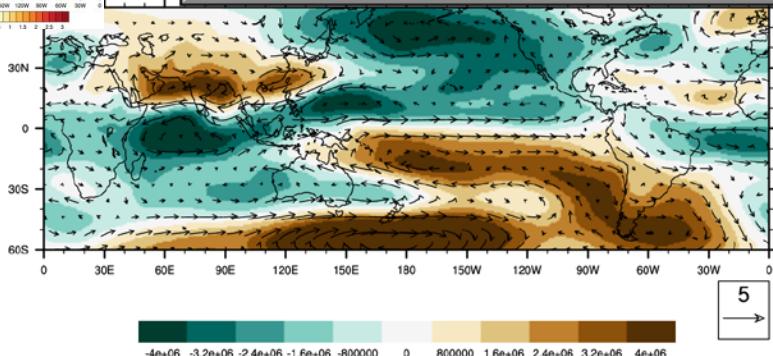


ECAHM5-SIT (T213), TC resolving





- Different SSTAs produce similar response
- but not El Niño-like SSTA
- Relatively insensitive to regional structure of SSTA



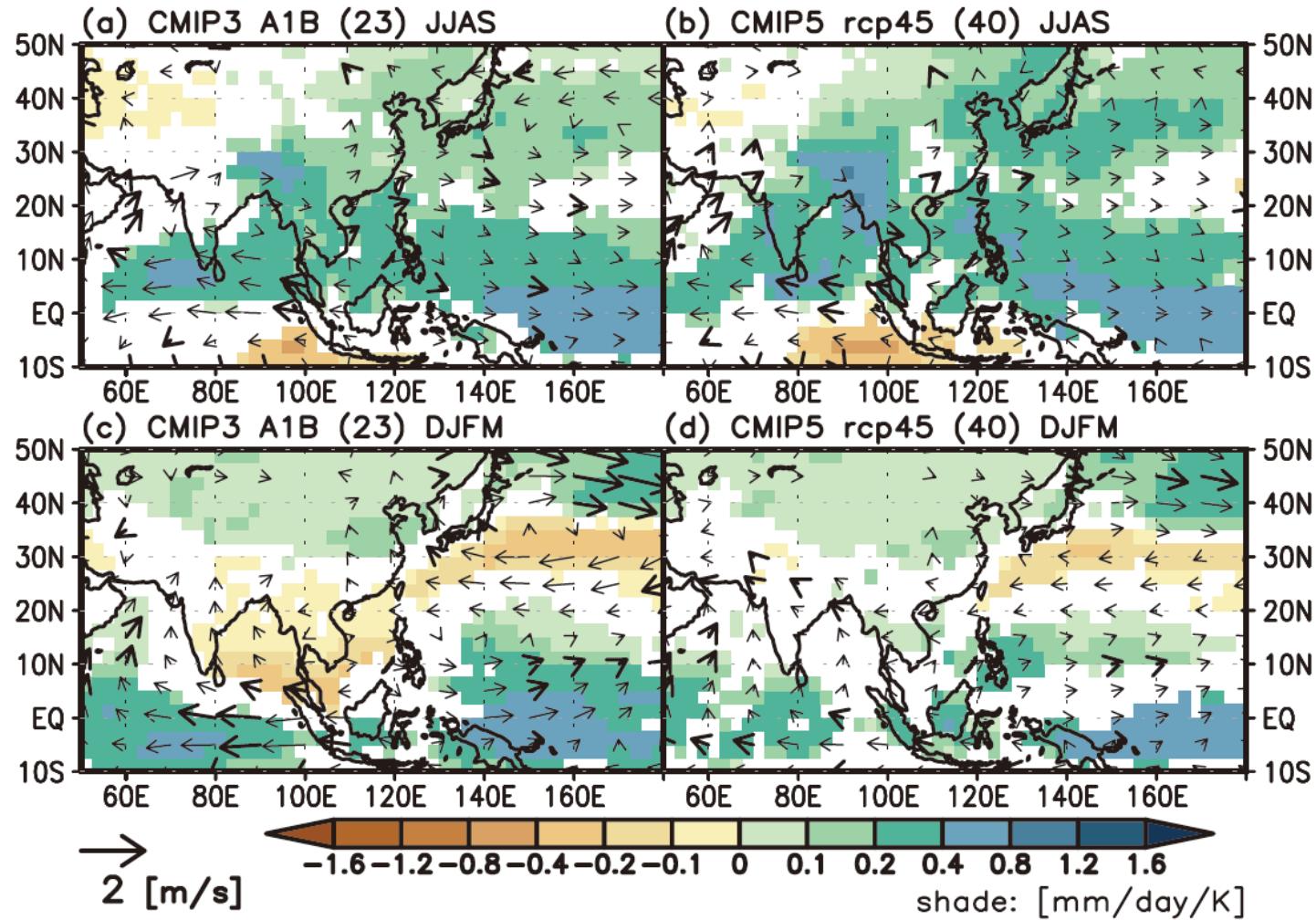
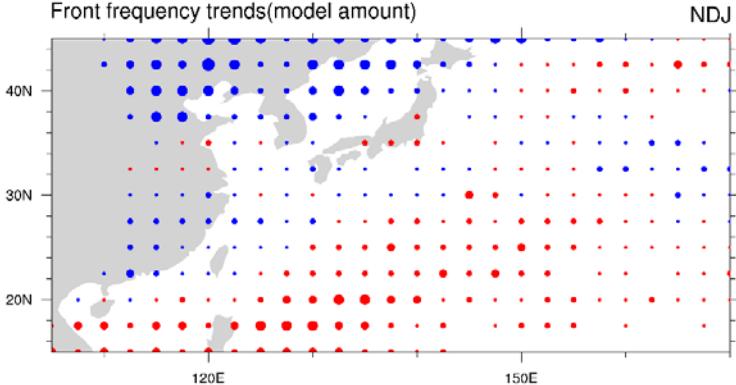


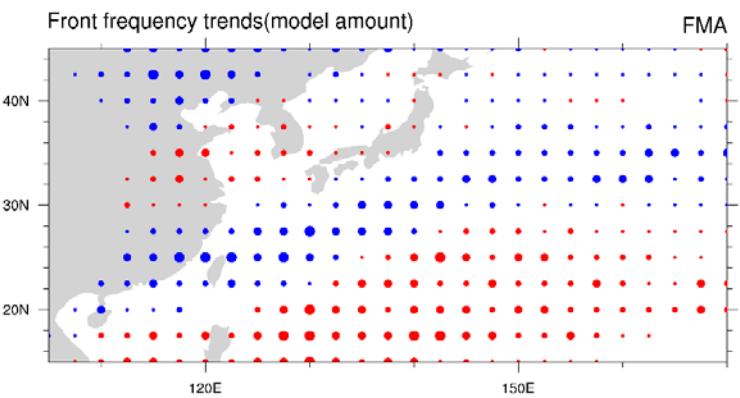
Fig. 8. Future changes in (a, b) June–September, and (c, d) December–March mean precipitation and 850 hPa wind fields. (a, c) A 23-model CMIP3 MME under the SRES A1B scenario, and (b, d) 40-model CMIP5 MME under the RCP4.5 scenario with a common 2.5° by 2.5° grid. Grid points where two-thirds of models agree on the sign of the changes are shaded for precipitation and shown in bold vectors for both zonal and meridional winds.

Front frequency change

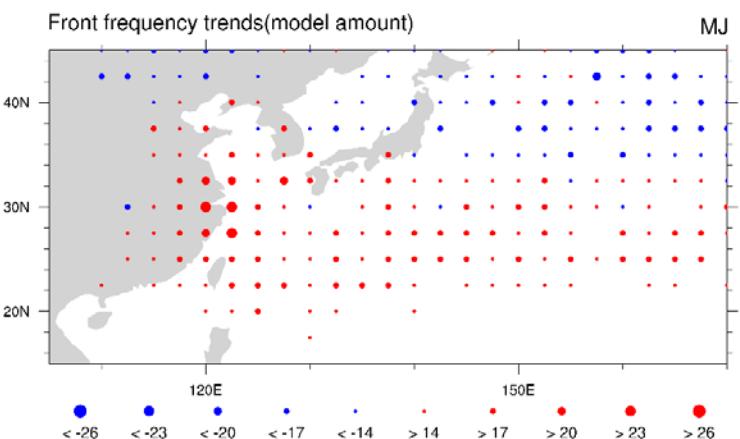
Front frequency trends(model amount)



Front frequency trends(model amount)

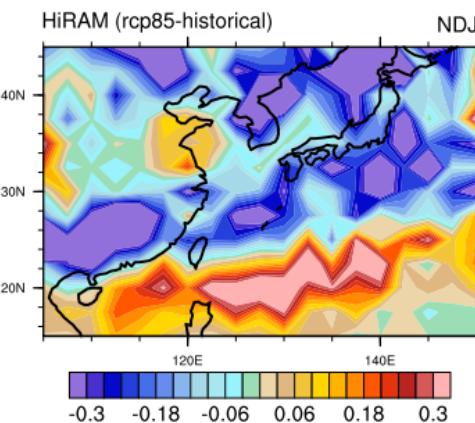


Front frequency trends(model amount)

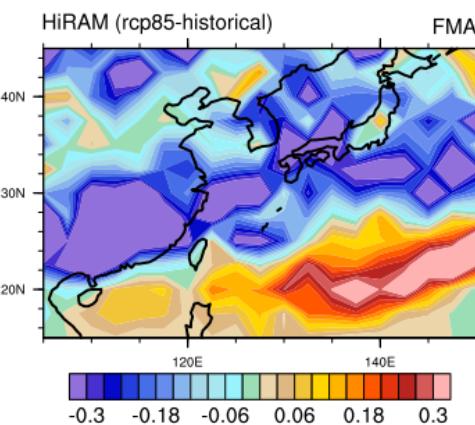


CMIP5 (RCP85-historical)

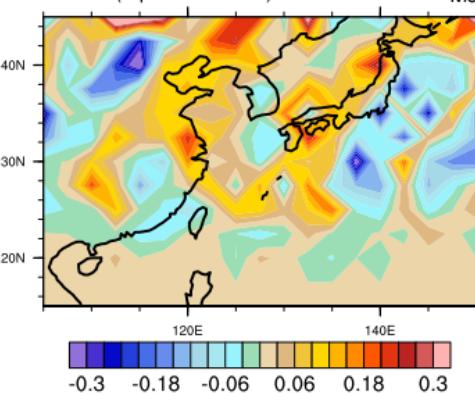
HiRAM(RCP8.5-historical)



HiRAM (rcp85-historical)

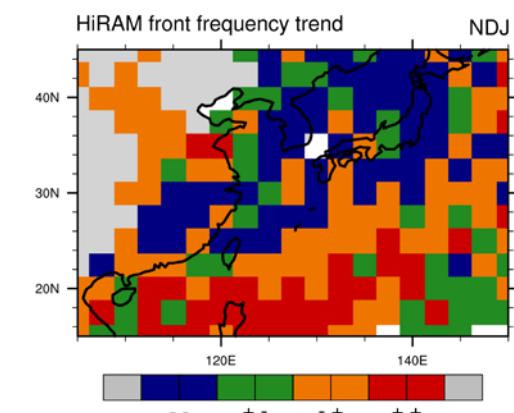


HiRAM (rcp85-historical)

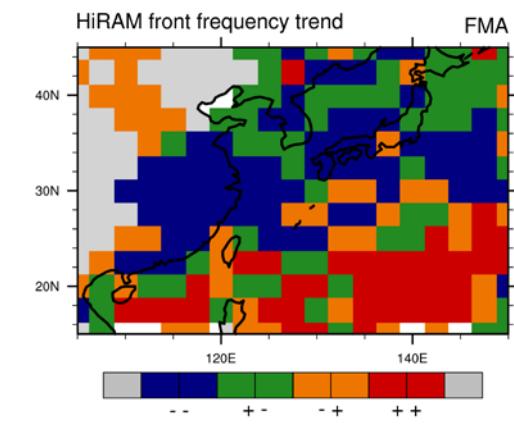


(NF-historical) (RCP8.5-NF)

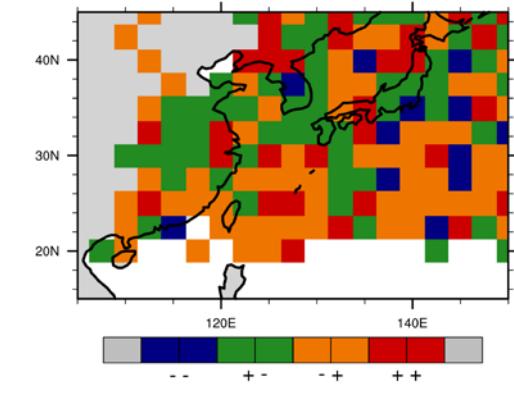
HiRAM front frequency trend



HiRAM front frequency trend



HiRAM front frequency trend

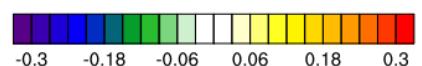
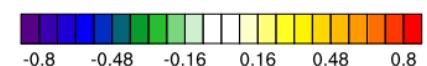
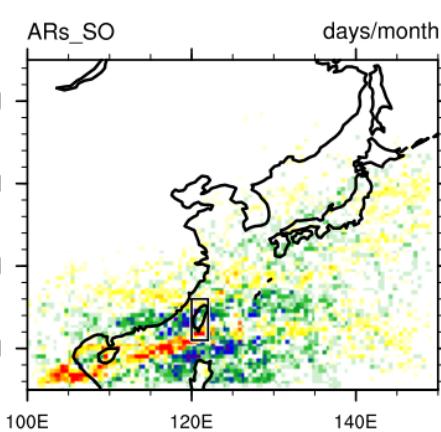
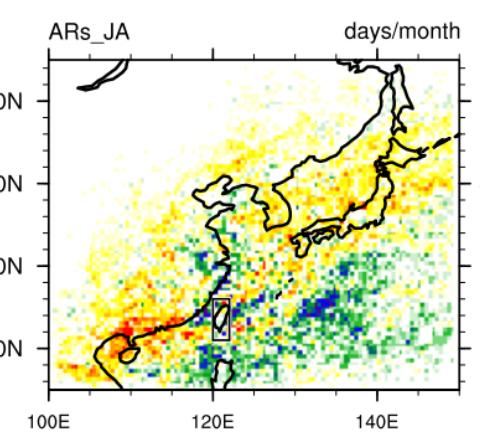
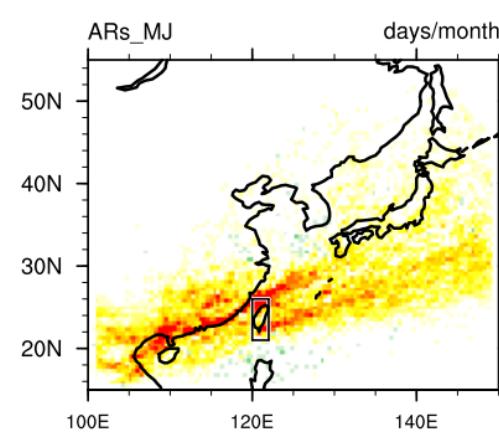
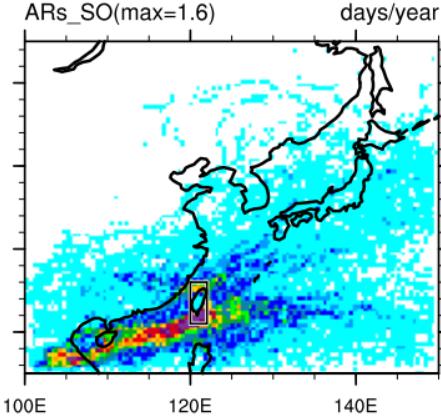
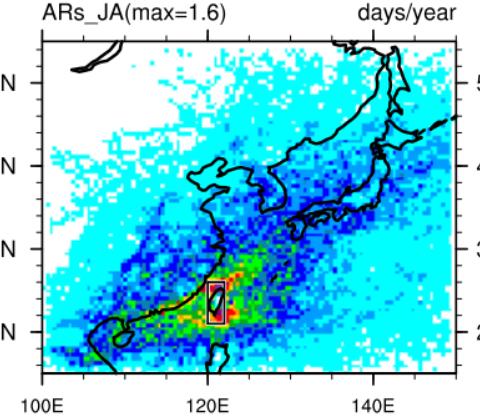
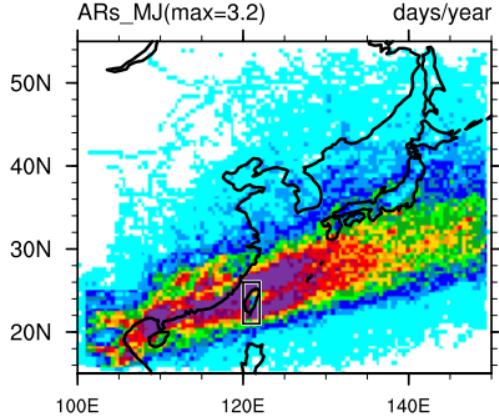
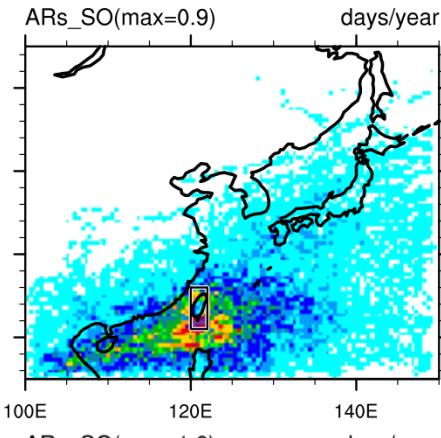
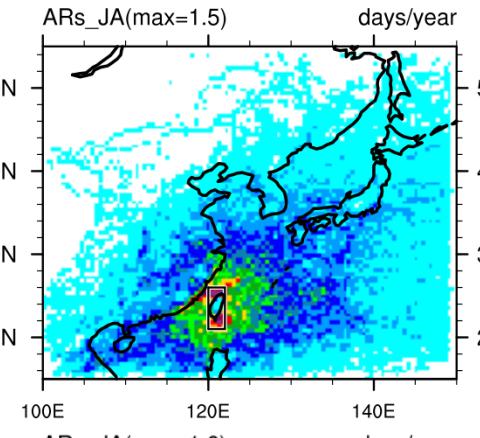
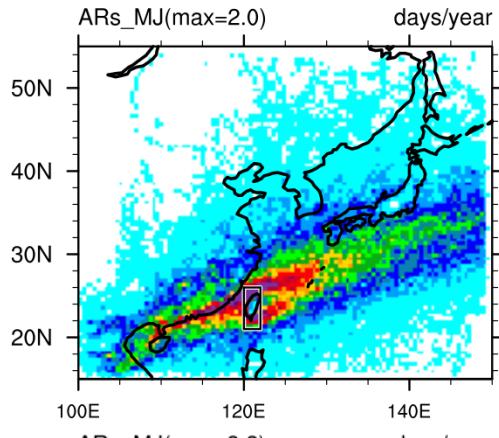


**Atmospheric
River
Frequency
(passed Taiwan)**

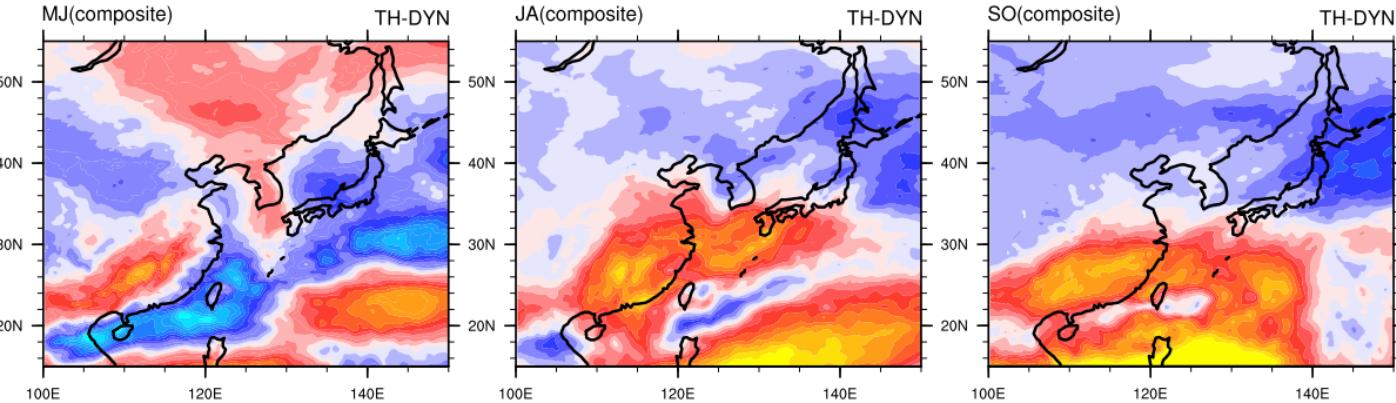
HiRAM
historical

HiRAM
rcp85

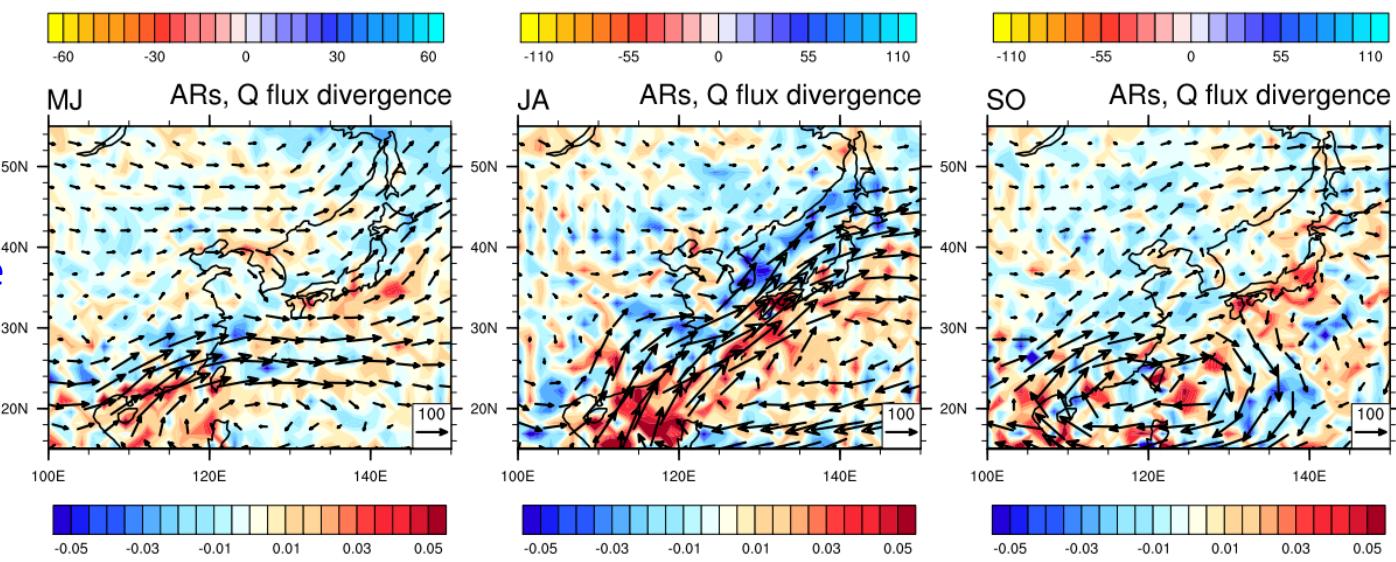
Rcp85-historical



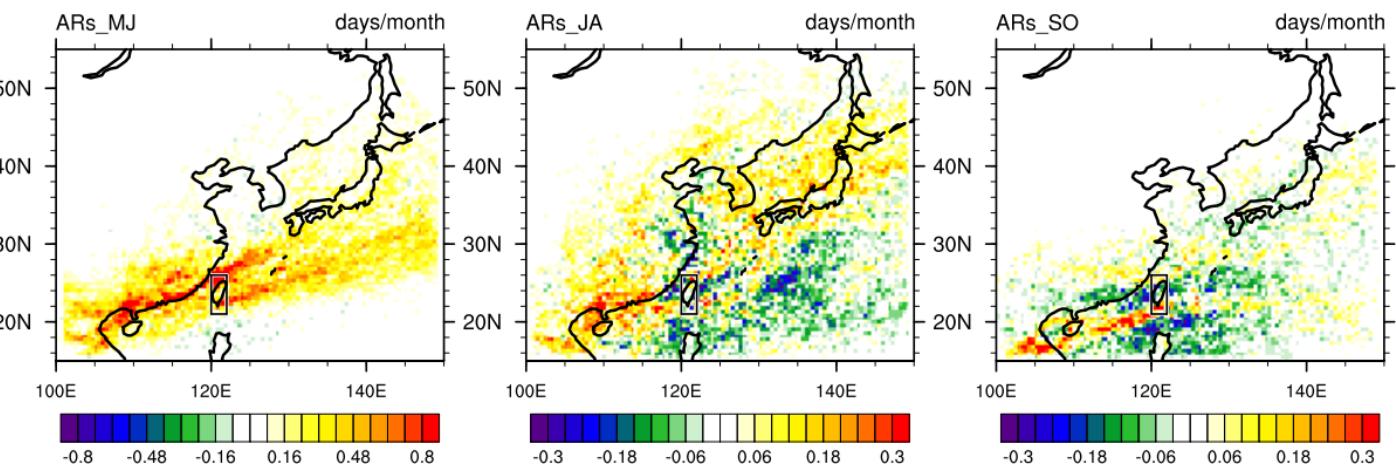
thermodynamic
- dynamic response



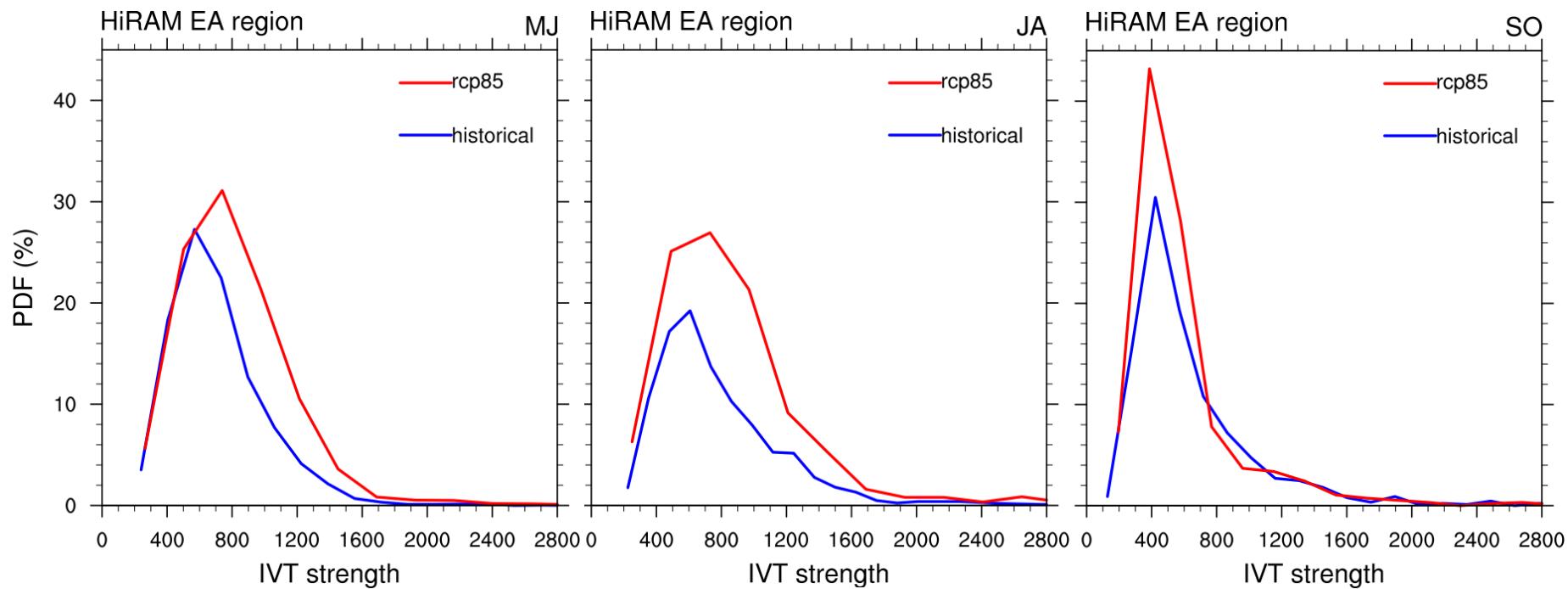
Moisture flux&divergence
Composite difference
(rcp85 minus historical)



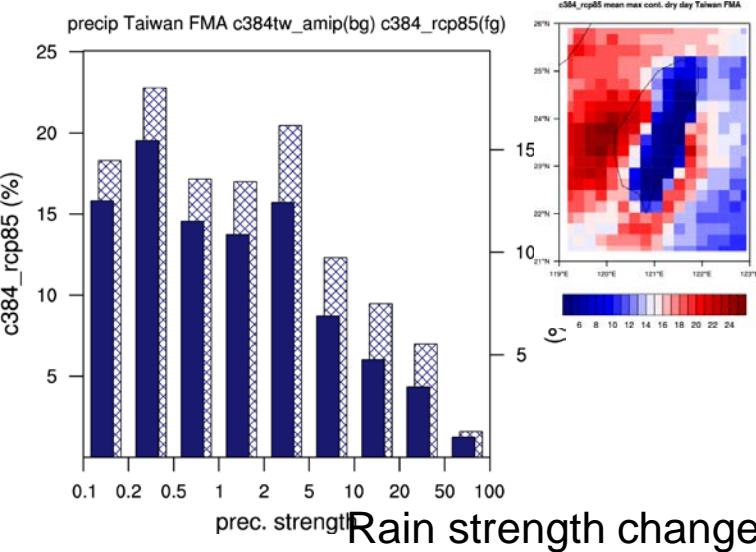
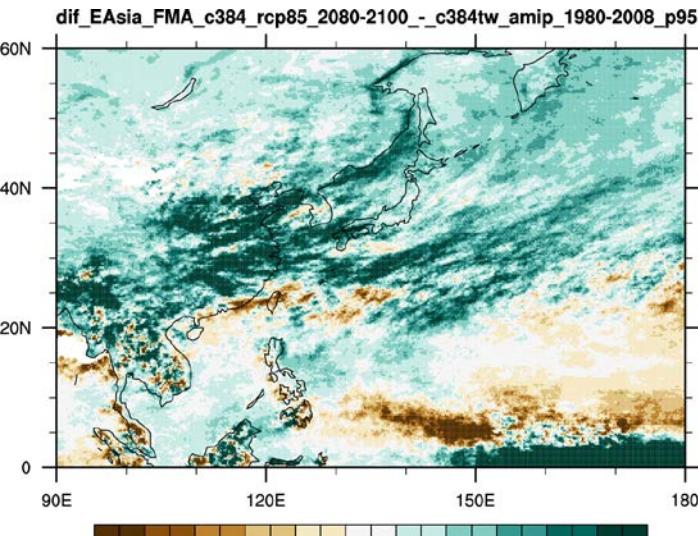
AR frequency
Rcp85-historical



IVT strength probability density function



Spring Rainfall in Warmer climate



Mean yearly max
continuous dry day
change

